MARYLAND GEOLOGICAL SURVEY



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CONVERSION FACTORS AND ABBREVIATIONS

For those readers who may prefer to use metric (International System) units rather than the inch-pound units used in this report, values may be converted using the following factors:

Multiply inch-pound unit	Ву	To obtain metric unit
foot (ft)	0.3048	meter (m)
foot per mile (ft/mi)	0.1894	meter per kilometer (m/km)
ineh (in.)	25.4	millimeter (mm)
mile (mi)	1.609	kilometer (km)
square mile (mi ²)	2.590	square kilometer (km²)
gallon (gal)	3.785	liter (L)
gallon per minute per foot [(gal/min)/ft]	0.2070	liter per second per meter [(L/s)/m]
million gallons (Mgal)	3,785	cubie meter (m³)
foot per day (ft/d)	0.3048	meter per day (m/d)
foot squared per day (ft²/d)	0.09290	meter squared per day (m²/d)
eubic foot per day (ft³/d)	0.02832	cubic meter per day (m³/d)
gallon per day (gal/d)	0.003785	cubic meter per day (m³/d)
gallon per minute (gal/min)	0.06309	liter per second (L/s)
million gallons per day (Mgal/d)	0.04381	eubic meter per second (m³/s)

In this report, chemical concentration in water is expressed in milligrams per liter (mg/L) or micrograms per liter (μ g/L); 1,000 μ g/L = 1 mg/L. The concentration unit of milliequivalents per liter (meq/L) takes into aeount the ionic charge and combining weight of an ion. In an analysis expressed in meq/L, unit concentrations of all ions are chemically equivalent. (See Hem, 1985, p. 56 for a further explanation and comprehensive listing of conversion factors.) Water temperature in degrees Celsius (°C) can be converted to degrees Fahrenheit (°F) using the following equation:

$$^{\circ}F = 1.8 (^{\circ}C) + 32$$

Specific electrical conductance of water is expressed in microsiemens per centimeter at 25 °C (μ S/cm). This unit is equivalent to micromhos per centimeter at 25 °C.

Sea level: In this report "sea level" refers to the National Geodetic Vertical Datum of 1929 (NGVD of 1929) — a geodetic datum derived from a general adjustment of the first-order level nets of both the United States and Canada, formerly called "Sea Level Datum of 1929."

Department of Natural Resources

MARYLAND GEOLOGICAL SURVEY

Kenneth N. Weaver, Director

BULLETIN 35

HYDROGEOLOGY AND GROUND-WATER RESOURCES OF SOMERSET COUNTY, MARYLAND

by William H. Werkheiser U.S. Geological Survey



Prepared in cooperation with the
United States Department of the Interior
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1990

COMMISSION OF THE MARYLAND GEOLOGICAL SURVEY

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HYDROGEOLOGY AND GROUND-WATER RESOURCES OF SOMERSET COUNTY, MARYLAND

by

William H. Werkheiser

ABSTRACT

Somerset County, Maryland, an area of about 597 square miles, relics on ground water for 84 percent of its water supply. Development in the county is expected to substantially increase the demand for water; hence, an assessment of the ground-water resources was conducted to collect baseline information against which future change can be measured. Specific goals were to (1) refine the understanding of the hydrogeologic framework; (2) describe the quality of ground water; and (3) evaluate the effects on the ground-water-flow system of projected ground-water withdrawals at Princess Anne and Crisfield.

Somerset County is in the Atlantic Coastal Plain physiographic province, and is underlain by a wedge of unconsolidated sediments that forms a series of aquifers and confining units. The aquifers and aquifer systems that supply water are (1) the surficial aquifer system; (2) the Pocomoke aquifer; (3) the Manokin aquifer; (4) the Choptank aquifer; (5) the Piney Point aquifer; (6) the Palcocene aquifer system; and (7) the Potomac aquifer system.

The surficial aquifer system consists of fine-grained sediments, is relatively thin throughout much of the county, and is used primarily for domestic water supply. The aquifer system may produce more water in the northeastern part of the county, where it is coarser-grained and thicker. Chemical analyses of four samples suggest the water is soft to moderately hard and slightly acidic. In areas containing anoxic water, dissolved iron may be present in elevated concentrations. In areas of oxygenated water, nitrate concentrations may be clevated where water comes into contact with nitrogen sources.

The Pocomoke aquifer is present only in the southeastern part of the county. The median reported specific capacity of 68 wells tapping the unit is 10 gallons per minute per foot [(gal/min)/ft]. Specific capacity from five 1-hour tests performed during the investigation ranges from 2.0 to 17.3 (gal/min)/ft. The most common water-quality problems associated with the Pocomoke aquifer are elevated concentrations of iron and manganese. All but 2 of 24 samples exceeded the U.S. Environmental Protection Agency (USEPA) secondary maximum contaminant level (SMCL) for iron and 18 of 24 exceeded the SMCL for manganese.

The principal aquifer of use in Somerset County is the Manokin aquifer. The median reported specific capacity of wells in the unit is 5.2 (gal/min)/ft. Hydraulic conductivity, calculated from five aquifer tests near Princess Anne, averages 13.2 feet per day (ft/d), and storage coefficients range from 0.0002 to 0.001. Throughout much of the county, water-level altitudes in the aquifer are below sea level, with the lowest near Princess Anne. Because ground-water flow is chiefly toward the pumping centers at Princess Anne, poor-quality water from the southern part of the county, or the Chesapeake Bay, could migrate in that direction.

There is marked areal variation in the quality of water in the Manokin aquifer. North of Westover, water is soft to moderately hard, relatively low in dissolved solids, and of the sodium bicarbonate type. South of Westover, water is moderately hard to hard, contains higher concentrations of dissolved solids, and is of the sodium chloride type. In the north-castern corner of the county, iron concentrations generally exceed the USEPA SMCL of 0.3

milligrams per liter (mg/L). In the southern part of the eounty, chloride concentrations generally exceed the USEPA SMCL of 250 mg/L. The 250-mg/L isochlor trends roughly from Pocomoke City to Deal Island and passes about 1 mile (mi) north of Westover. The 500-mg/L isochlor is located about 2 mi southwest of the 250-mg/L isochlor.

Few wells produce water from the Choptank or Piney Point aquifers. Water from the Choptank aquifer is reported to contain chloride concentrations in excess of 900 mg/L, and water from the Piney Point aquifer contains dissolved solids in excess of 1,000 mg/L.

The Paleoeene aquifer system is used only by the town of Crisfield as a source of water supply. Reported specific capacity of two wells screened entirely in the aquifer system is nearly 2 (gal/min)/ft. The Potomae aquifer system supplies water to several municipalities along the Chesapeake Bay. Reported specific capacities of four wells range from 1 to 7 (gal/min)/ft. Transmissivity and storage coefficient, estimated from a multiple-well aquifer test, are 2,140 ft²/d (feet squared per day) and 0.0002, respectively. Transmissivity, estimated from a single-well aquifer test, is 1,280 ft²/d.

Water from the Palcocene and Potomac aquifer systems is soft, has concentrations of dissolved solids ranging from 475 to 1,070 mg/L, and has the highest pH of any ground water in the eounty. The water is of the sodium-bicarbonate type, with sodium eomprising more than 95 percent of the cations. Seven of 10 water-quality samples eontained fluoride eon-centrations above the SMCL and two of these exceeded the USEPA primary maximum eon-taminant level (MCL). The extent and water quality of the two aquifer systems east of Crisfield are not known.

A digital, steady-state, ground-water-flow model was used to evaluate the effects of projected increases in pumpage of 600,000 gallons per day (gal/d) from the Manokin aquifer near Princess Anne. Simulated water levels in the Manokin aquifer ranged from 15 to 70 feet (ft) below those measured in November 1986. Using model-derived ground-water velocities, it would take about 50 years for ground water to move from the vicinity of the 250-mg/L isochlor to the nearest simulated pumped well at Princess Anne. A travel time of 300 years was estimated for the distance from the 500-mg/L isochlor to the same well. The analysis did not include the effects of dispersion, which could hasten the first arrival of brackish water (chloride concentration greater than 250 mg/L).

A non-steady analytical solution was used to estimate additional water-level declines that may result from projected pumpage from the Paleocene and Potomac aquifer systems in the Crisfield area. Using a range of transmissivity and a storage eoefficient from aquifer tests, additional water-level declines were estimated to range from 7 to 31 ft. It is not known if poor-quality water is migrating toward the pumping centers, because the nature of the aquifer systems east of Crisfield is unknown.

INTRODUCTION

Somerset County, a predominantly rural area located on the Eastern Shore of Maryland, depends on ground water for about 84 percent of its water supply. Although demand for ground water in the county is expected to increase because of population growth, the effects of increased withdrawals are not known. Of particular concern are the Princess Anne area, where a major State facility (Eastern Correctional Institution) recently has been constructed, and the harbor area of Crisfield. An assessment of the ground-water resources and the effects of anticipated ground-water withdrawals was conducted to aid in the long-term management of the resource. The study was done by the U.S. Geological Survey in cooperation with the Maryland Geological Survey and Somerset County.

PURPOSE AND SCOPE

This report presents the findings of the assessment; specifically, the report (1) refines the description of the hydrogeologic framework of Somerset County; (2) describes the quality of ground water; and (3) describes the effects on the ground-water-flow system of projected increases in ground-water pumpage at Princess Anne and Crisfield.

Lithologie logs from well-completion records and published reports, along with geophysical logs, were analyzed for information on the hydrogeologic framework. Water levels measured in 65 wells from April 1986 through September 1987 were used to determine potentiometric surfaces of the aquifers. Eighty-one water-quality samples, collected from April 1986 through September 1987, were analyzed to assess ground-water quality. Ground-water-flow models were applied, to two sites, to evaluate the effects of proposed pumpage increases on the ground-water-flow system.

DESCRIPTION OF STUDY AREA

Somerset County occupies part of the Atlantic Coastal Plain physiographic province which extends from Long Island, N.Y., to the Gulf of Mexico. It is located in the southwestern part of the Eastern Shore of Maryland (fig. 1). The land area of the county is about 597 mi² (square miles). Of this total area, 46 percent is covered by water, 21 percent is forested, 16 percent is used for agriculture, 16 percent is wetland, and 1 percent is developed (J. Windsor, Somerset County Department of Technical and Community Services, oral commun., 1988). Topographically, Somerset County is essentially a low-lying plain, with altitudes generally less than 20 ft (feet) above sea level. In the northeastern corner near the Somerset County-Wicomico County line, altitudes are somewhat greater, reaching 50 ft above sea level.

Population of the county was 19,188 in 1980 (U.S. Department of Commerce, 1981), with the greatest concentration along the highway corridor between Salisbury, Md. and Pocomoke City, Md., and in the harbor area of Crisfield.

Figure 2 shows the annual precipitation for the period of record at Princess Anne. The average annual precipitation for the periods 1949-57, 1959-64, and 1966-84 is 44.44 in. (inches) (National Oceanie and Atmospherie Administration, 1984). Data were incomplete for 1958 and 1965, and are not shown on figure 2. Precipitation is fairly uniform throughout the year, although there tends to be more rainfall during July, August, and September. The driest year was 1966, with slightly more than 30 in. of precipitation; the wettest year was 1979 when more than 63 in. of precipitation fell.

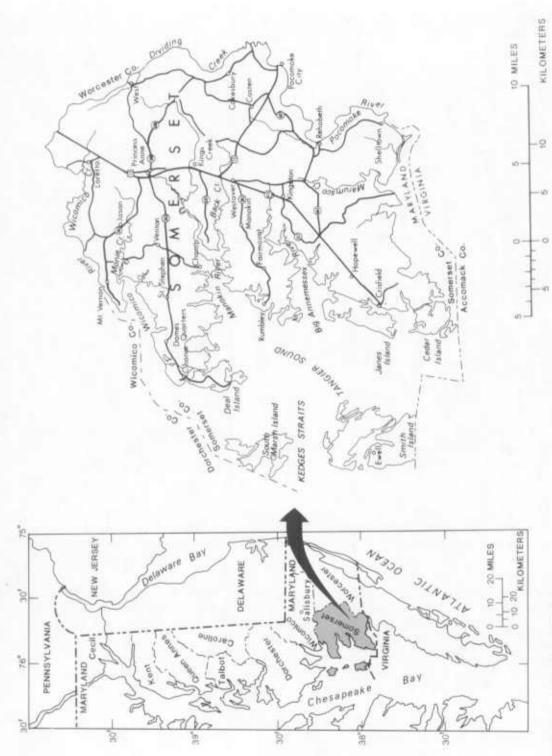


Figure 1.- Location of study area.

INTRODUCTION

5

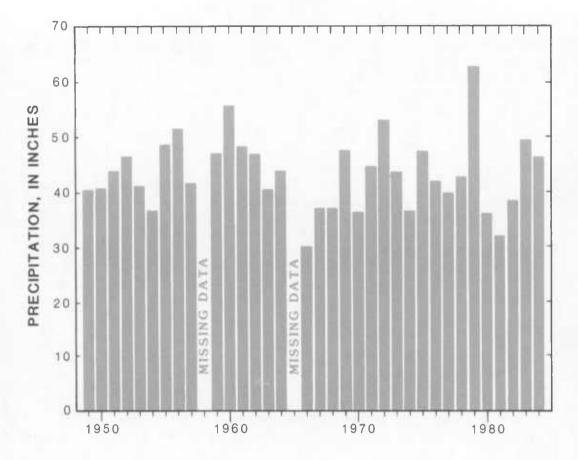


Figure 2.— Precipitation at Princess Anne, 1949-84.

PREVIOUS INVESTIGATIONS

Rasmussen and Slaughter (1955) described the hydrogeologic framework and ground-water resources of Somerset County as part of an investigation of the water resources of Somerset, Wicomico, and Worcester Counties. Hansen (1967) presents hydrogeologic data from a test well drilled into the Cretaceous sediments at Janes Island State Park. The hydrogeology and stratigraphy of a 1,515 foot well drilled at the Eastern Correctional Institution in 1989 is described by Hansen and Wilson (1990). Several studies have examined the ground-water resources of Somerset County as part of regional hydrogeologic investigations. Cushing and others (1973) describe the water resources of the Delmarva Peninsula. Bachman and Wilson (1984) included the northeastern part of Somerset County in their study of the surficial Columbia aquifer of the Eastern Shore of Maryland. A study of regional ground-water flow in the Atlantic Coastal Plain of Maryland, Delaware, and the District of Columbia, includes Somerset County (D.A. Vroblesky and W.B. Fleck, U.S. Geological Survey, written commun., 1986). Water-quality data from Somerset County wells are included in Knobel (1985).

The geology of Somerset County was mapped by Owens and Denny (1984). Regional geologie investigations that include the county were conducted by Owens and Denny (1979), Brown and others (1972), and Mixon (1985). Hansen (1981) discussed the relationship of the Columbia Group to the underlying Miocene aquifer complex; and Hansen (1978) investigated Upper Cretaceous and Paleocene pinchouts in the Maryland part of the Salisbury Embayment.

WELL-IDENTIFICATION SYSTEM

The locations of wells described in this report are shown on 5-minute quadrangle maps at the end of the report. An index map of the 7 1/2-minute quadrangle maps used to derive these maps is given in figure 41 at the end of the report. Wells are identified in this report in accordance with the Maryland Geological Survey numbering system. Each identifier consists of two pairs of letters followed by a number (for example, SO Bf 19). The first pair of letters indicates the county (SO for Somerset); the second pair designates one of the 5-minute quadrangles of latitude and longitude into which each county has been subdivided (fig. 42 at end of report). The number identifies a specific well within a 5-minute quadrangle.

ACKNOWLEDGMENTS

The many homeowners, farmers, and businesses who allowed access to their wells for water-level measurements and collection of water-quality samples are gratefully acknowledged. Valuable information on the location and characteristics of wells in Somerset County was provided by Philip Pryor and Linda Collier, of the Somerset County Health Department. Staff of the Maryland Water Resources Administration are acknowledged for responding to various information requests and for installing two piezometers. Dr. Harry J. Hansen, of the Maryland Geological Survey, reviewed and edited the hydrogeologic sections contained in this report. His suggestions are much appreciated.

HYDROGEOLOGY AND GROUND-WATER RESOURCES

GENERAL DESCRIPTION

Somerset County is underlain by a wedge of generally uneonsolidated sand, silt, and clay strata that dip to the southeast. This wedge is about 4,225 ft thick at Crisfield, Md., and thickens to the east at a rate of about 90 ft/mi (feet per mile) (Hansen, 1978, p. 4). In describing the water-bearing properties of these sediments, the terms aquifer and aquifer system are used. An aquifer is a body of earth material capable of yielding significant quantities of water to wells and springs. An aquifer system is a heterogeneous body of material that comprises two or more aquifers separated, at least locally, by confining units that impede ground-water movement but do not greatly affect the regional hydraulic continuity of the system (Laney and Davidson, 1986, p. 6-7). The aquifers and aquifer systems present in the county are: the surficial aquifer system, the Pocomoke aquifer, the Manokin aquifer, the Choptank aquifer, the Piney Point aquifer, the Paleocene aquifer system, and the Potomac aquifer system.

Table 1 summarizes the relation between geologic units, based on the lithology and paleontology of the sediments; and hydrogeologie units, based on the water-bearing properties of the sediments. The stratigraphy presented in this report is a compilation from previous geologie investigations. The nomenclature of Owens and Denny (1984) is used for the Holocene and Pleistocene units, and the Beaverdam Sand of the Pliocene Series. The designation "Yorktown and Eastover Formations (undifferentiated)" is used rather than the designation "Yorktown and Cohansey Formations" (Rasmussen and Slaughter, 1955, p. 43) because the age equivalency of the Yorktown Formation and the Cohansey Formation has been seriously questioned (Owens and Denny, 1979, p. A7; Hansen, 1981, p. 128; and Mixon, 1985, p. G14). The nomenclature of Rasmussen and Slaughter (1955) is used for the remainder of the Miocene units (St. Marys Formation, Choptank Formation, Calvert Formation). The nomenclature of Hansen (1978) is used for the Oligocene through Cretaceous units, except for the Paleocene units, which are not differentiated in this report. The hydrogeologic unit term "Paleocene aquifer system" follows the usage of the Maryland Geological Survey.

Plates 1-4 are hydrogeologic sections through Somerset County. Many of the geophysical logs shown (SO Bg 3, SO Ce 6, SO De 3, SO De 27, SO De 28, SO Ea 11, and WI Dd 23) are published in other reports (Hansen, 1967, fig. 13; Boggess and Heidel, 1968, p. 9; Hansen, 1978, fig. 10; and Hansen, 1981, fig. 2). Stratigraphic interpretations are based on these logs, and correlations are extended to other logs on the sections.

Plate 1 is a south-north hydrogeologic section that roughly parallels the strike of the units; plates 2-4 are west-east hydrogeologic sections that roughly parallel the dip of the aquifers. Plate 2 is the northernmost section along the dip and plate 4 is the southernmost. The surficial aquifer system is about 20 ft thick and directly overlies the Poeomoke aquifer along much of hydrogeologic section A-A' (pl. 1). The Pocomoke aquifer is absent in the northwestern part of the county where it is truncated by the surficial aquifer system (pls. 2-4). Downdip, the surficial aquifer system and the Pocomoke aquifer are separated by a confining unit of silt, clay, and fine sand that thickens to the southeast.

The Manokin aquifer underlies the surficial aquifer system and the Pocomoke aquifer, and is separated from them by a confining unit that appears to be thick and laterally continuous. The Manokin aquifer is a distinct unit along hydrogeologic section B-B' (pl. 2) and in the northern part of hydrogeologic section A-A' (pl. 1), but is poorly developed along much of hydrogeologic section C-C' (pl. 3), in the western part of D-D' (pl. 4), and in the southern part of A-A' (pl. 1). At some locations, such as at SO Bg 3 (pl. 2), gamma logs indicate that the Manokin aquifer grades into the underlying confining unit; whereas at other places (SO Ce 42, pl. 1), the contact between the two units is more sharply defined. Where the Manokin aquifer is poorly developed (SO De 27, pl. 3), it is difficult to distinguish the Manokin from overlying and underlying confining units.

The Manokin aquifer is separated from the underlying Choptank aquifer by a thick confining unit (St. Marys Formation). The Choptank aquifer is present at all locations where wells have been drilled to that depth (pls. 1-4); and it appears that the Choptank aquifer is present throughout Somerset County. Another thick confining unit separates the Choptank aquifer and the Piney Point aquifer (pls. 1, 3-4). The Piney Point aquifer appears to be laterally continuous in the southern part of the county. No wells have been drilled deep enough in the northern part of the county to determine the characteristics of the aquifer there.

The Paleoeene aquifer system lies below the Piney Point aquifer. At Janes Island State Park and at Fairmount, the aquifer system appears to consist of three aquifers separated by confining units (pl. 1). Between Smith Island and Crisfield, however, the relation is more

TABLE 1
GEOLOGIC AND HYDROGEOLOGIC UNITS IN SOMERSET COUNTY
[correlation of units shown in each column under previous investigations not implied]

				Previo	ue investigati	one			Thie report		
System	Sariaa		Geologic v	nita				Hydrogaologic unite			
		Sla	muaean nd ughter 955)	Hensen (1978)	Hansan (1981)	Owene and Danny (1984)	Mixon (1985) aoutharn Somereat County	Banaen (1967) Criefield eree	Geologic unite	Hydrogaologic unite	
	Holocene	Undiff	arentieted	Not atudied		Alluvium, awamp, and tidel mereh dapoeits	Estuarina mereh depoeite		Alluvium, ewamp, and tidal marah deposite		
Queternary	Plaiatocana	Talbo	eonaburg Sand t and Pamlico ormetione eton Silt	Not atudiad	Shoralina Complax	Personsburg Sand Kent Island	Sand <u>alluvium</u> Kent point-ber	Undifferen- tieted	Pereonaburg Sand Kent Teland	Surficial aquifar eyetem	
		Beaverdem Sand			Pageworks.	Formetion Omer Formation	Kent Island Formetion		Formetion Omer Formation		
		Beacon H	, Bryn Mawr, and ill grevels velly Sand)	Not	Send	Beaverdam Send			Beevardam Sand	??	
Tertiery	Pliocene	(Red Oreverly Smid)	etudiad	"Red Gravelly Send" Casses		Yorktown Formation	Not praeant	Yorktown Formetion end	Confining unit Pocomoka equifar		
		Yorktown and Cohaneay Formationa (?)	Upper equicluda Pocomoke equifar Lower equiclude Manokin equifer		St. Marye (7)	Yorktown (?) and Cohaneay (?) Formetione	Eastovar Formation	Confining unit Pocomoka aquifar Confining unit Manokin equifer	Eastover Formation (undiffaren- tisted) St.Marya Formation	Miocene aeriae unit Manok undiffaran-Confin tiated	
	Miocene	St. Marye Formetion Choptank Formation Celvert Formation	Undifferen- tisted	Not atudiad	Not atudied	Not atudiad	Confining unit Choptenk aquifar Confining unit	Choptenk Formation	Choptank aquifar Confining		
	Oligocene	Not prese		Not present	Not etudied	Not etudied	Not	Not present	Formation Not present	Hot present	
	Eocena	Chickeho Piney Po	int Formation y Formation	Pinay Point Formation	Not atudiad	Not atudied	Not etudied	Piney Point aquifar	Pinay Point Formation	Piney Point aquifer	
	Palaocana		Greeneand (?) Formetion	Aquia Formation Brighteest Formetion	Not etudied	Not etudied	Not etudied	Confining unit "Peleccena" equifer	Undifferentieted	Confining unit Palaccana equifor ayatem	
	Upper		Formation Formation	Not present	Not atudied	Not atudied	Not etudied	Confining unit Confining unit Megothy aquifer Confining unit "Upper Reriten" aquifer	Not present	Not present	
Cratacaoua	Cretaceoue	Raritan Patapaco	Formation Formation and Arundal	Potomac Group (undiffaren- tiated)	Not atudied	Not studied	Not atudied	Confining unit Not studied Not studied	Potomac Group (undiffaran- tiated)	Potomac aquifar eyetam	
	Lower Crataceous	Petuxent	ations Formation		atudiad	atudiad	atudied	atudiad	tiated)		

complex, with the position and number of aquifers changing from Smith Island to Crisfield (pl. 4). Therefore, the vertical or lateral distribution of aquifers and confining units in the Paleocene aquifer system is difficult to predict. The contact between the Paleocene units and the Potomac Group was determined by Hansen (1978, fig. 10), based on paleontological evidence and geophysical logs from well SO Dc 3. That determination was used in this report and correlation was made from well SO Dc 3 to other wells on the basis of geophysical logs only. Therefore, the position of the contact between the Paleocene aquifer system and the Potomac aquifer system is uncertain. Also, in places the Potomac aquifer system and the Paleocene aquifer system may be hydraulically connected (pl. 4). There are limited data on the characteristics of the Potomac aquifer system (pls. 1, 3-4). However, because the Potomac Group is of nonmarine origin, it is likely that the aquifer system consists of discontinuous sand bodies of variable vertical and horizontal extent, separated by confining units of fine sand, silt, and clay. The characteristics of the various aquifers and aquifer systems are more fully explained in subsequent sections of this report.

GROUND-WATER FLOW

Water enters the ground-water-flow system primarily by the infiltration of precipitation. Although water budgets were not developed for Somerset County, other studies on the Delmarva Peninsula indicate that on an annual basis, between 37 and 55 percent of the precipitation that falls eventually enters the ground-water system (Rasmussen and Andreasen, 1959; Johnston, 1976). Figure 3 schematically shows typical flow paths of water in the subsurface of Somerset County, Most water that infiltrates to unconfined aquifers (surficial aquifer system, and parts of the Pocomoke aquifer) discharges to nearby surface-water bodies or is utilized by plants and transpired to the atmosphere. Some water, however, moves downward from the unconfined aquifers to deeper, confined aquifers. Water also enters the eonfined aquifers of Somerset County by lateral flow from areas outside the county. Water in these deeper aquifers is discharged primarily by vertical leakage through confining units to other aquifers and by pumping. In some areas, vertical leakage into an aquifer is increased by ground-water withdrawals. The rate of ground-water movement through a confining unit separating two aquifers depends on the head difference between the two aquifers and the vertical hydraulic conductivity and thickness of the confining unit. If the head difference increases because water levels in an aquifer have declined in response to pumping, the rate of flow through the confining unit increases.

GROUND-WATER AVAILABILITY

The quantity of water an aquifer theoretically can yield to wells depends primarily on the hydraulic characteristics of the aquifer. Information on reported well yields alone usually does not adequately characterize the water-producing capacity of an aquifer. Yields given in well-completion reports depend on the method of well construction, type of pump, the yield required, duration of pumping period, and the drawdown allowed during the yield test. A better means for comparing the water-producing characteristics of wells is specific capacity, which is the well discharge divided by the water-level drawdown in the well at the end of the yield test. Specific capacity more accurately reflects the hydraulic characteristics of the aquifer, but also is strongly affected by well construction.

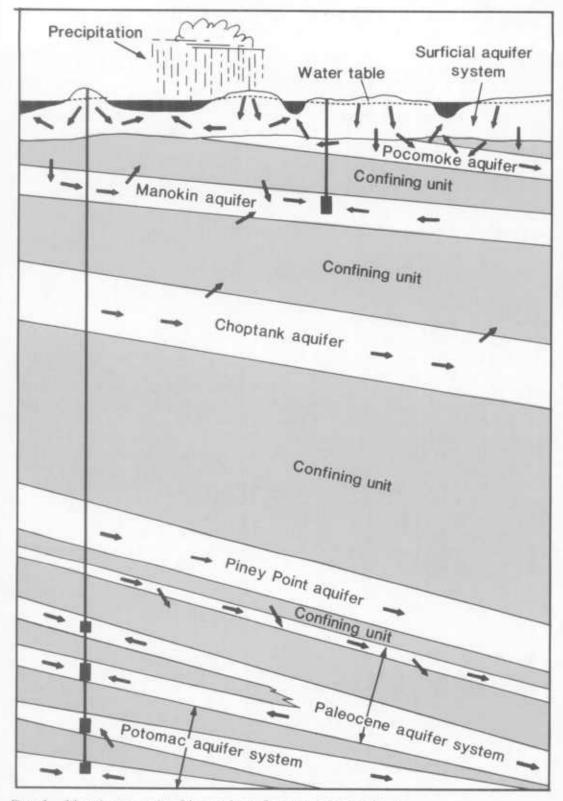


Figure 3.— Schematic representation of the ground-water-flow system in Somerset County.

The two terms used most often to describe the hydraulic characteristics of a confined aquifer are transmissivity and storage coefficient. Transmissivity is defined as the rate that water is transmitted through a unit width of aquifer under a unit hydraulic gradient (Lohman, 1972, p. 6). Transmissivity in a homogeneous, isotropic aquifer is related to the hydraulic conductivity and thickness of the aquifer by the equation:

$$T = Kb. (1)$$

where:

T = transmissivity, in feet squared per day;

K = hydraulic conductivity, in feet per day; and

b = thickness of the aquifer, in feet.

The storage coefficient of an aquifer is defined as the volume of water the aquifer releases from or takes into storage per unit surface area per unit change in head (Lohman, 1972, p. 8). In an unconfined aquifer, the storage coefficient is approximately equal to the specific yield, which is defined as the ratio of the volume of water an aquifer will yield by gravity drainage to the total volume of the aquifer (Lohman, 1972, p. 6).

The transmissivity and storage properties of an aquifer are most frequently evaluated by conducting controlled tests that involve pumping a well for a specific period of time and measuring changes in water levels at nearby observation wells. In Somersct County, however, very few aquifer tests have been performed. More commonly, a specific-capacity test is conducted when a production well is installed. Transmissivity of a confined aquifer can be estimated from reliable specific-capacity data by using the following equation derived from the equation in Driscoll (1986, p. 1,021), which assumes a 6-in.-diameter well and a 24-hour test:

$$T = Q/s \times 270,$$
 (2)

where:

T = transmissivity, in feet squared per day; and

Q/s = specific capacity, in gallons per minute per foot of drawdown.

For unconfined aquifers the approximation is:

$$T = Q/s \times 200.$$
 (3)

Except where noted, the specific-capacity values in this section were obtained from well-completion reports. Reported values may differ considerably because of:

- (1) Changes in hydraulic properties of the aquifer;
- (2) differences in well construction:
- (3) differences in the thickness of aquifer penetrated;
- (4) inaccuracies in reported discharge;
- (5) inaccuracies in reported drawdown; and
- (6) differences in duration of the specific-capacity tests.

WATER USE

Total water use in Somerset County has tripled since 1950, when estimated usage was 1.56 Mgal/d (million gallons per day). In 1986, water use was estimated to be 4.85 Mgal/d

(J.C. Wheeler, U.S. Geological Survey, written commun., 1988). Figure 4 illustrates water use by category, and shows that the largest amount of water used in Somerset County is for public supply, followed by irrigation, livestock, and commercial and industrial uses. Self-supplied, domestic water use accounted for only 12 percent of all water used in 1986. This contrasts with the distribution in 1950, when 34 percent of water use in the county was for self-supplied, domestic purposes (Rasmussen and Slaughter, 1955, p. 135).

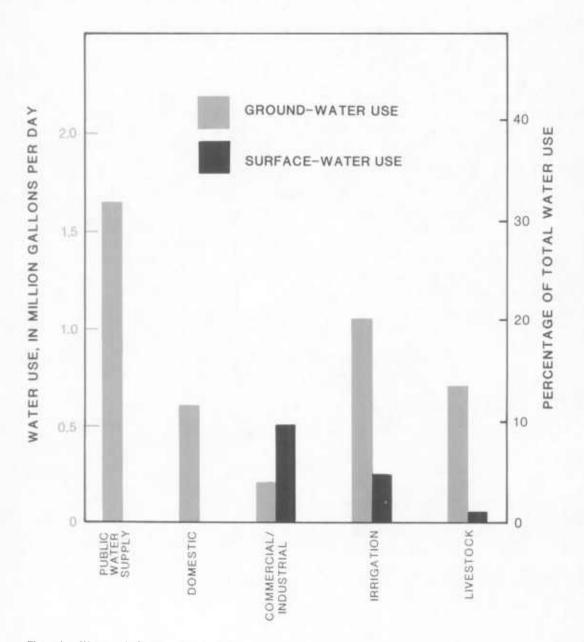


Figure 4.— Water use in Somerset County, 1986.

Figure 4 also shows that, in 1986, 84 percent of water use in Somerset County was from ground-water sources. Virtually all of the potable water in the county is ground water because most streams and rivers are tidal and contain brackish water. Figure 5 shows the locations of the public water-supply wells in the county, and the amounts of ground water each system produces. The largest municipal user of ground water is Crisfield, followed by Princess Anne, Smith Island, Fairmount, and a mobile home park near Eden. The public system that uses the least amount of ground water is the Rumbley-Frenchtown system.

Crisfield withdraws water from the Paleocene and Potomae aquifer systems, whereas Smith Island, Fairmount, and Rumbley-Frenchtown withdraw water from the Potomae aquifer system only. Princess Anne withdraws water from the Manokin aquifer, and the system near Eden withdraws water from both the surficial aquifer system and the Manokin aquifer. Table 2 lists users of more than 10,000 gal/d (gallons per day). In addition to municipal users, a plywood-manufacturing plant, and a poultry business also use more than 10,000 gal/d. The largest increase in ground-water use is expected to be in the Princess Anne area, where a State correctional facility has recently (1987) been constructed. This facility is expected to withdraw about 450,000 gal/d of ground water from the Manokin aquifer, when it is in full operation (J. Windsor, Somerset County Department of Technical and Community Services, written commun., 1986).

AQUIFERS AND CONFINING UNITS

Although the principal sand units underlying Somerset County are water bearing, several of them are not used for water supply because they contain poor-quality water. The focus of this report is on those aquifers and aquifer systems that supply water suitable in quantities and quality for human needs. They are the surficial aquifer system, the Pocomoke aquifer, the Manokin aquifer, the Paleocene aquifer system, and the upper sands of the Potomac aquifer system. Aquifers that yield marginal or poor-quality water are the Choptank aquifer, the Piney Point aquifer, and the lower sands of the Potomac aquifer system. A number of wells (54) in the Crisfield area are screened in deposits that do not belong to any of the aquifer units listed above. The deposits are individual sands in confining units below the Pocomoke aquifer and above the Choptank aquifer. Because the Manokin aquifer may be poorly developed where these deposits are present, it is difficult to assign the deposits to a particular hydrogeologic unit. In this report, these deposits are undifferentiated and assigned to the Miocene Series.

Surficial Aquifer System

Description

In this report, the uppermost aquifer system in Somerset County is called the surficial aquifer system and includes Holocene alluvium, and swamp and tidal-marsh deposits; the Parsonsburg Sand; the Kent Island Formation; the Omar Formation; and the Beaverdam Sand (table 1). As such, the surficial aquifer system is stratigraphically complex, and commonly exhibits numerous lithologic changes over short distances both laterally and vertically. Table 3 briefly describes the individual units of the surficial aquifer system and figure 6 shows the areal distribution of these units. In general, the alluvium, swamp deposits,

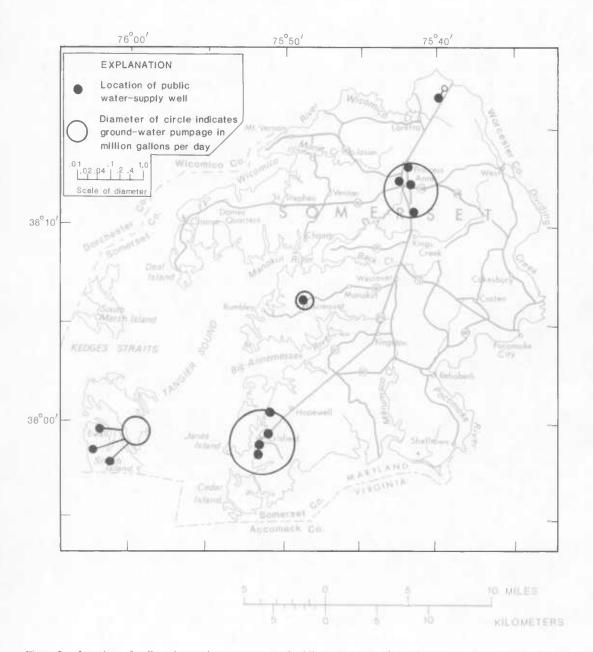


Figure 5.— Locations of wells and ground-water pumpage of public water systems that withdrew more than 10,000 gallons per day in 1986.

TABLE 2
GROUND-WATER USERS IN SOMERSET COUNTY, THAT WITHDREW MORE THAN 10,000 GALLONS PER DAY IN 1986
[PCPC = surficial aquifer system; PNPN = Piney Point aquifer; PCMK = Pocomoke aquifer; PLCN = Paleocene aquifer system; MNKN = Manokin aquifer; PTMC = Potomac aquifer system]

User	Well no.	Aquifer	Withdrawal rate in thousand gallons per day
Chesapeake Plywood, Inc.	SO Df 10	MNKN	94.67
	SO Df 11	PCMK	
	SO Df 12 SO Dg 4	PCMK PCMK	
City of Colodials	9		994 90
City of Crisfield	SO Dc 4 SO Ec 1	PTMC PLCN	884.89
	SO Ec 2	PLCN	
	SO Ec 3	PTMC	
	SO Ec 4	PNPN	
	20 20 4	PLCN	
		PTMC	
	SO Ec 42	PLCN	
	20 20 12	PTMC	
	SO Ec 48	PTMC	
	SO Ec 49	PLCN	
		PTMC	
Eden Mobile Home Park	SO Af 22	MNKN	22.77
	SO Af 23	PCPC	
	SO Af 24	MNKN	
	SO Af 25	MNKN	
Perdue, Inc.	SO Ce 85	MNKN	22.36
	SO Ce 86	MNKN	
Smith Island (7 ayatems)	SO Ea 3	PTMC	58.50
	SO Ea 4	PTMC	
	SO Ea 6	PTMC	
	SO Ea 7	PTMC	
	SO Ea 8	PTMC	
	SO Ea 12	PTMC	
	SO Ea 13	PTMC	
Somerset County Sanitary Diatrict (Fairmount)	SO Cd 41	PTMC	33.20
Someraet County Sanitary	SO Be 51	MNKN	413.94
Diatrict (Princess Anne)	SO Be 54	MNKN	
	SO Be 56	MNKN	
	SO Be 78	MNKN	

TABLE 3
GEOLOGIC UNITS IN THE SURFICIAL AQUIFER SYSTEM (from Owens and Denny, 1984)

Geologic Unit	Description
Alluvium, swamp, end tidal-mersh deposits, undifferentieted	Silt and clay, interbedded with some fine send Found along the Chesepeake Bey end large streams. Typically less than 5 feet thick.
Parsonsburg Send	Predominently loose, light-colored quartz sand. Largest occurrence is a small uplend in the northeastern corner of the county. Thickness ranges from 4 to 20 feet.
Kent Island Formetion	Interbedded grevel, send, silt, end clay. Sendy end gravelly in the eestern pert of the county, where it overlies the Beaverdam Sand; clayey and silty in the western pert of the county. Thickness ranges from 3 to 10 feet, except in chennels, where it is up to 43 feet thick.
Omar Formation	Dark-gray silty send, silt, end silty clay. Occurs in a small eree neer the eastern border of the county. Meximum thickness is 25 feet.
Beaverdam Sand	Sand and silty sand, locally inter- bedded with grevelly sand, cley and silt. Underlies the northeastern part of the county. The unit may be as much as 100 feet thick in channels near the northeestern border of the county.

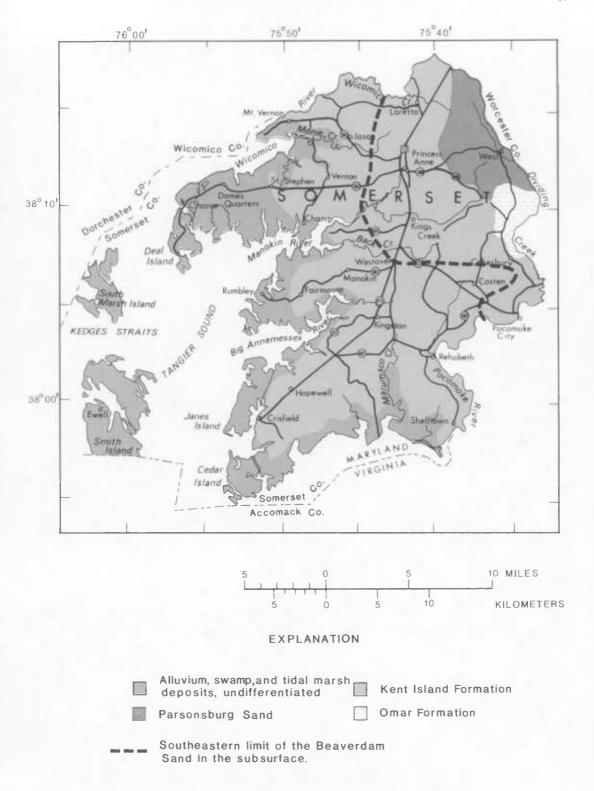


Figure 6.— Generalized geologic map of Somerset County (modified from Owens and Denny, 1984).

tidal-marsh deposits, and Omar Formation consist of silty sand, silt, and clay; the Parsonsburg Sand is composed of medium sand; and the Beaverdam Sand is sand and silty sand. The Kent Island Formation is gravelly and sandy where it overlies the Beaverdam Sand and silty and clayey where it overlies finer-grained sediments. As a whole, the aquifer system generally is coarser grained in the northeastern part of the county where it includes the Beaverdam Sand, and finer grained in the southwestern part of the county where the Beaverdam Sand is absent (fig. 6).

Plate 5 shows the thickness of the surficial aquifer system in Somerset County. The aquifer system is about 20 ft thick throughout most of the county, except in the northeastern part, where it is more than 40 ft thick. The geologic units that comprise the aquifer system were deposited on an erosional surface that contained numerous channels. Therefore, where these paleochannels are present, the aquifer system thickens. Hansen (1966) describes a paleochannel in the Salisbury area that is about 175 ft deep, and Owens and Denny (1979, p. A13) show paleochannels in the Maryland and Delaware portions of the Delmarva Peninsula. The thickening of the surficial aquifer system in the northeastern part of the county may be due, in part, to the infilling of a paleochannel. Generally, however, paleochannels in Somerset County are not of the magnitude of those described in other areas by Hansen (1966), and Owens and Denny (1979, p. Al3). Mixon (1985, p. Gl8) and Owens and Denny (1984) show channels in the Yorktown and Eastover Formations and Yorktown and Cohansey Formations, respectively (Yorktown and Eastover Formations, in this report), that have maximum depths of 30 to 40 ft. Where the Pocomoke aguifer subcrops beneath the surficial aquifer system (pl. 6), sand-on-sand contacts make paleochannel identification difficult. In these areas, the base of the surficial aquifer system was determined from nearby wells that have a significant (5-10 ft) confining unit separating the aquifer system and the Pocomoke aquifer. Therefore, sediments in paleochannels in the recharge areas are included in the Pocomoke aquifer.

The surficial aquifer system is used primarily as a source of domestic water supply, as the generally thin deposits of the aquifer system preclude its use for public water supply. In the northeastern part of the county, however, where the aquifer system is thicker, larger quantities of water may be obtainable. Also, in areas where the surficial aquifer system directly overlies the Pocomoke aquifer (pl. 6), the two units in combination may yield significant quantities of water.

Well yield and specific capacity

Reported yields of 12 wells in the surficial aquifer system range from 4 to 80 gal/min (gallons per minute). The median reported well yield is 12 gal/min. Reported specific capacity ranges from 1.3 to 27 (gal/min)/ft (gallons per minute per foot), with a median value of 3.4 (gal/min)/ft. Although the data are not sufficient to adequately describe areal distributions of well yield or specific capacity, yields are generally greater in the northcastern part of the county where the aquifer system is thicker.

Water levels

Somerset County is highly dissected by streams and drainage ditches that drain the surficial aquifer system. Consequently, water levels in the surficial aquifer system are adjusted

to water-surface altitudes in the streams and ditches. In the northeastern part of the county where stream altitudes are highest, water levels in the aquifer system also are high. Likewise, in the southwestern part of the county, streams are tidally influenced and water levels in the aquifer system are only a few feet above sea level (fig. 7).

Water levels in the surficial aquifer system respond to seasonal changes in recharge and evapotranspiration. Figure 8 shows water levels in well SO Cf 2 for 1949-88. Seasonal fluctuations are evident as water levels rise to about 1 ft below land surface in the spring and decline to 6 ft below land surface in late summer. These seasonal fluctuations occur because, from fall to early spring, evapotranspiration is low and precipitation is moderate, resulting in recharge to the unconfined aquifers. During the late spring and summer growing season, evapotranspiration is high, so little precipitation reaches the water table. Accordingly, water levels decline until the growing season ends, and then begin their seasonal recovery.

Confining Unit between the Surficial Aquifer System and Pocomoke Aquifer

In the eastern and southeastern part of Somerset County, a confining unit retards the vertical movement of water between the surficial aquifer system and the Pocomoke aquifer (pls. 1-4). This confining unit is in the upper part of the Yorktown and Eastover Formations (undifferentiated) and eonsists primarily of elay and silt (Mixon, 1985, p. Gl0-Gl4) except in the vicinity of Costen. In this area, there are no fine-grained sediments separating the surficial aquifer system and the Pocomoke aquifer. The confining unit in this area may be primarily sand, or the unit may have been eroded and replaced with sands of the surficial aquifer system. In either ease, vertical ground-water flow between the surficial aquifer system and the Pocomoke aquifer is probably greater in this area than in areas where the clay and silt separate the two aquifers. Thickness of the eonfining unit ranges from zero where it is truncated by the surficial aquifer system to about 65 ft at Rehobeth (pls. 1-4).

Pocomoke Aquifer

Description

The Pocomoke aquifer, which is part of the Yorktown and Eastover Formations (undifferentiated) (table 1), is present only in the southeastern two-thirds of the county (pl. 6). It eonsists primarily of gray, fine- to medium-grained fossiliferous sand with stringers of gravel and coarse sand (Rasmussen and Slaughter, 1955, p. 101). Locally, the aquifer contains glaueonitie sand and elay and silt interbeds (Owens and Denny, 1979, p. 9).

Plate 6 shows recharge areas and altitude of the top of the Pocomoke aquifer. Recharge to the aquifer occurs primarily where it directly underlies the surficial aquifer system. This is in a 1- to 5-mi (mile) wide band that trends from southwest to northeast through Somerset County. The Pocomoke aquifer may receive additional ground-water recharge from leakage from the surficial aquifer system in the vicinity of Costen.

The top of the aquifer generally slopes to the southeast, with altitudes ranging from about 20 ft below sea level at the northwestern limit of the aquifer to 70 ft below sea level near the county line at Pocomoke City (pl. 6). The number and vertical position of silt and clay units

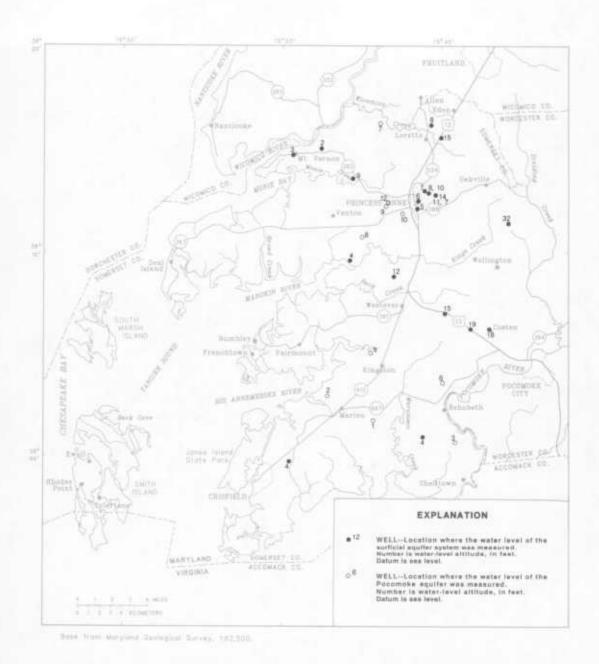


Figure 7.— Water-level altitudes in the surficial aquifer system and the Poeomoke aquifer, April 1987.

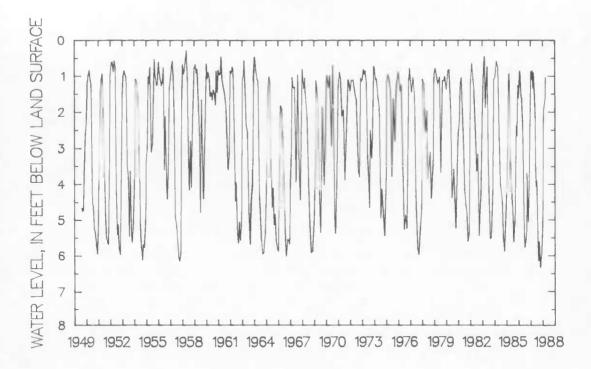


Figure 8.— Water levels in observation well SO Cf 2, 1949-88 (well location shown on quadrangle map Cf at end of report).

in the aquifer differs, making identification of the top and bottom of the aquifer difficult in places. This may account for some of the variation in the altitudes shown on pl. 6. Thickness of the aquifer also is variable (pl. 7), and ranges from zero at its northwestern limit to more than 75 ft in the southeastern part of the county.

The Pocomoke aquifer supplies water for domestie, irrigation, and industrial use in Somerset County and is the sole source of water for Poeomoke City, which is located southeast of the study area. In areas where both the Pocomoke aquifer and Manokin aquifer contain potable water, the Manokin is usually the aquifer used for drinking-water supplies, because water in the Pocomoke aquifer frequently contains greater concentrations of iron. In the southern part of the county, however, chloride concentrations in the Manokin aquifer exceed the secondary maximum contaminant level (SMCL) of 250 mg/L (milligrams per liter) (U.S. Environmental Protection Agency, 1986b) and the Pocomoke aquifer is preferred for potable water. Nearly all irrigation wells in Somerset County are sereened in the Pocomoke aquifer because it is encountered at shallower depths and tends to contain more gravel beds than the Manokin aquifer (Owens and Denny, 1979, p. A9).

Well yield and specific capacity

Well yields reported in water-well completion reports for 104 wells in the Pocomoke aquifer range from 6 to 800 gal/min, with a median of 20 gal/min. Reported well yields for irrigation wells are generally greater than the median for all wells, and probably more ac-

curately reflect the short-term, water-producing capacity of the aquifer. Reported yields for 24 irrigation wells range from 10 to 800 gal/min, with a median of 60 gal/min.

Specific capacities reported for 68 water wells range from 1 to 50 (gal/min)/ft with a median value of 10 (gal/min)/ft. Five 1-hour specific-capacity tests were conducted during this investigation. Specific capacity computed from these data ranges from 2.0 to 17.3 (gal/min)/ft (table 4).

Figure 9 shows the reported specific-capacity data for the Pocomoke aquifer. Throughout most of the county, reported specific capacity ranges from 5 to 10 (gal/min)/ft. In the center of the county, however, specific capacity increases to over 20 (gal/min)/ft. Because of the variability in reported values, the specific-capacity data in figure 9 are approximate and cannot be used to determine the value of specific capacity at a particular site.

Hydraulic properties

Transmissivity and storage coefficient of the Pocomoke aquifer are unknown as no aquifer-test data are available for Somerset County. In Pocomoke City, which is located southeast of Somerset County, transmissivities calculated from two aquifer tests are 1,070 ft²/d (feet squared per day) and 5,350 ft²/d (Rasmussen and Slaughter, 1955, p. 145-150). Storage coefficients calculated from these aquifer tests are 0.003 and 0.0002, respectively. These values indicate that the water-yielding capacity of the Pocomoke aquifer may change considerably over short distances.

Water levels

There are no long-term observation wells screened in the Pocomoke aquifer in Somerset County, and water levels were measured in only nine wells in the Pocomoke aquifer during this investigation (fig. 7). Therefore, temporal and spatial water-level trends in the Pocomoke aquifer are unknown. However, in recharge areas (pl. 6) the Pocomoke aquifer is unconfined and water levels probably are controlled strongly by local streams and drainage

TABLE 4
SPECIFIC CAPACITY OF SELECTED WELLS IN THE POCOMOKE AQUIFER [(gal/min)/ft = gallons per minute per foot]

Well no.	Specific capacity [(gal/min)/ft]	
SO Be 77	3.2	
SO Be 88	17.3	
SO Dd 58	3.3	
SO De 36	5.7	
SO Df 21	2.0	

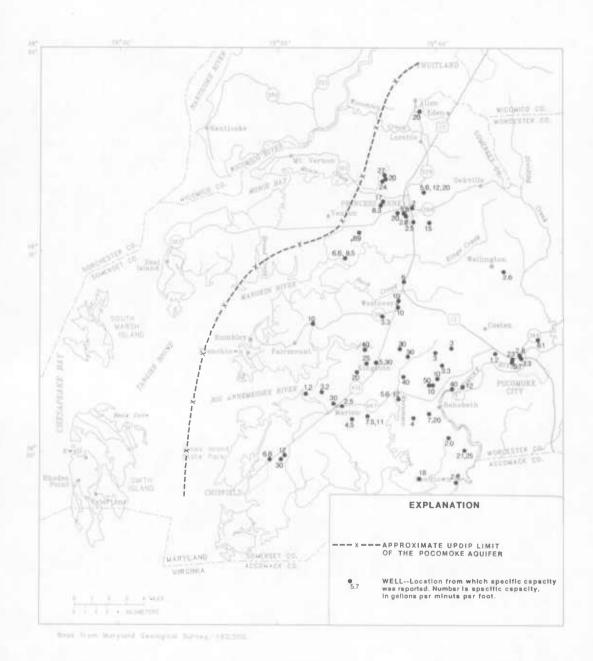


Figure 9.— Reported specific capacity of wells in the Pocomoke aquifer.

ditches (fig. 7). Where the Pocomoke aquifer is overlain by a confining unit, the vertical movement of water to and from the aquifer is inhibited and water levels are probably less influenced by local surface-water bodies.

Confining Unit between the Pocomoke and Manokin Aquifers

A confining unit consisting of silt, elay, and fine-grained sand separates the Pocomoke aquifer from the underlying Manokin aquifer (pls. 1-4). The thickness of the confining unit differs, ranging from less than 40 ft near Eden to more than 100 ft near Princess Anne (fig. 10). In the Crisfield area, the lithology of the confining unit changes from predominantly silt and clay to one characterized by more interbeds of sand. Several of these interbeds are transmissive enough to supply water to domestic wells in the area. The change in character of the confining unit in the Crisfield area also may allow more water to move vertically between the Pocomoke aquifer and the Manokin aquifer.

Manokin Aquifer

Description

The Manokin aquifer is the primary aquifer used for water supply in Somerset County. It is in the Eastover Formation and consists principally of gray, finc- to medium-grained sand and contains some shell material. In the western part of the county, in the area encompassing Fairmount, Kingston, Janes Island State Park, and Smith Island, the unit becomes finer grained and is no longer recognizable as a distinct aquifer in geophysical logs (pls. 1-4).

The Manokin aquifer subcrops beneath the Chesapeake Bay west of Deal Island and trends northeast through Dorehester and Wieomieo Counties into Delaware near Seaford (Hansen, 1981, p. 129; Pickett, 1976). The top of the aquifer slopes southeast at about 9 ft/mi, with the altitude of the top of the unit ranging from about 75 ft below sea level at Deal Island to about 190 ft below sea level near Wellington (fig. 11). Figure 12 shows the thickness of the Manokin aquifer, which ranges from zero, where the aquifer becomes finer grained, to more than 80 ft in the northeastern corner of the county.

Most domestic wells in Somerset County withdraw water from the Manokin aquifer. The aquifer also serves as the source of water for the town of Princess Anne and the Eastern Correctional Institution. In addition, several poultry-raising operations and a plywood-manufacturing plant withdraw water from the Manokin aquifer. South of the town of Westover, however, chloride concentrations in the aquifer exceed the 250-mg/L SMCL of the U.S. Environmental Protection Agency (USEPA) (1986b), and the Manokin aquifer usually is not used as a source of potable water in this area.

Well yield and specific capacity

Reported yields for 269 wells in the Manokin aquifer range from 2 to 350 gal/min, with a median of 20 gal/min. Reported specific capacities for 198 wells range from 0.1 to 75

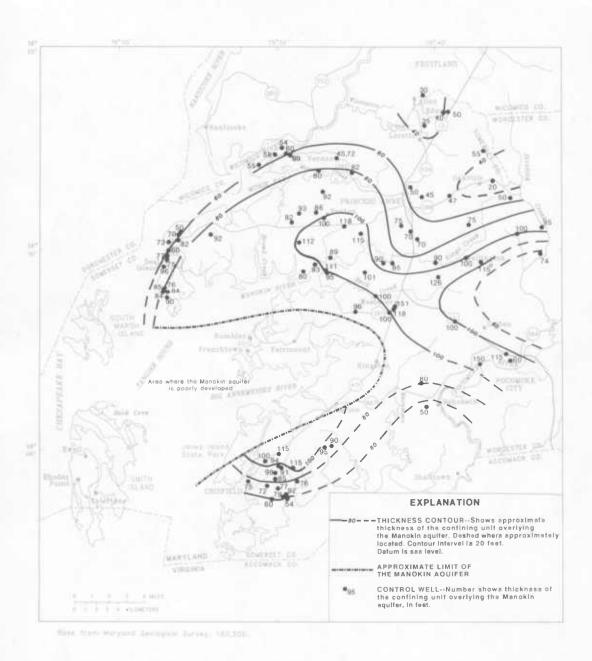


Figure 10.— Thickness of the confining unit overlying the Manokin aquifer.

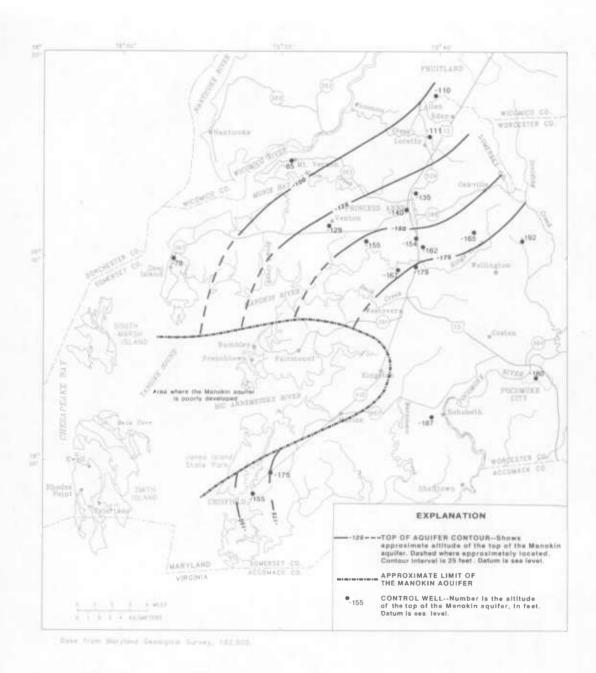


Figure 11.- Altitude of the top of the Manokin aquifer.

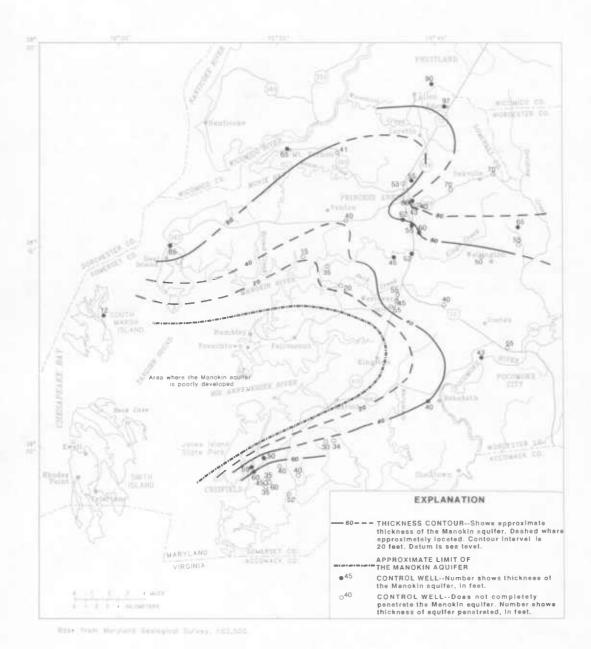


Figure 12.— Thickness of the Manokin aquifer.

(gal/min)/ft, with a median specific capacity of 5.2 (gal/min)/ft. No areal trends are evident in the distribution of well yield or specific capacity.

Hydraulie properties

The hydraulie characteristics of the Manokin aquifer are not well known throughout the county, but four multiple-well aquifer tests and a 48-hour, single-well aquifer test (at well SO Be 55) were performed in the Princess Anne area. Table 5 includes transmissivities, hydraulie conductivities, and storage coefficients estimated from the Manokin test data. Transmissivity ranges from 500 to 940 ft²/d. Hydraulie conductivity, which is determined by dividing transmissivity by aquifer thickness, ranges from 10.9 to 14.7 ft/d (feet per day) and has an average value of 13.2 ft/d. Storage coefficients calculated from the four multiple-well aquifer tests range from 0.0002 to 0.001.

Sediments of the Manokin aquifer appear to be relatively uniform with respect to grain size and sorting. Most driller's logs describe the aquifer as fine- or medium-grained sand, with some silty sand. Therefore, as a first approximation of the areal distribution of transmissivity, the average hydraulic conductivity (13.2 ft/d) was multiplied by aquifer thickness as determined from figure 12.

These initial values were modified during calibration of the digital model of the Manokin aquifer, as described later in this report. Figure 13 shows the modified transmissivity distribution. The largest values of transmissivity are in the northeastern corner of the county where values may exceed 1,000 ft²/d. Transmissivity decreases to the south and east to about 200 ft²/d in the vicinity of Crisfield. North of Crisfield, and encompassing the towns

TABLE 5
TRANSMISSIVITIES, HYDRAULIC CONDUCTIVITIES, AND STORAGE COEFFICIENTS
FOR AQUIFERS, AS DETERMINED FROM AQUIFER TESTS
[gal/min = gallons per minute; ft=feet; ft²/d = feet squared per day;
ft/d = feet per day; -- = no data]

Pumped	Date and duretion	Pumping rete				relle and pumping	Aquifer	Anelyeia	Method of anelysis	Trane- missivity	Bydreulic conductivity	Storege coefficient
well	of test	(gel/min)		We	11 (1	(t)	aquifer eyetem	by		(ft ² /d)	(ft/d)	
SO Be 51	4/12/54 17 daya	115	SO	Be	42	1,300	Manokin	Raamuasen and Sleughter		940	10,9	0.0002
SO Be 55	6/11/85 48 hours	128					Manokin	W. B. Werkheiser	Cooper and Jacob eemi- log	920	14.2	
SO Ba 56	5/29/85 24 houre	503	SO	Be	110	10	Manokin	W. H. Werkheiser	Bantush-Jecob leaky erteeian	770	13.1	.001
SO Cd 41	5/28/79 24 houre	100					Potomec	W. H. Werkheiser	Cooper and Jecob eemi- log	1,280		
SO Ce 44	7/ 1/85 48 hours	180		Co Co		600 2,500	Manokin	W. H. Werkheiear	Hantush-Jecob leaky artaaian	760 740	14.9 14.5	.0002
SO Ca 48	5/28/85 48 houre	125	SO	Ce	51	600	Manokin	W. H. Werkheiser	Hantush-Jacob laaky erteaian	500	13.3	,0002
SO Dc 4	12/21/70 24 houra		SO	Dc	3	200	Potomac	I. H. Kantrowitz	Theie non- leaky arteeien	2,140	~-	.0002
SO Ec 49	7/17/84 24 hours	800					Peleo- cana and Potomac	W. H. Warkhaisar	Cooper and Jacob aami- log	1,050	e 0	

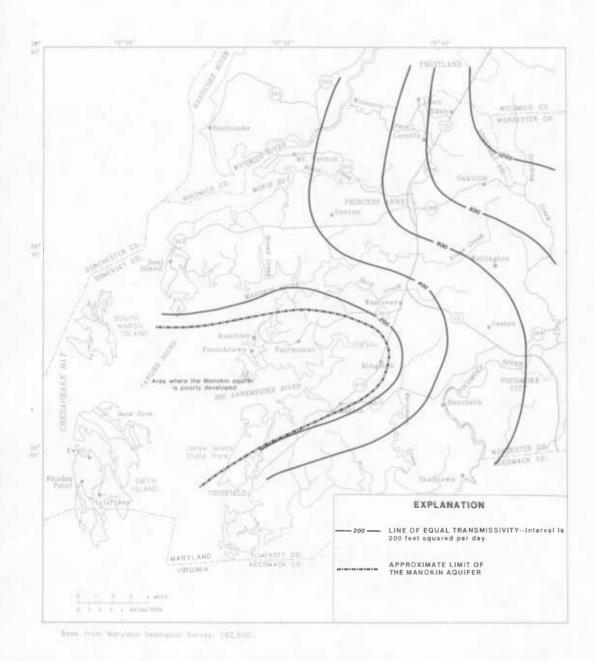


Figure 13.— Transmissivity of the Manokin aquifer used in the model of the Princess Anne area.

of Rumbley and Kingston, is an area where the sediments that are equivalent to the Manokin aquifer consist primarily of very fine sand and silt and no longer function as an aquifer. In this area the primary source of water is either the overlying Pocomoke aquifer or the deeper Palcocene and Potomae aquifer systems.

Water levels

During the early 1950's, before the Manokin aquifer was developed extensively, water-level altitudes in the subcrop of the Manokin aquifer were highest near the Delaware-Maryland border at Delmar (Rasmussen and Slaughter, 1955, p. 99) and were probably near sea level beneath the Chesapeake Bay. Ground-water levels for the aquifer in Somerset County ranged from about 20 ft above sea level near Eden to about 5 ft above sea level near Crisfield (Rasmussen and Slaughter, 1955, p. 208-238). This suggests that under prepumping conditions, the higher heads in the vicinity of the Delaware-Maryland border would cause ground water to flow across Somerset County from northeast to southwest and west.

In the 1950's, large-eapacity wells were drilled in the Manokin aquifer for public and industrial water supply. Due to this pumping, water levels in the aquifer have declined by as much as 45 ft. Figure 14 shows the altitude of water levels for the aquifer during April 14-16, 1987. The lowest altitudes, about 20 ft below sea level, arc in the vicinity of Prineess Anne where several large-capacity wells are located. The highest altitudes, about 10 ft above sea level, are in the northeastern corner of the county.

The water-level declines in the vicinity of Princess Anne have had several effects on the ground-water-flow system. First, the reduction in head in the Manokin aquifer has increased the leakage of water from overlying and underlying units into the aquifer. Second, ground-water-flow directions have changed from a regional (northeast to southwest) to a local pattern, with ground water moving radially toward the pumping centers at Princess Anne. There is concern that pumpage in the Princess Anne area will cause movement of the high-chloride (chloride concentration in excess of 250 mg/L) water south of Westover. Figure 14 shows that water levels in the Manokin aquifer throughout most of the county are below sea level, with the deepest water levels occurring near Princess Anne. Therefore, the potential exists for the high-chloride water south of Westover to migrate toward Princess Anne. In addition to migration of high-chloride water from the southern part of the county, migration of high-chloride water from west of the county also is possible. The Manokin aquifer subcrops beneath the Chesapeake Bay to the west of Somerset County. With the change in ground-water-flow directions associated with ground-water withdrawals near Princess Anne, brackish water from the Chesapeake Bay could intrude into the aquifer and migrate toward the pumping centers at Princess Anne.

Water levels shown in figure 14, when compared with the top of the aquifer shown in figure 11, suggest that withdrawals from the Manokin aquifer could be increased without dewatering the aquifer. However, the increased hydraulic gradient associated with increased pumping would result in more rapid migration of high-chloride water toward pumping centers. The effects of increased withdrawals from the Manokin aquifer will be more fully addressed in later sections.

Figure 15 depicts water-level fluctuations in observation well SO Be 42. This well is located near several pumped wells at Princess Anne that cycle on and off throughout the day, causing water levels in the observation well to vary substantially in a short time period.

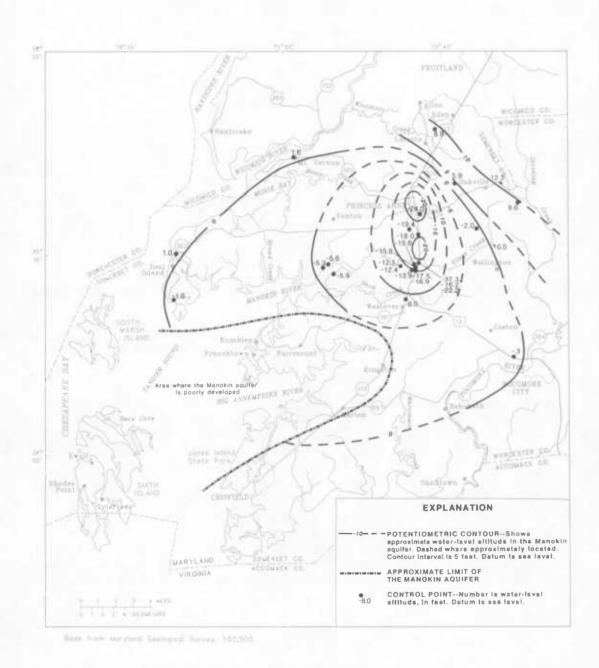


Figure 14.— Water-level altitudes in the Manokin aquifer, April 14-16, 1987.

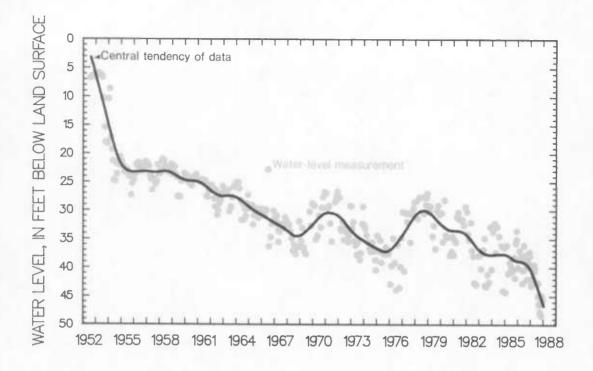


Figure 15.— Water levels in observation well SO Be 42, 1952-88 (well location shown on quadrangle map Be at end of report).

Therefore, the water levels shown in figure 15, which were measured at about 6-week intervals, show eonsiderable seatter. The line drawn through the data shows the central tendency of the data and was obtained by using a locally weighted, scatterplot smoothing routine. The sharpest water-level decline was during 1953-55, in response to the initial pumping of the high-capacity wells at Princess Anne. In the late 1950's as the ground-water system adjusted to this pumpage, water levels began to stabilize at about 23 ft below land surface. During the 1960's water levels again started to decline, probably in response to increased ground-water withdrawals. Throughout the 1970's and early 1980's water levels oscillated, possibly due to changes in recharge and total ground-water pumpage. The oscillations also may have resulted, in part, from the redistribution of ground-water pumpage that occurred when additional production wells were installed in 1967 and 1976. In 1988, water levels again declined as ground-water withdrawals increased at Princess Anne and the Eastern Correctional Institution.

Confining Unit between the Manokin and Choptank Aquifers

A confining unit composed of gray clay and silty clay (St. Marys Formation) separates the Manokin aquifer from the underlying Choptank aquifer (Rasmussen and Slaughter, 1955, p. 93). Although data are inadequate to define the extent of this unit in Somerset County, in those wells that fully penetrate the St. Marys Formation, the thickness ranges from 70 ft on Deal Island to 190 ft near Rehobeth (pls. 1-4). The unit is present at every lo-

cation where wells have been drilled deep enough to encounter it. Therefore, it appears that it is laterally continuous in Somerset County. The thickness, fine-grained nature, and lateral extent of the unit probably allow little water to exchange between the Manokin aquifer and the Choptank aquifer.

Choptank Aquifer

The Choptank aquifer, which consists of the Choptank Formation, is composed of gray, coarse- to fine-grained sand, with shell beds and lenses of gray clay (Rasmussen and Slaughter, 1955, p. 86). The top of the aquifer occurs at about 225 ft below sea level on Smith Island, and dips to the east at about 10 ft/mi (fig. 16). Based on limited data, the aquifer appears to thicken from about 70 ft at Smith Island eastward to 150 ft at Rehobeth (pls. 1-4).

Although the aquifer is capable of supplying adequate quantities of water to wells, chloride concentrations in excess of 900 mg/L and dissolved solids in excess of 3,000 mg/L preclude its use as a source of water for most purposes. Rasmussen and Slaughter (1955, p. 208-238) report several wells finished in the Choptank aquifer at Deal Island and Crisfield, but presently no wells in Somerset County are known to produce water from the aquifer. Rasmussen and Slaughter (1955, p. 89) suggest that the quality of water in the aquifer improves toward the subcrop area, which is about 30 mi north of Somerset County.

Confining Unit between the Choptank and Pincy Point Aquifers

A confining unit consisting of the Calvert Formation separates the Choptank and Piney Point aquifers. The confining unit in Somerset County consists predominantly of blue, green, and brown clay, but may contain local sand interbeds (Rasmussen and Slaughter, 1955, p. 83-85). In Wicomico County, the sands in the Calvert Formation become extensive enough to be termed the Nanticoke aquifer by Rasmussen and Slaughter (1955, p. 85), but in Somerset County the entire Calvert Formation appears to be a confining unit. The unit appears to be laterally continuous throughout Somerset County and is up to 400 ft thick (pls. 1-4). The fine-grained nature, lateral continuity, and thickness of the unit probably allow little water to exchange between the Choptank aquifer and the Piney Point aquifer.

Piney Point Aquifer

Description

The Piney Point aquifer is in the Piney Point Formation and is predominantly green, fine-to medium-grained glaueonitie sand and gray, coarse-grained quartzose sand in a greenish-gray clay matrix (Hansen, 1967, p. 3). It occurs at depths ranging from 580 ft below land surface on Deal Island to 950 ft below land surface at Rehobeth (pls. 1-4). At Crisfield, the Piney Point is encountered at a depth of about 730 ft (pl. 4). The thickness of the aquifer ranges from 50 ft at Rehobeth to 85 ft on Smith Island (pls. 1-4). The characteristics of the Piney Point aquifer are not known in the northern part of the county, as no wells have been

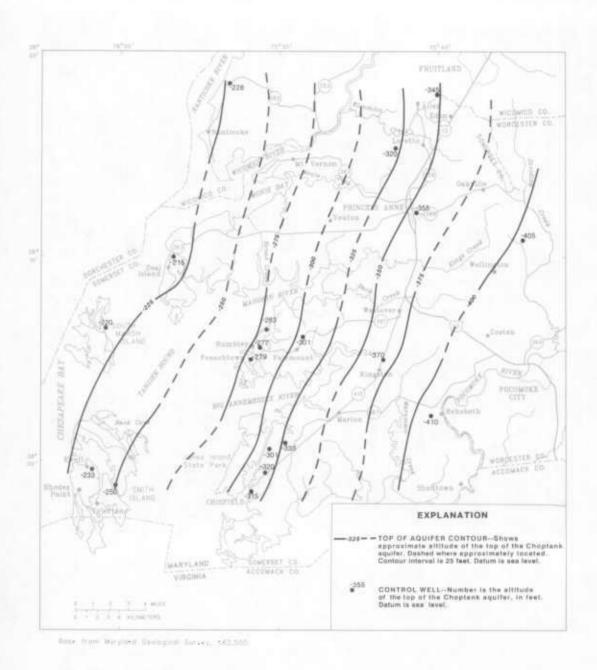


Figure 16.— Altitude of the top of the Choptank aquifer.

drilled in the unit. A single chemical analysis (from well SO Bb 19) collected during this investigation indicates that water in the aquifer may contain dissolved solids in excess of 1,000 mg/L, which may make the aquifer undesirable as a source of potable water. A municipal well in Crisfield, SO Ee 4, derives water from the Piney Point and deeper aquifers so that the more mineralized water from the Pincy Point aquifer is mixed with the less mineralized water from the other aquifers. This well currently is not in use.

Well yield

The water-producing capabilities of the Piney Point aquifer are virtually unknown, as reported well yields are available for only three wells. The reported yields are 4, 10, and 30 gal/min. Rasmussen and Slaughter (1955, p. 79) report that the aquifer is absent east of Somerset County, apparently grading into a fine-grained unit. Therefore, the water-yielding potential of the Piney Point aquifer may be lower in the eastern part of the county.

Water levels

Water levels in the Piney Point aquifer are unknown as there are no observation wells screened solely in the aquifer. Rasmussen and Slaughter (1955, p. 210, 217) report water levels in two wells, SO Bb 1 and SO Cc 1, that were 5 ft and 7 ft above sea level, respectively. As the aquifer is unstressed in Somerset County, 1988 water levels are likely to be similar to those of the early 1950's. Observation well SO Ec 4 is screened in the Piney Point aquifer as well as deeper aquifers. The water level in this well varies from about 12 ft to 18 ft below land surface (fig. 20). This indicates that the Piney Point aquifer does not greatly influence the water level in the well. This may be because the screen is elogged in the vicinity of the aquifer, or it may indicate that the transmissivity of the Piney Point aquifer at this location is lower than the transmissivity of the other aquifers in which the well is screened.

Confining Unit between the Piney Point Aquifer and Paleocene Aquifer System

A confining unit of gray, green, or black sand and clay separates the Piney Point aquifer from the underlying Paleocene aquifer system. In this report the confining unit is considered part of the Piney Point Formation, although other investigators have placed these sediments in the Eocene Nanjemoy Formation (Hansen, 1978, fig. 10), or have considered them as part of the Paleocene Series (Rasmussen and Slaughter, 1955, p. 311; Hansen, 1967, p. 11). Thickness of the unit ranges from 10 to 25 ft (pls. 1, 3-4). Depth to the top of the unit ranges from 690 ft below land surface at Deal Island to 1,000 ft below land surface at Rehobeth (pl. 4). The characteristics of the confining unit in the northern part of the county are not known.

Paleocene Aquifer System

Description

Underlying the Piney Point aquifer is a series of aquifers and confining units collectively called the Palcocene aquifer system in this report, following the usage of the Maryland Geo-

logical Survey. The aquifers generally consist of fine- to medium-grained glauconitic and quartzose sand, whereas, the confining units are generally composed of gray to green elay and sandy clay (Hansen, 1967, p. 3). The aquifer system is encountered at depths of about 720 ft on Smith Island and about 800 ft at Crisfield (pl. 4). Total thickness of the system is about 90 ft on Smith Island and about 175 ft near Crisfield (pl. 4). In those wells that obtain water from the Paleocene aquifer system, screens are usually set in the lower sands of the aquifer system. The sands in the Paleocene aquifer system appear to thin or change facies to the north and east of Crisfield, causing the Paleocene sediments to act as a confining unit in other areas of the lower Eastern Shore (Rasmussen and Slaughter, 1955, p. 74).

In 1988, the city of Crisfield was the only municipality using water from the Paleocene aquifer system. Municipal wells SO Ec 1 and SO Ec 2 are screened solely in the aquifer system and three others—SO Ec 4, SO Ec 42, and SO Ec 49—are screened in multiple aquifers including the Paleocene aquifer system. On Smith Island, and at the towns of Rumbley, Frenchtown, and Fairmount, the Paleocene aquifer system is bypassed, and wells are screened in the more productive sands of the Potomac aquifer system.

Well yield and specific capacity

Because few wells are screened solely in the Paleocene aquifer system, little is known about its water-yielding capabilities. The specific capacity of wells SO Ec 1 and SO Ec 2 are 1.8 and 1.9 (gal/min)/ft, respectively. The yield of well SO Ec 2 was reported to be 300 gal/min during the specific-capacity test. These wells are adjacent to each other and both are finished in the lower sands of the aquifer system. Therefore, these specific-capacity values may not be representative of the entire aquifer system.

Hydraulic properties

No aquifer-test data are available for wells screened only in the Palcocenc aquifer system. A single-well aquifer test was conducted on well SO Ec 49, which is screened in two lower sands in the Paleocene aquifer system and the three upper sands of the Potomac aquifer system (pl. 4). The composite transmissivity calculated for the screened sands of these two aquifer systems is 1,050 ft²/d (table 5).

Water levels

Water levels in the Paleocene aquifer system are not known, because no observation well is screened entirely in the system. Rasmussen and Slaughter (1955, p. 59) report static water levels in newly drilled wells in 1950 to be above land surface. In 1986, the water level in well SO Ec 42, screened in both the Paleocene aquifer system and the underlying Potomae aquifer system, was 2 ft below land surface.

The fact that available drawdown in the aquifer system is about 800 ft suggests that greater yields should be possible without dewatering the aquifers. However, because the aquifer system becomes finer grained between Crisfield and Salisbury, it is not known if the Paleocene aquifer system could sustain significantly larger ground-water withdrawals without producing adverse effects, such as excessive water-level declines in the aquifer.

Potomac Aquifer System

Description

The deepest hydrogeologic units in Somerset County that produce water of acceptable quality are in the aquifer system of the Potomac Group. The aquifer system, as reported in descriptions of well cuttings, consists of white, yellow, and gray, fine- to coarse-grained sand intercalated with gray, green, and red clay (Hansen, 1967, p. 11). The upper sand units in the aquifer system were identified as belonging to the Magothy Formation by Rasmussen and Slaughter (1955, p. 53-55). Later work by Hansen (1978) suggests that marine Upper Cretaeeous beds are absent in Somerset County in the Crisfield area, and that the strata underlying the Paleoeene aquifer system are part of the Potomac Group.

The sediments of the Potomae Group are of deltaic to fluvial origin (Rasmussen and Slaughter, 1955, p. 45; Hansen, 1982, p. 3). As such, individual beds of sand, silt, and elay may have restricted areal and vertical extent. Because of this, correlation of individual beds over more than a few miles is difficult.

The top of the Potomac aguifer system is encountered at about 820 ft below land surface on Smith Island and at about 1,000 ft below land surface near Crisfield (pl. 4). A geothermal test well, SO Dd 47, was drilled through the Potomae Group sediments and encountered basement rock at a depth of about 4,225 ft (Hansen, 1982, p. 12). Although the data from this well indicate that the Potomac Group is over 3,000 ft thick at Crisfield, only the sands in the upper several hundred feet of the unit are thought to contain potable water. The deepest aquifer yielding potable water occurs at a depth of 1,295 ft in SO Ec 49 (pl. 4). The multi-point electric log for SO Dd 47 (fig. 17) suggests that one or two deeper aquifers may contain potable water (H.J. Hansen, Maryland Geological Survey, written commun., 1988). Electrical resistivity of an aquifer is inversely proportional to the dissolved-solids eontent of water in an aquifer (Keys and MacCary, 1983, p. 42). The sand that occurs from about 1,410 to 1,460 ft in SO Dd 47 has a resistivity of about 14 ohm-meters (fig. 17), which is slightly less than the resistivities of the shallower freshwater-producing aquifers (1,050 to 1,200 ft) in the Potomac aquifer system. Assuming that conditions are similar in the aquifers, the dissolved-solids concentration of the deeper sand (1,410 to 1,460 ft) should be comparable to the concentrations (600 to 750 mg/L) found in the upper sands. The sand that occurs from 1,510 to 1,540 ft has about one-half the resistivity of the upper sands, suggesting a greater concentration of dissolved solids.

The primary purpose for drilling well SO Dd 47 was to test the geothermal potential of the Somerset County area. Figure 18 is a temperature log of SO Dd 47. The geothermal gradient in the Crisfield area is about 2.25 °F per 100 ft to a depth of about 1,000 ft. From 1,000 ft to the top of the basement the gradient is somewhat less at 1.75 °F per 100 ft.

East of Fairmount and Crisfield, the characteristics of the aquifer system are not well known but a test well, 66M23, was drilled in 1987 in the upper sands of the Potomac Group near Jenkins Bridge, Virginia, which is about 14 mi east of Crisfield. The well is 1,298 ft deep and is screened in the interval from 1,288 to 1,298 ft. The chloride concentration in the water from well 66M23 is 1,500 mg/L (A. Meng, U.S. Geological Survey, oral commun., 1988). If the sand of 1,288 to 1,298 ft is correlative with the freshwater-producing sands at Crisfield, it appears likely that water in the Potomae aquifer system becomes brackish east of Crisfield.

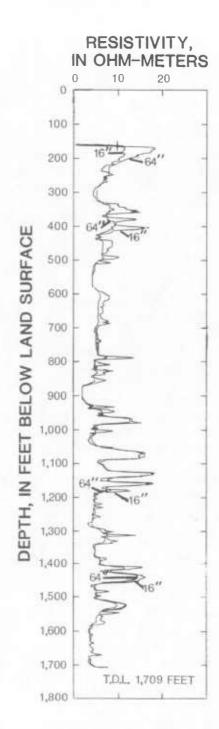


Figure 17.— Electric log for upper part of geothermal test well SO Dd 47 (well location shown on quadrangle map Dd at end of report).

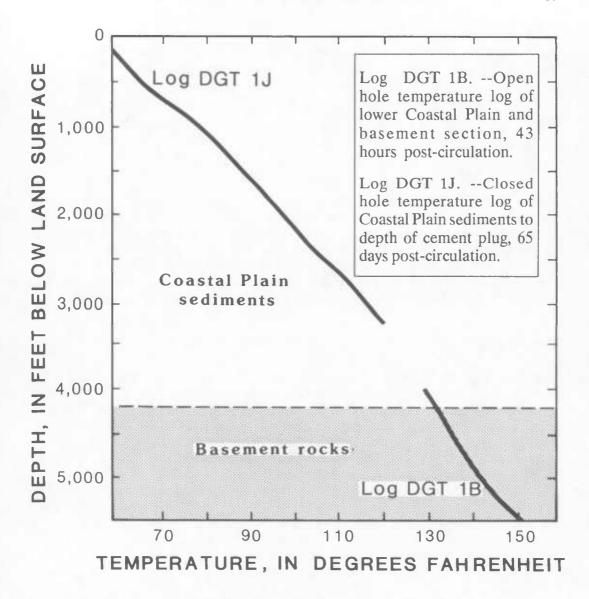


Figure 18.— Temperature profile of geothermal test well SO Dd 47 (modified from Dashevsky and McClung, 1979; well location shown on quadrangle map Dd at end of report).

The upper sands of the Potomac aquifer system are the most heavily used aquifer system for public water supply in the eastern part of the county. They produce water for municipal supplies at Crisfield, Rumbley, Frenchtown, Fairmount, and Smith Island.

Well yield and specific capacity

Reported yields for 17 wells producing from the Potomac aquifer system range from 10 to 500 gal/min. The median well yield is 50 gal/min. In general, wells with lower reported

yields are located on Smith Island and were drilled to supply small water systems. Wells on the mainland of Somerset County that produce from the Potomac aquifer system generally supply larger water systems and have larger reported yields. The range of reported yields for 10 of these mainland wells is from 30 to 500 gal/min, with a median value of 175 gal/min. Reported specific capacity for four wells on the mainland ranges from 1 to 7 (gal/min)/ft.

Hydraulic properties

Because of the complex nature of the Potomac aquifer system, transmissivity values and storage coefficients calculated from aquifer tests may be representative of the tested aquifer only in the vicinity of the test site. In 1970, an aquifer test was performed at Janes Island State Park, using SO Dc 4 as the pumping well and SO Dc 3 as the observation well. Transmissivity of the aquifer (1,100-1,138 ft) is 2,140 ft²/d and the storage coefficient is 0.0002 (table 5). Transmissivity of the aquifer at 1,100-1,140 ft at Fairmount, in well SO Cd 41, is 1,280 ft²/d, as estimated from a 24-hour, single-well aquifer test in 1979 (table 5).

Water levels

Water levels in the Potomac aquifer system at Smith Island were reported as high as 15 ft above land surface in 1953 (Rasmussen and Slaughter, 1955, p. 54). Figure 19 shows mean daily water levels in well SO De 3 at Janes Island State Park during 1985-87. The well is influenced significantly by a nearby pumped well, but the mean daily water level varies from 8 to 12 ft below land surface. Figure 20 shows mean daily water levels during 1986-88 in observation well SO Ec 4, which is screened in the Piney Point aquifer, the Paleocene aquifer system, and the Potomac aquifer system. In this well, water levels are influenced by several production wells screened in the Paleocene and Potomac aquifer systems, rather than by any single well. The mean daily water level was about 18 ft below land surface from February 1986 through March 1987. In March 1987 water levels rose dramatically due to a temporary shutdown of a nearby production well. The generally upward trend from 1987 to 1988 may be due to changes in pumping rates in nearby wells.

Water levels in the Potomac aquifer system in the Crisfield area range from about 8 to 18 ft below land surface. This suggests that greater yields could be achieved from the aquifer system by utilizing more of the 950 ft of available drawdown. However, as with the Paleocene aquifer system, the areal extent and quality of water in the Potomac aquifer system are not well known east of the Crisfield area. It is possible that the aquifer system could not sustain withdrawals greatly in excess of present pumpage or that such pumpage may cause the migration of poor-quality water toward pumping centers.

WATER QUALITY OF THE PRINCIPAL AQUIFERS

Ground-water quality in Somerset County differs considerably, both areally and with depth. Much of the county contains ground water suitable for drinking, but in some areas it is not possible to obtain water that meets drinking-water standards. Chemical analyses were performed for many of the inorganic constituents for which the USEPA has estab-

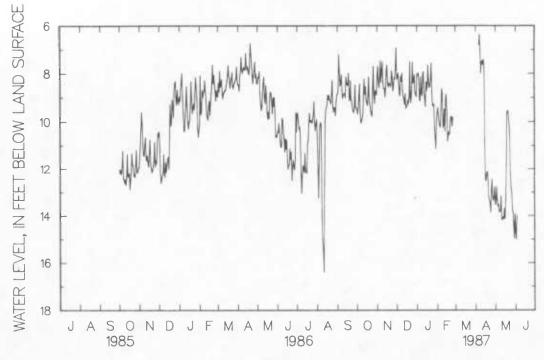


Figure 19.— Water levels in observation well SO Dc 3, 1985-87 (well location shown on quadrangle map Dc at end of report).

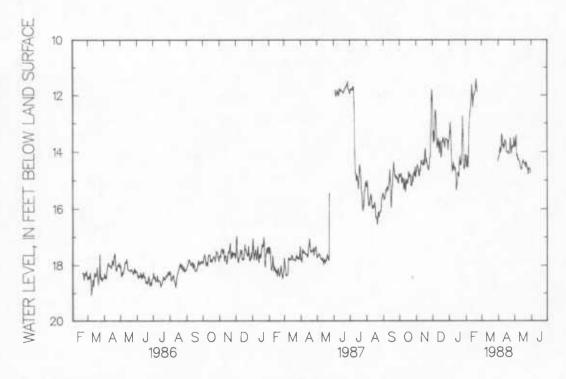


Figure 20.— Water levels in observation well SO Ec 4, 1986-88 (well location shown on quadrangle map Ec at end of report).

lished maximum contaminant levels (MCL). These constituents are listed in table 6. Maximum contaminant levels have been established for those substances that have an associated health risk, whereas secondary maximum contaminant levels (SMCL) are for those substances that primarily affect the aesthetic quality of water (U.S. Environmental Protection Agency, 1986a, 1986b).

Chemical analyses of water from 96 wells in Somerset County are listed in table 13 (at the end of report). Of those wells, 84 were sampled during this investigation and 17 were sampled prior to 1980. Because sample-collection and analytical techniques have changed substantially since 1980, only the recent analyses performed during this investigation (1986-87) are used to describe ground-water quality in Somerset County. Figure 21 shows the location of all sampled wells and the aquifer in which the wells are screened.

Little information exists concerning the quality of water in the Choptank and Pincy Point aquifers. Few wells are screened in these aquifers, and no wells screened entirely in either aquifer are used to supply potable water. Rasmussen and Slaughter (1955, p. 202) report that chloride concentrations in the Choptank aquifer in the Crisfield area exceeded 939 mg/L. The quality of water in the Choptank aquifer in other parts of the county is not well known. A well near Loretto (SO Ae 21) was originally finished in the Choptank aquifer, but was eventually screened in the Manokin aquifer because chloride concentration in water from the Choptank aquifer was 518 mg/L (P. Pryor, Somerset County Health Department, oral commun., 1988). Water from well SO Bb 19, which is finished in the Piney Point aquifer, had acceptable chloride concentrations (180 mg/L), but contained dissolved solids in excess

TABLE 6

MAXIMUM CONTAMINANT LEVELS AND SECONDARY MAXIMUM CONTAMINANT LEVELS FOR SELECTED INORGANIC CONSTITUENTS IN DRINKING WATER (from U.S. Environmental Protection Agency, 1986a, 1986b)

[mg/L = milligrams per liter; MCL = maximum contaminant level; SMCL = secondary maximum contaminant level]

Constituent	Contaminant level (mg/L)	Type of standard
Barium	1.00	MCL
Cadmium	.01	MCL
Chloride	250	SMCL
Copper	1	SMCL
Dissolved Solids		
(total residue)	500	SMCL
Fluoride	4	MCL
Fluoride	2	SMCL
Iron	.30	SMCL
Lead	.05	MCL
Manganese	.05	SMCL
Nitrate (as N)	10	MCL
Sulfate	250	SMCL
Zinc	5	SMCL

of 1,000 mg/L. Rasmussen and Slaughter (1955, p. 202) reported chloride concentrations in water from wells SO Bb 1 and SO Cc 1, which are finished in the Piney Point aquifer, of 250 and 562 mg/L, respectively.

The following discussion of ground-water quality addresses the principal aquifers and aquifer systems in the county; the surficial aquifer system, the Pocomoke aquifer, the Manokin aquifer, the Paleocene aquifer system, and the Potomac aquifer system. Discussion of the Paleocene and Potomac aquifer systems is combined in this section because several wells are screened in both aquifer systems, and water from the two aquifer systems may mix in the area of use.

Surficial Aquifer System

The quality of water in the surficial aquifer system is known only from samples obtained from four wells. The dominant cations in water from wells SO Ce 83 and SO Cf 20 are sodium and iron, the dominant cations in water from SO Bf 20 are calcium and magnesium, and the dominant cation in SO Ae 17 is magnesium. Anion composition also differs in the four analyses. In water from wells SO Bf 20 and SO Cf 20, the dominant anions are sulfate and chloride, in water from well SO Ce 83 the major anions are sulfate and bicarbonate, and in water from SO Ae 17 the major anion is sulfate. Concentrations of dissolved solids range from 108 to 252 mg/L, below the SMCL of 500 mg/L (U.S. Environmental Protection Agency, 1986b). Water from the wells is soft to moderately hard and slightly acidic. The data available are not adequate to describe the quality of water in the surficial aquifer system. The complex geology and ground-water-flow field, and effects of land use, probably cause water quality in the aquifer system to differ considerably from place to place.

The most commonly reported water-quality problem associated with the surficial aquifer system is excessive concentrations of iron (P. Pryor, Somerset County Health Department, oral commun., 1988). Two of the four samples had concentrations of iron greater than the SMCL of 0.3 mg/L (U.S. Environmental Protection Agency, 1977b). Samples from wells SO Ce 83 and SO Cf 20 had iron concentrations of 48 mg/L and 27 mg/L, respectively. Iron is a common constituent in anoxic ground water in the surficial aquifers on the Delmarva Peninsula. In Delaware, Denver (1986) noted that dissolved iron is a significant component of shallow ground water associated with poorly drained soils and in water from wells screened near the base of the unconfined aquifer. Similar conditions probably exist in Somerset County. Iron also may be present in high concentrations in oxygenated water when unstable conditions exist in a well or aquifer. Boggess and Heidel (1968, p. 23) observed high concentrations of iron in oxygenated water in the Salisbury area and attributed these high concentrations to unstable conditions in the aquifer.

The surficial aquifer system also is susceptible to nitrate contamination. Nitrate concentrations in ground water that approach or exceed 10 mg/L as nitrogen usually are derived from nitrogen fertilizers or animal wastes. Records from the Somerset County Health Department show that elevated nitrogen concentrations occur chiefly in shallow ground water underlying well-drained soils. Figure 22 shows areas of the county underlain by well-drained soils. In these areas, shallow ground water in the vicinity of nitrogen sources may be susceptible to nitrate contamination.

The surficial aquifer system comprises both unconfined and confined aquifers. Unconfined aquifers generally are more susceptible to contamination from land-usc practices than

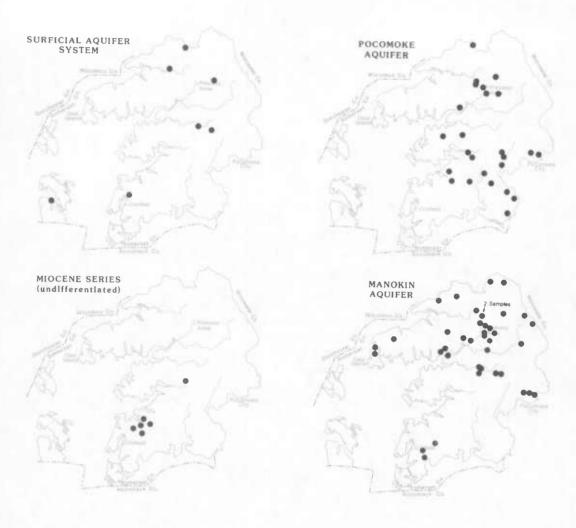
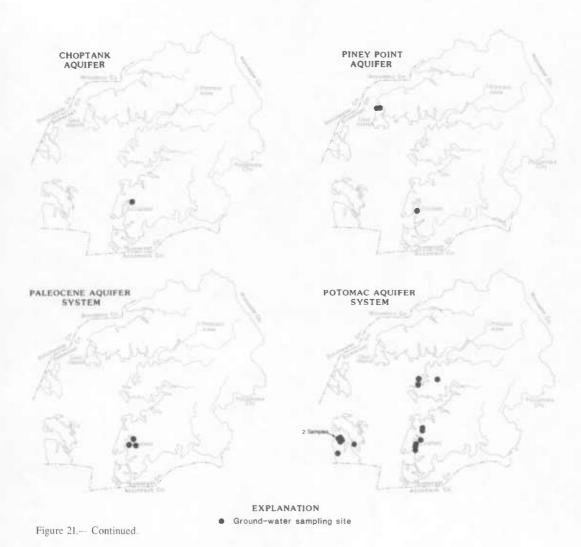


Figure 21.— Location of ground-water sampling sites.

confined aquifers because they are not overlain by low permeability units that inhibit the downward migration of contaminants. In 1988, Somerset County was considering regulations that would require new wells to have a minimum depth of 50 ft (P. Pryor, Somerset County Health Department, oral commun., 1988), because of the susceptibility of unconfined aquifers to contamination. Throughout much of the county the thickness of the surficial aquifer system is less than 50 ft; therefore, new wells drilled under the proposed regulation would be cased through the surficial aquifer system and be screened in deeper, confined aquifers.

Pocomoke Aquifer

The quality of water in the Pocomoke aquifer is quite variable, as shown by the 24 chemical analyses in the trilinear diagram of figure 23. Overall, there appears to be little similar-



ity between the analyses. However, when the analyses are grouped by hydrogeologic characteristics, similarities can be seen. The Pocomoke aquifer in recharge areas is generally unconfined and water chemistry is influenced chiefly by the composition of precipitation, aquifer mineralogy, land use, soil type, and position in the ground-water-flow system. Therefore, chemical data for water from the recharge areas are scattered on figure 23. In contrast, analyses of water from the confined parts of the aquifer are grouped more closely on figure 23, probably due to the influence of mineral dissolution on water chemistry. The dominant cations in water from the confined part of the Pocomoke aquifer are calcium and sodium and the dominant anion is bicarbonate.

Plate 8 presents Stiff diagrams of major ions in 24 water samples from the Pocomoke aquifer. Generally, water is least mineralized in the recharge areas. The amount of dissolved constituents, especially calcium and bicarbonate, increases as water moves from subcrop areas to the confined parts of the aquifer. The increase in calcium and bicarbonate probably



Figure 22.— Areas underlain by well-drained soils (adapted from Maryland Department of State Planning, 1974).

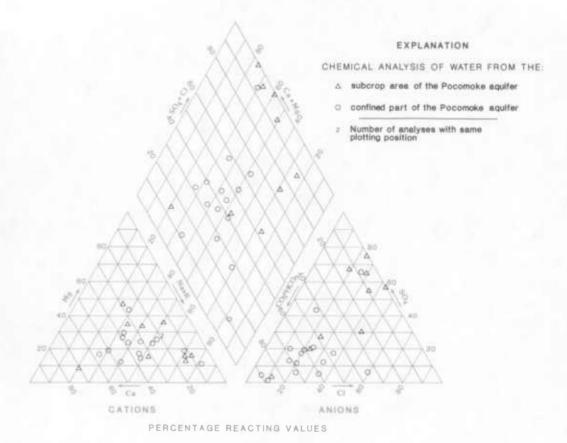


Figure 23.— Tri inear diagram for water from the Pocomoke aquifer.

is due to the dissolution of ealeite or aragonite, which are the principal minerals dissolved by ground water on the Delmarva Peninsula (Cushing and others, 1973, p. 7).

Figure 24 shows box plots of the distribution of pH, hardness, and dissolved-solids concentration of water from the Poeomoke aquifer. The pH of water from recharge areas ranges from 4.5 to 6.4, and has a median of 5.1. In the confined part of the aquifer, pH of the water tends to be higher, ranging from 5.1 to 7.3, and has a median of 6.4 (fig. 24). Water from the confined part of the aquifer tends to be harder than water from recharge areas. Hardness of water from recharge areas ranges from 11 to 100 mg/L (as calcium carbonate) and has a median value of 33 mg/L. In the confined part of the aquifer hardness of the water ranges from 22 to 410 mg/L, and has a median value of 140 mg/L (fig. 24). Dissolved-solids concentrations of water from recharge areas range from 90 to 177 mg/L, and have a median of 131 mg/L. In water from the confined part of the aquifer, dissolved-solids concentrations range from 116 to 1,440 mg/L, and have a median of 287 mg/L (fig. 24). Overall, water in the recharge areas is more acidic, softer, and contains lower concentrations of dissolved solids than water in confined parts of the aquifer. These distributions likely reflect the longer flow paths in the confined system and the longer contact time between water and minerals of the aquifer matrix.

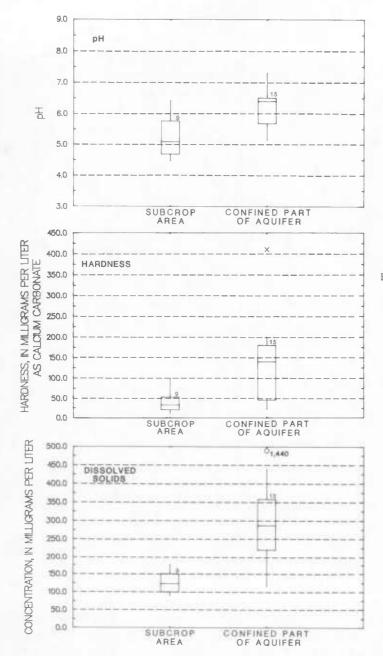
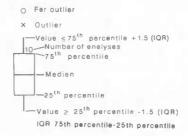


Figure 24.— pH, hardness, and dissolvedsolids concentrations of water from the Pocomoke aquifer.

EXPLANATION



The most common quality problems in water from the Pocomoke aquifer are iron and manganese concentrations in excess of the SMCL's of 300 and 50 μ g/L (micrograms per liter), respectively (U.S. Environmental Protection Agency, 1986b). Iron concentrations range from 50 to 41,000 μ g/L with a median of 5,000 μ g/L. Manganese concentrations range from 12 to 900 μ g/L, with a median of 110 μ g/L. Water from 22 of the 24 samples exceeds the SMCL for iron and 18 exceed the SMCL for manganese. One sample, from well SO Ef 6, exceeds the SMCL's for chloride and dissolved solids. This well is located near the Pocomoke River, and may be receiving brackish water from the river. Although no samples exceeded the primary drinking water standard for nitrate, water from two wells, SO Ae 16 and SO Df 14, contains nitrate concentrations of 6.3 and 6.5 mg/L respectively, which are greater than background (table 13, at end of report).

Manokin Aquifer

Thirty-seven water-quality samples were collected from wells in the Manokin aquifer during this investigation. The distribution of major ions in the water samples is shown in figure 25. The dominant cation is sodium in every sample but one (from well SO Af 20), and the dominant anions are bicarbonate and chloride.

There is a marked areal difference in the chemical quality of water from the Manokin aquifer. North of Westover, the water is a sodium-bicarbonate type, dissolved solids range from 173 to 620 mg/L, and hardness ranges from 2 to 97 mg/L. South of Westover, the water is a sodium-chloride type, dissolved solids range from 807 to 1,860 mg/L, and hardness ranges from 76 to 260 mg/L. Stiff diagrams of the major ions in 16 water samples from the aquifer are presented in plate 9.

Back (1966, p. A38) proposed two explanations for sodium being the dominant cation in ground water in the Coastal Plain scdiments of Maryland and Virginia: (1) saltwater underlies the area; and (2) cation exchange occurs between calcium in ground water and sodium in clay minerals. Ground water in the Manokin aquifer in Somerset County may obtain sodium by both mechanisms. In the northeastern corner of the county, chemical analysis of water from well SO Af 20 shows that the major cations are calcium and sodium. Toward Westover, sodium becomes the dominant cation, calcium becomes subordinate, and bicarbonate concentration increases. This may be the result of mineral dissolution, which produces calcium and bicarbonate ions, and the available calcium exchanging with sodium from clay minerals as the water moves through the aquifer. South of Westover, both sodium and chloride concentrations increase, probably due to the presence of brackish water in the aquifer.

The areal distribution of chloride in the Manokin aquifer is shown in figure 26, which was constructed using data from water samples collected for this investigation during the summer of 1986 and 1985-86 chloride data from files of the Somerset County Health Department (table 7). From the northern part of the county, chloride concentrations gradually increase southwestward to the vicinity of Westover. Here, perpendicular to a line that roughly trends from Pocomoke City to Deal Island, the concentration gradient is steep, with chloride concentration increasing from 150 mg/L to more than 500 mg/L in a distance of about 2 mi. Chloride concentrations are not known in the central part of the county because few wells are drilled in the Manokin aquifer there. Local drillers, however, report the water in this area as being "very salty." Therefore, it is likely that chloride concentrations exceed 500 mg/L. In the Crisfield area, chloride concentrations decrease slightly. Here the

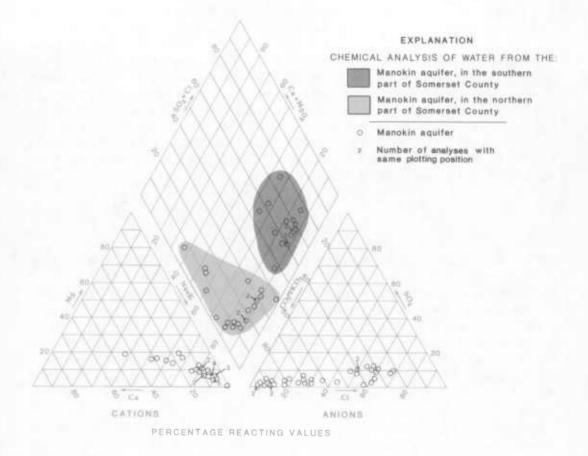


Figure 25.— Trilinear diagram for water from the Manokin aquifer.

eonfining units overlying the Manokin aquifer are more permeable than in other parts of the eounty, and water containing lower concentrations of chloride may migrate downward from overlying aquifers. A short distance from Crisfield, however, chloride concentrations approach 500 mg/L.

There are several possible explanations for the occurrence of water containing high levels of chloride in the Manokin aquifer. One explanation is that the lowering of hydraulie head in the Princess Anne area has resulted in the migration of brackish water from the Chesapeake Bay toward pumping centers. If this were the case, analyses for samples collected in the area south of Westover prior to ground-water development should have eonsiderably lower chloride concentrations than recent analyses. In fact, analyses from the 1950's, when the principal direction of ground-water flow was toward the Chesapeake Bay, also show high chloride concentrations in the area south of Westover (Rasmussen and Slaughter, 1955, p. 202-203). Therefore, it is doubtful that the high chloride concentrations have resulted from ground-water pumpage.

Other explanations for the observed ehloride concentrations include incomplete flushing of the aquifer and the presence of a transition zone. Because the aquifer subcrops beneath Chesapeake Bay, an interface between freshwater in the aquifer and brackish water of the

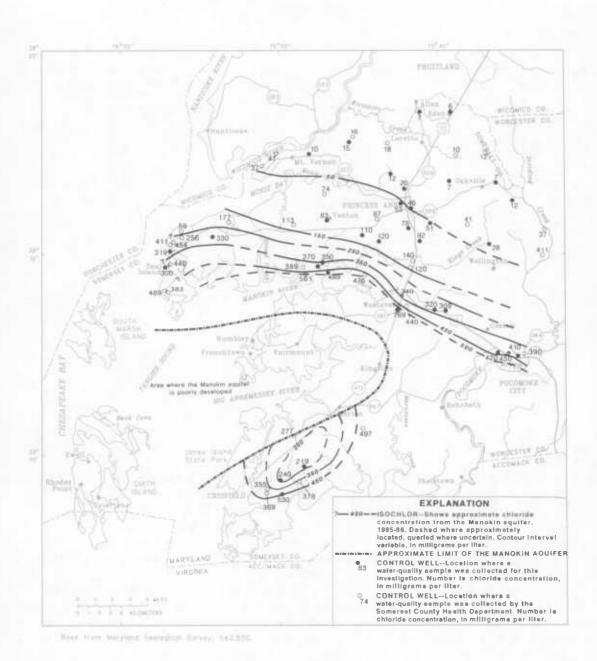


Figure 26.- Chloride concentration in water from the Manokin aquifer, 1985-86.

TABLE 7
CHLORIDE CONCENTRATIONS IN WATER FROM WELLS SAMPLED BY THE SOMERSET COUNTY HEALTH DEPARTMENT
[ft = feet; mg/L = milligrams per liter]

Well no.	Well depth (ft)	Chloride concentration (mg/L)
SO Ad 15	150	16
SO Ae 21	160	18
SO Bb 20	150	454
SO Bb 21	140	58
SO Bb 22	160	256
SO Bb 24	150	411
50 Bc 17	120	177
SO Bc 18	140	47
SO Bc 19	125	37
50 Bd 43	150	113
SO Bd 46	160	74
O Be 107	190	87
SO Bf 22	230	41
SO Bf 24	240	10
SO Bf 25	235	12
SO Bg 5	260	37
SO Cb 25	160	383
50 Cb 26	140	489
SO Cb 27	140	440
SO Cd 53	180	389
SO Cd 54	160	561
SO Ce 44	240	140
SO Ce 48	230	120
SO Ce 90	230	440
SO Ce 94	185	476
SO Dd 64	185	277
SO De 42	210	497
SO Dg 10	220	390
SO Ec 53	200	369
SO Ec 55	200	355
SO Ed 48	205	376

bay would develop naturally. This interface, rather than being a sharp boundary, is more likely to be a zone of diffusion where brackish water and freshwater mix. Possibly, the natural position of this zone is in the southern part of the county. Zones of incomplete flushing may occur north of Crisfield, where the aquifer becomes less permeable. This low-permeability area north of Crisfield restricts ground-water flow, which may result in localized stagnation zones and areas where saltwater trapped during times of higher sea level has not been completely flushed from the system.

Although the zone of water containing high concentrations of chloride may have resulted from natural processes, it has the potential to migrate toward pumping centers at Princess Anne. In the early 1950's, regional ground-water flow in the Manokin aquifer was from northeast to southwest, and thus stabilized the poorer-quality water in the vicinity of Westover. Ground-water flow in the county is presently toward Princess Anne, and water containing elevated chloride concentrations may be migrating toward the Princess Anne area. The rate of this possible migration is addressed in later sections of this report.

Dissolved iron is a problem in the Manokin aquifer in the northern part of the county, where concentrations in water samples reach 4.1 mg/L. Figure 27 shows dissolved-iron concentrations at sampled locations. Iron concentrations are generally greatest in the northeastern corner of the county and progressively decrease toward the south, in the direction



Figure 27.— Iron concentration in water from the Manokin aquifer, 1986.

of pre-development ground-water flow. Conversely, pH tends to increase in the direction of pre-development ground-water flow (northeast to west and southwest), probably because bicarbonate is brought into solution from mineral dissolution (Cushing and others, 1973, p. 9). Figure 28 is a graph of pH versus total iron concentration and shows a general trend of decreasing iron concentration with increasing pH. Cushing and others (1973, p. 9) noted a downgradient decrease in iron concentration in confined aquifers in other parts of the Delmarva Peninsula. They attributed this decrease to increasing pH along a flow path, causing precipitation of siderite. Langmuir (1969) observed a downgradient decrease in iron concentrations in water from Cretaceous-age sediments in New Jersey. He explained the decrease as a result of increasing stability of suspended amorphous material due to aging, coupled with adsorption of ferrous iron by oxyhydroxides at pH greater than 6.5.

Paleocene and Potomac Aquifer Systems

The Paleocene and Potomac aquifer systems contain the deepest aquifers used in Somerset County. For this investigation, one water-quality sample was collected from a well screened in the Paleocene aquifer system (SO Ec 1), one sample was collected from a well screened in both aquifer systems (SO Ec 49), and eight samples were collected from wells in the Potomac aquifer system. Several wells in the Crisfield area are screened in both the Paleocene and Potomac aquifer systems. The multiple screens in these wells may allow water in the two aquifer systems to mix.

Water in the two aquifer systems is a sodium-bicarbonate type, with sodium accounting for more than 95 percent of the cations (fig. 29). Alkalinities are the greatest of any of the aquifers in Somerset County (fig. 30). The dominance of sodium and bicarbonate is probably best explained by ion-exchange processes (Foster, 1950). As ground water moves downgradient it becomes enriched in calcium and bicarbonate through mineral dissolution. As water continues in the flow system, calcium is depleted as it is exchanged for sodium in clay minerals. Because calcium is depleted, the ground water cannot achieve equilibrium with calcite and more bicarbonate is introduced into the ground water, resulting in a sodium-bicarbonate type water.

Minerals present in the two aquifer systems that may supply calcium and bicarbonatc include calcite, aragonite, and silicate minerals (principally feldspar). Because the Potomac Group is nonmarine, the principal sources of calcium and bicarbonate are probably silicate minerals.

Figure 31 shows Stiff diagrams depicting the distribution of major ions in water from 10 wells in the two aquifer systems. In the area of use, the composition of the water is relatively uniform. Some samples contain greater concentrations of chloride, but no areal pattern is apparent. Overall, the water in these two aquifer systems is soft, has dissolved-solids concentrations ranging from 475 to 1,070 mg/L, and has the highest pH of any aquifer in the county (fig. 32).

The major water-quality problem in the Paleocene and Potomac aquifer systems is elevated concentrations of fluoride (Kula and Hansen, 1989). Seven water samples collected for this study exceed the SMCL of 2.0 mg/L and two of these exceed the MCL of 4.0 mg/L (U.S. Environmental Protection Agency, 1986a). Fluoride occurs in accessory minerals in most sediments, although the low solubility of these minerals usually causes fluoride concentrations to be less than 1.0 mg/L (Cushing and others, 1973, p. 8). Primary

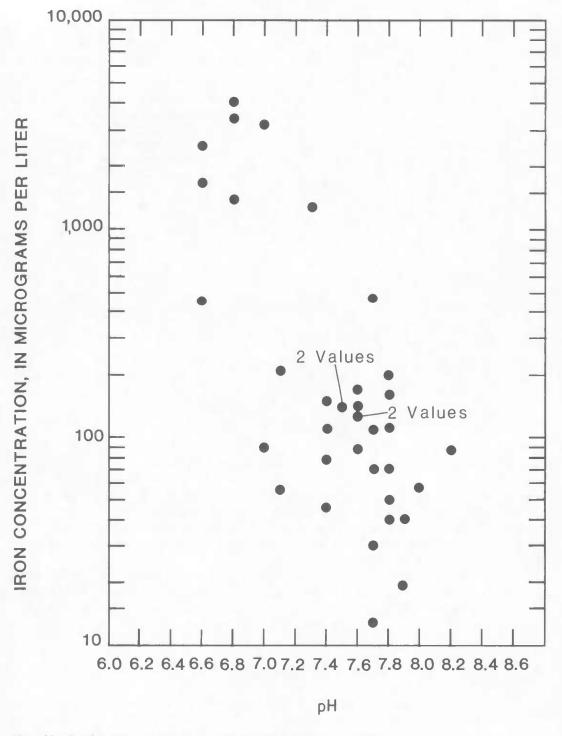


Figure 28.— Relation of iron concentration to pH in water from the Manokin aquifer.

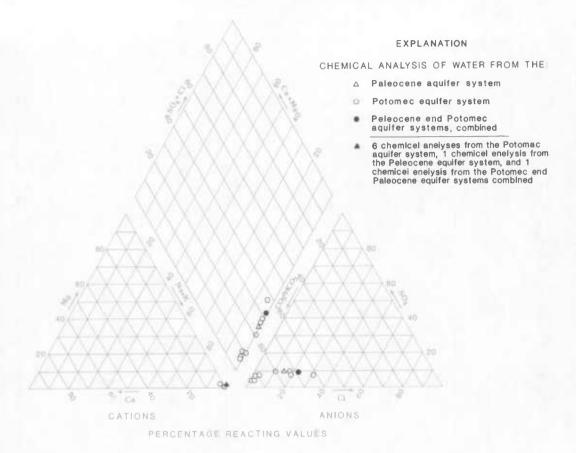


Figure 29.— Trilinear diagram for water from the Paleocene and Potomac aquifer systems.

sources of fluoride are fluorite and apatite (Hem, 1985, p. 121). It is also a minor component in such hydroxide-bearing minerals as biotite, museovite, kaolinite, and horneblende. The relatively high concentrations of fluoride in water in the Paleocene and Potomac aquifer systems may be due to ion exchange between hydroxide and fluoride ions, or to changes in the mineralogy of the sediments.

The water quality in the area of use does not appear to have changed appreciably since the 1950's. However, the position of the saltwater interface in the Paleocene and Potomac aquifer systems east of Crisfield is not known. Because the hydraulic heads in the aquifer systems have been lowered due to pumping, the interface may be migrating westward.

EFFECTS OF GROUND-WATER DEVELOPMENT

Ground-water demand in Princess Anne and Crisfield (fig. 33), where most pumping is eurrently located, is expected to increase in the near future. Meeting this demand will involve putting new production wells into service, as well as increasing the pumping rate of existing wells. The purpose of this section is to describe the probable effects of additional withdrawals in these two areas on the ground-water-flow system.

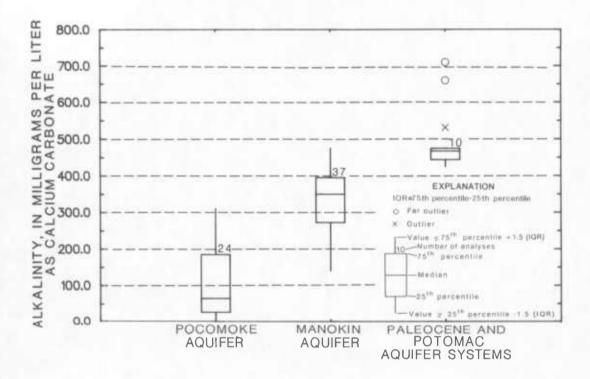


Figure 30.— Alkalinity of water from the principal aquifers.

Effects of Projected Ground-Water Pumpage in the Princess Anne Area

The Manokin aquifer is the principal source of ground water in the Princess Anne area. Users of more than 10,000 gal/d include the town of Princess Anne, a plywood-manufacturing plant, a poultry-raising operation, and the Eastern Correctional Institution. Ground-water-level altitudes in the area range from about 5 ft above sea level to about 20 ft below sea level. Ground-water pumpage is expected to increase by about 600,000 gal/d by the year 2000 (J. Windsor, Somerset County Department of Technical and Community Services, written commun., 1986), primarily to meet the water demands of the Eastern Correctional Institution. Water levels are expected to decline further as a result of the increased withdrawals. The present ground-water-flow system and the effects of increased withdrawals from the Manokin aquifer on the ground-water-flow system were evaluated with a digital ground-water-flow model.

Model description and grid

This investigation used the U.S. Geological Survey's modular three-dimensional finite-difference ground-water-flow model (McDonald and Harbaugh, 1988) to simulate flow in the Manokin aquifer in the Princess Anne area. The model was used to simulate steady-state conditions (that is, water levels and fluxes do not change with time). Simulations of transient conditions were not conducted for the following reasons:

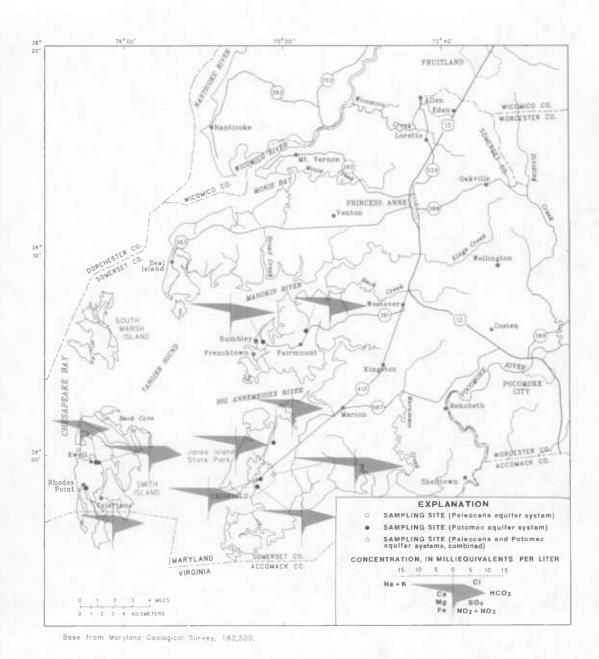


Figure 31.— Chemical characteristics of water from the Paleocene and Potomac aquifer systems.

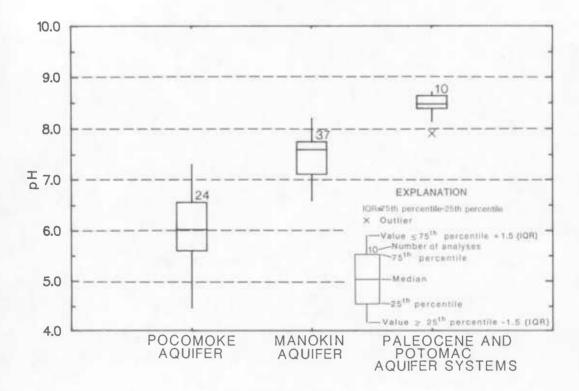


Figure 32.—pH of water from the principal aquifers.

- (1) The purpose of the modeling was not to predict water levels at specific times, but to estimate the long-term effects of increased withdrawals; and
- (2) data such as ground-water withdrawals over time and long-term water-level measurements that are necessary to properly calibrate the model under transient conditions were unavailable.

Briefly, the model operates by using finite-difference approximations of the partial-differential equation that describes ground-water flow. Instead of being treated as a single continuous system, the ground-water-flow system is represented by a grid of rectangular blocks. Each block, or cell, is considered to have uniform properties, and a finite-difference approximation is formulated for every cell. The unknown variable, hydraulic head, is found by solving the finite-difference approximation by an iterative procedure. If the inputs to the model are of sufficient accuracy, the model-generated water levels will compare favorably to observed water levels. The model is then considered calibrated. The calibrated model can be used to evaluate the response of the ground-water-flow system to imposed stresses, such as increased ground-water withdrawals. McDonald and Harbaugh (1988) present a detailed description of the features and mathematical development of the model used in this investigation.

The grid representing the ground-water-flow system in the Princess Anne area eonsists of two layers, each containing 57 rows and 47 columns of cells. The size of the grid cells differs, ranging from 1,000 ft on a side near pumping wells to 8,000 ft on a side at the boundaries of the model. The grid (fig. 34) is larger than the area of interest (fig. 33) because

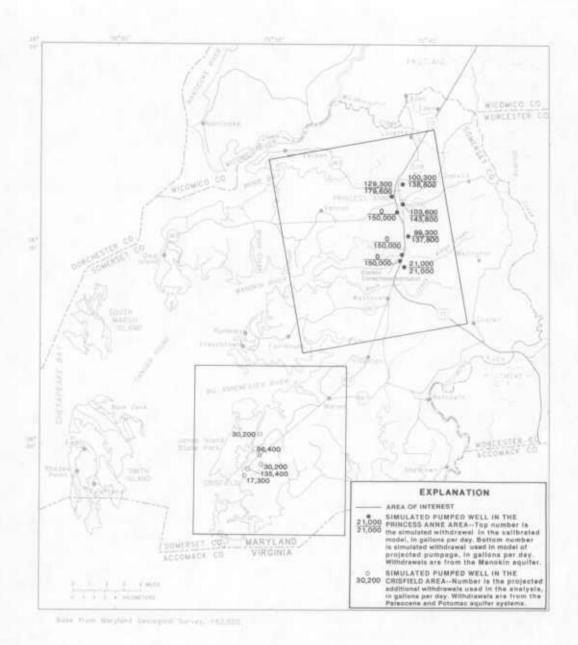


Figure 33.— Location of modeled areas and simulated pumpage distributions.

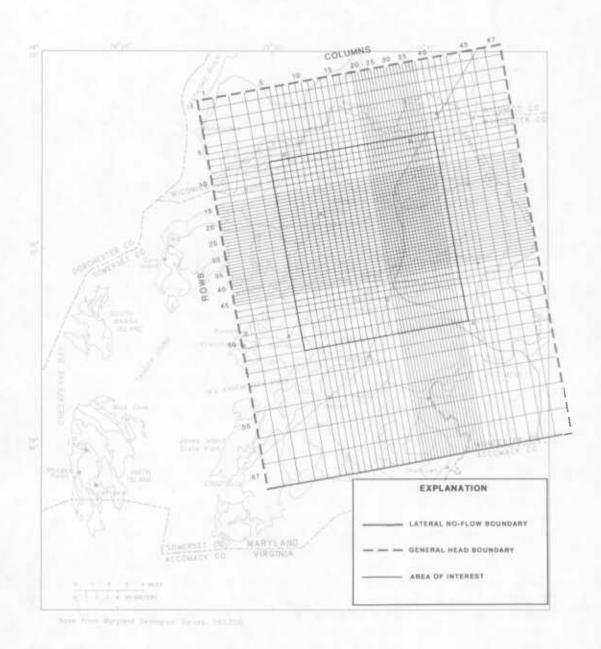


Figure 34.- Model grid for the Princess Anne area.

water levels within the area of interest are affected by conditions beyond it. By convention, the numbering system for the rows and columns begins in the upper left corner of the grid. Each cell is referenced by layer, row, and column. For example, cell 2, 10, 20 is located in the second layer, at the intersection of row 10 and column 20.

Conceptual model and boundary conditions

To make the ground-water-flow system amenable to mathematical analysis, geologic and hydrologic conditions in the study area must be simplified into a conceptual model. The conceptual model for the Princess Anne area (fig. 35) consists of the following characteristics and assumptions:

- (1) The aquifers under consideration are represented by two layers. The upper layer (layer 1) comprises the surficial aquifer system and the Pocomoke aquifer. The lower layer (layer 2) consists of the Manokin aquifer;
- (2) the confining unit overlying the Manokin aquifer is not modeled as an active layer. Leakance is not calculated by the model, but is determined independently by dividing vertical hydraulic conductivity of the confining unit by the thickness of the confining unit. These values are then supplied to the model for calculation of vertical flow through the confining unit;
- (3) ground-water flow in the Manokin aquifer is horizontal;
- (4) water levels in the surficial aquifer system and the Pocomoke aquifer are constant, thus the top layer is not an active layer;
- (5) ground water may flow laterally across the northern, eastern, and western boundaries of the modeled area. This lateral flow is simulated by using the general-head-boundary package in the model (McDonald and Harbaugh, 1988);
- (6) the confining unit underlying the Manokin aquifer contributes negligible amounts of water to the aquifer and is represented as a no-flow boundary; and
- (7) the ground-water-flow system is at steady state.

Several components of the conceptual model require further explanation. The generalhead-boundary package simulates a source of water outside the modeled area that supplies water to a cell at a rate dependent on the water-level difference between the source and the cell (McDonald and Harbaugh, 1988, p. 11-1). The water level at the source is assigned as a constant value and the aquifer material between the source and the cell is represented by a hydraulic conductance (transmissivity of the material multiplied by width of the cell divided by distance to the source from the cell). Sources of water to the boundary cells along the western and northern sides of the model are the subcrop areas of the Manokin aquifer, which are located outside the modeled area. The source of water for the castern boundary cells is an area of relatively high heads east of the modeled area that is indicated on published maps of the Manokin aquifer (Cushing and others, 1973; Hodges, 1984). Water levels were assigned to the sources based on published reports (Bachman and Wilson, 1984; Hodges, 1984) and data from observation wells. Hydraulic conductances were assigned on the basis of information from published reports (Rasmussen and Slaughter, 1955; Achmad and Weigle, 1979; and Hodges, 1984). Table 8 lists the water levels and hydraulic conductances used in the model. Little is known about water levels to the south of the modeled area. The southern model boundary is located far enough away that the boundary does not greatly affect heads in the area of interest; it was simulated as a no-flow boundary.

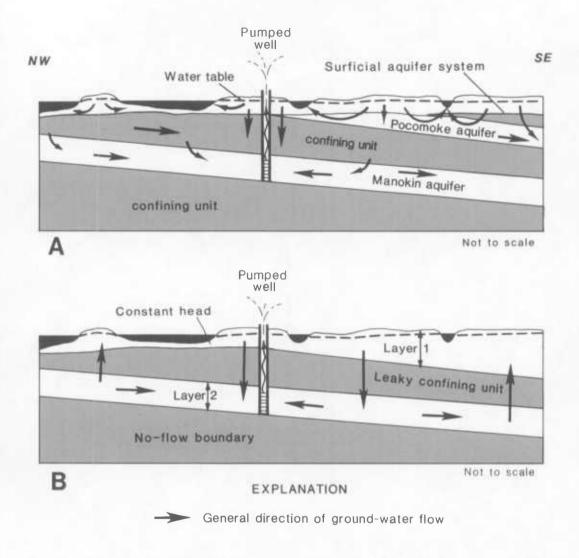


Figure 35.— Schematic representation of the actual ground-water-flow system in the Princess Anne area (A), and the simplified conceptual model (B).

The other component of the conceptual model requiring further explanation is the assumption that ground-water flow is at steady-state conditions. Although figure 15 indicates that the general decline of water level in well SO Be 42 was much less from 1983 to 1986 than in the 5 previous years, there is still a slight downward trend in the 1983-86 data. If the water levels to which the steady-state model is calibrated are under transient rather than steady-state conditions, the aquifer properties derived from the calibration may be inaccurate.

Figure 36 shows continuous water levels in observation well SO Ce 42 and periodic water levels in observation well SO Cf 2 for January 1986 to June 1987. SO Ce 42 is screened in the Manokin aquifer, about 2 mi southwest of the nearest production well, whereas SO Cf

TABLE 8
WATER LEVELS AND HYDRAULIC CONDUCTANCES USED FOR GENERAL-HEAD BOUNDARIES
IN THE MODEL OF THE PRINCESS ANNE AREA
[ft = feet above sea level; ft²/d = feet squared per day]

Model layer	Row	Column	Water level (ft)	Hydraulic conductance (ft ² /d)	Model layer	Row	Column	Water level (ft)	Hydraulic conductance (ft ² /d)
16 TH 26 TH 16 TH	11111111111111111111111111111111111111	1234567890123426789012345678901234567840123456777777777777777777777777777777777777	01112345555555555555555555555555555555555	1.5005500550055550055500555005550055500		3456789012345678901234234567890124567890124567890124567890124567890124567890100000000000000000000000000000000000	777777777777777777777777777777777777777	1555555555560005555444000000000000000000	80 80 85 85 85 80 80 80 80 80 80 80 80 80 80 80 80 80

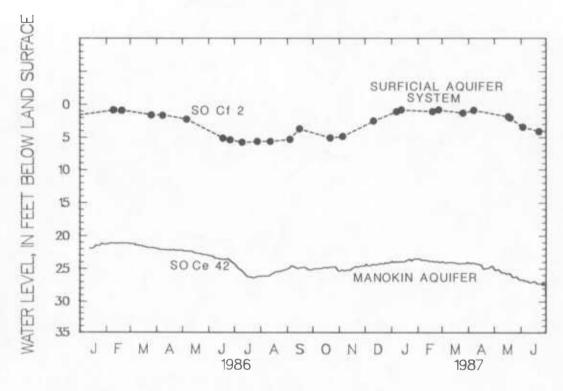


Figure 36.— Water levels in observation wells SO Ce 42 and SO Cf 2, 1986-87 (well locations shown on quadrangle maps Ce and Cf at end of report).

2 is screened in the surficial aquifer system and is not affected by nearby pumping. Scasonal fluctuations of water levels occur in each well, which indicates that some of the fluctuations in SO Cc 42 may be caused by natural conditions. Scasonal fluctuations in water levels due to natural causes tend to be cyclical, so that the system may approximate steady-state conditions over the long term. If this is the case in the Princess Anne area, the aquifer properties obtained by calibrating the model to water levels midway between the annual high and low water levels may be appropriate. Figure 37 shows water levels in SO Ce 42 and total monthly pumpage of the production wells at Princess Anne. Water levels in the observation well do not directly reflect changes in monthly pumpage, but there is a general downward trend in water levels, and a general increase in total monthly pumpage at Princess Anne. This indicates that the system probably is not at steady-state conditions. The effects of inaccurate aquifer properties are more fully addressed in the sections describing results and sensitivity analysis.

Data requirements

The data necessary to model the Princess Anne area are aquifer transmissivity, vertical hydraulic conductivity and thickness of the confining unit overlying the Manokin aquifer, and ground-water-withdrawal rates from the Manokin aquifer. Initial estimates of aquifer transmissivity for the Manokin aquifer were determined by multiplying the thickness of the

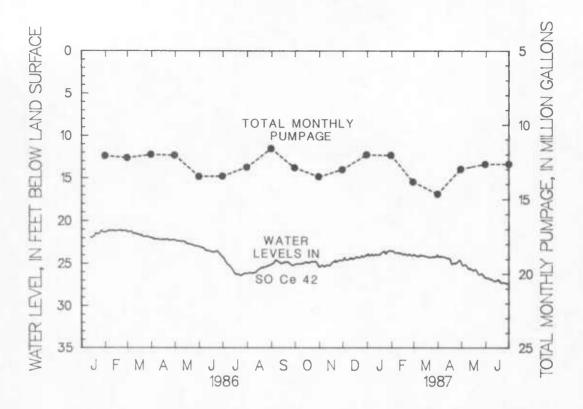


Figure 37.— Water levels in observation well SO Ce 42, and total monthly pumpage at Princess Anne, 1986-87 (well location shown on quadrangle map Ce at end of report).

aquifer (fig. 12) by an assumed constant hydraulic conductivity of 13.2 ft/d. During calibration, transmissivities were adjusted to the distribution shown in figure 13. Vertical hydraulic conductivity of the confining unit overlying the Manokin aquifer in Somerset County is not known. In Salisbury, Wolff (1970, p. 202) measured vertical hydraulic conductivity of between 2.8×10^{-5} and 5.7×10^{-3} ft/d for Miocene clays. In Ocean City, Achmad and Weigle (1979, p. 11) used a model-derived value of 1.9×10^{-3} ft/d. The initial uniform value used in the Princess Anne area model was 1×10^{-4} ft/d. This was adjusted during calibration to a final uniform value of 1×10^{-5} ft/d. Thickness of the confining unit was obtained from figure 10. Withdrawal rates used in the model were the average reported pumpage for the town of Princess Anne from November 1985 to November 1986, and estimated pumpage for the other large users.

Model calibration

The model was calibrated to water levels measured in the Manokin aquifer during the last week of November 1986. Water levels during this time were between the low water levels in July 1986 and the high water levels of January 1987 (fig. 36), and were thought to best approximate steady-state conditions.

During calibration, transmissivity of the Manokin aquifer (layer 2), vertical hydraulic conductivity of the overlying confining unit, and the hydraulic conductances and water

levels of the general-head boundaries were adjusted independently until model-generated heads agreed approximately with measured water levels. The amount each input was adjusted from the initial value was dependent on the reliability of the original estimate. For example, transmissivity of the Manokin aquifer (layer 2) is relatively well known in the area of interest. Therefore, transmissivity was not allowed to differ greatly from the initial value. Accordingly, vertical hydraulic conductivity of the confining unit above the Manokin was allowed to range two orders of magnitude, and the water levels and hydraulic conductances in the general-head boundaries were allowed to differ by as much as 100 percent.

Calibration of the model was aided by comparing differences between observed and computed water levels and by minimizing the root mean square error (RMSE) of the differences between the observed and simulated water levels. The differences were compared to determine if the model was underestimating or overestimating water levels. The ealibrated model overestimated water levels at 12 locations and underestimated water levels at 9 locations. The RMSE was calculated using the following formula:

$$RMSE = \left[\frac{\sum (h_m - h_x)^2}{N}\right]^{1/2}$$
 (4)

where: N = the number of measured water levels:

 h_m = the water level measured in an observation well; and

 h_s = the simulated water level at the eenter of the eell that eontains the observation well.

The calibrated model had a RMSE of 1 ft, with a maximum difference of 2 ft between observed and simulated water levels. Because many of the observation wells are grouped in specific locations, greater emphasis was placed on matching the areal distribution of observed water levels than on minimizing the RMSE.

Figures 38 and 39, respectively, show the model-generated water levels from the calibration simulation, and the water levels measured in the Manokin aquifer in November 1986. In general, water levels agree satisfactorily in the area of interest and differ somewhat in other parts of the modeled area.

The water budget generated by the model shows that 74 percent of the water entering the Manokin aquifer (layer 2) is derived from the general-head boundaries and 26 percent is from downward leakage through the overlying confining unit. Ground-water pumpage accounts for 97 percent of the water leaving the Manokin aquifer (pumped layer). About 1 percent leaves through upward leakage to the surficial aquifer system and Pocomoke aquifer (layer 1), and 2 percent leaves laterally to general-head boundaries.

Sensitivity analysis

Scnsitivity analysis involves changing one model input during a simulation, while keeping all other inputs constant. This gives an indication of which input most affects model results. Sensitivity analysis was performed on the calibrated steady-state model. During the sensitivity analysis, model inputs were changed by amounts inversely proportional to the reliance of their values. For example, vertical hydraulic conductivity of the upper confining

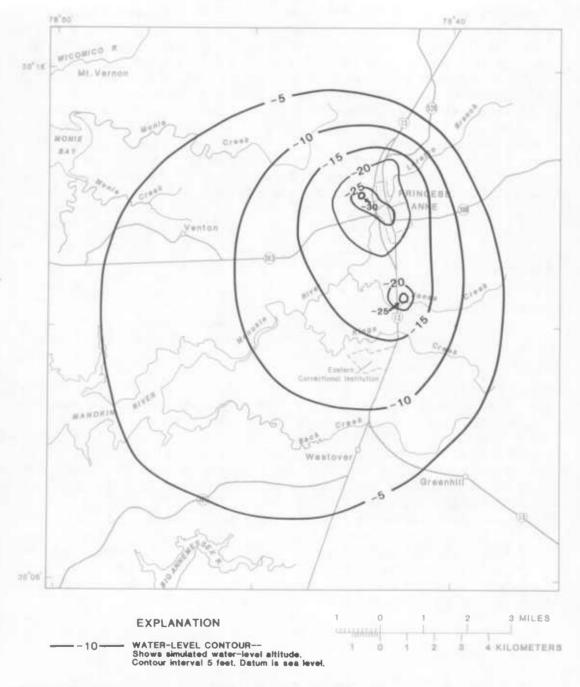


Figure 38.— Simulated steady-state water-level altitudes in the Manokin aquifer using average daily pumpage from November 1985 to November 1986.

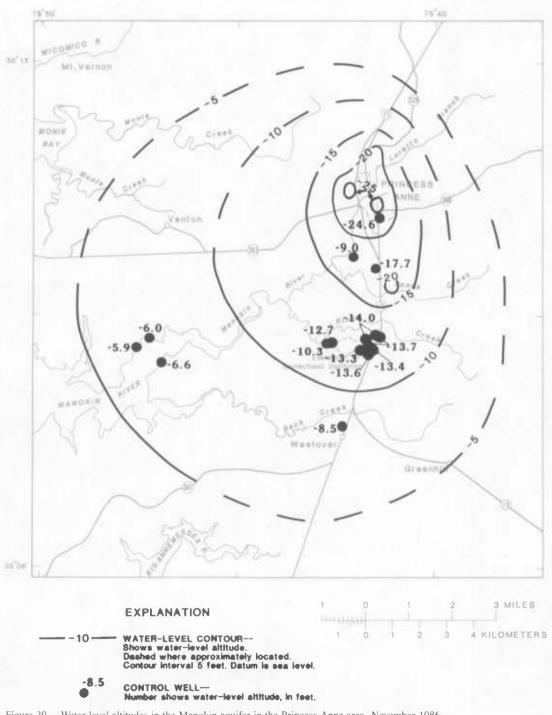


Figure 39.— Water-level altitudes in the Manokin aquifer in the Princess Anne area, November 1986.

unit, which was largely model derived, was changed by an order of magnitude, but transmissivity of the Manokin aquifer (layer 2) was changed by only 50 percent.

Table 9 lists the model inputs that were changed, the amount of variation, and the approximate water-level change in the area of interest resulting from the input change. The inputs that had the greatest effect on model results were vertical hydraulic conductivity of the confining unit and transmissivity of the Manokin aquifer (layer 2). Changes in these inputs resulted in water-level changes of up to 12 ft. The water-level changes generally were greatest near simulated pumping wells. The input that had the least effect on model results was the constant head specified for the surficial aquifer system and Pocomoke aquifer (layer 1); the magnitude of the water-level change was less than 2 ft.

Water levels and hydraulic conductances in the general-head boundaries also were varied to determine the effects of the boundaries on water levels in the area of interest. This is of particular concern in the model simulations involving increased ground-water withdrawals, because water levels in the external boundaries could change in response to stresses not accounted for in the calibrated model. Water levels in the general-head boundaries were varied by a factor of one-half. The resulting water-level change in the area of interest was between 1 and 7 ft. Changes in hydraulic conductance resulted in water-level changes of 1 to 5 ft.

TABLE 9 SENSITIVITY ANALYSIS OF MODEL INPUTS FOR THE CALIBRATED MODEL FOR THE PRINCESS ANNE AREA [ft = feet; < = less than]

	Change	Approximate range of water- level change in layer 2, in
Model input	in input	ft (negative value indicates water-level decline)
Vertical hydraulic	10x	5 to 12
conductivity of	. 1x	-3 to -5
confining unit		
Transmissivity	1,5x	1 to 12
	.75x	-1 to -12
Constant water level	1.5x	1 to 2
specified in layer 1	. 5x	-1 to -2
Water level in general-	1.5x	2 to 9
head boundaries	.5x	-2 to -9
Hydraulic conductance in	2x	1 to 4
general-head boundaries	.5x	-1 to -5
Southern boundary	no-flow to	
condition	constant head	< 1
	at 1 ft	

Finally, the boundary condition along the southern border of the model was changed from no-flow to a constant head of 1 ft. The resulting difference in water levels in the area of interest was less than 1 ft.

Steady-state simulation of increased ground-water withdrawals from the Manokin aquifer

Ground-water withdrawals from the Manokin aquifer in the Princess Anne area are expected to increase by about 600,000 gal/d by the year 2000. This includes increased pumpage from the Princess Anne municipal wells, as well as pumpage from new wells installed for the Eastern Correctional Institution. Simulations of steady-state conditions were conducted for this increased pumping rate in the Manokin aquifer (layer 2). The pumpage distribution used for the simulation is shown in figure 33. Figure 40 depicts the resulting model-generated water levels in the Manokin. In the area of interest, water-level altitudes range from 80 ft below sea level near pumped wells to 10 ft below sea level at the model boundary. This represents water-level declines of 15 to 70 ft from 1986. Associated with the water-level declines is an increased hydraulic gradient between Westover and the pumped wells near Princess Anne. This suggests that water containing elevated chloride concentrations in the vicinity of Westover could migrate more quickly toward the pumping wells. Because of possible inaccuracies in aquifer properties and boundary conditions, and the doubt concerning the assumption of steady-state conditions, these results should be regarded as approximations only. Sensitivity analysis performed on the model indicates that errors in model inputs may impact simulated water levels eonsiderably.

Table 10 lists the results of the sensitivity analysis for the model when increased pumpage from model layer 2 is included. The inputs that had the greatest effect on simulated water levels were vertical hydraulie conductivity of the overlying confining unit and transmissivity of the Manokin aquifer (layer 2). Variation of these inputs within reasonable limits results in water-level changes of up to 33 ft. For both inputs, water-level changes are greatest near the pumped wells and least near the boundaries of the area of interest. This indicates that errors in these inputs would affect the hydraulic gradient in the area of interest, as well as water levels.

The average linear velocity of ground water depends on the hydraulic conductivity of the sediments, their porosity, and the hydraulic gradient. They are related by the equation:

$$v = KI/n_{e}, (5)$$

where: v = average linear velocity, in feet per day;

K = hydraulic conductivity, in feet per day;

I = hydraulic gradient; and

 $n_{r} = effective porosity.$

The amount of time required for ground water to travel a specified distance may be determined by dividing the distance by the average linear velocity. The times required for high-chloride water in the vicinity of Westover to reach the nearest simulated pumped well at the Eastern Correctional Institution (well SO Ce 48) were estimated for the simulation of inereased pumpage. An assumed effective porosity of 30 percent, a hydraulic conductivity of 13.2 ft/d, and hydraulic gradients derived from simulated water levels were used in the cal-



Figure 40.— Simulated water-level altitudes in the Manokin aquifer using projected pumpage.

TABLE 10
SENSITIVITY ANALYSIS OF MODEL INPUTS FOR THE SIMULATION OF INCREASED PUMPAGE IN THE PRINCESS ANNE AREA

[ft = feet]

Model input	Change	Approximate range of water- level change in layer 2, in ft (negative value indicate				
	in input	water-level decline)				
Vertical hydraulic	10x	15 to 33				
conductivity of	.1x	-6 to -16				
confining unit						
Transmissivity	1.5x	2 to 27				
	.75x	-1 to -26				
Water level in layer 1	1.5x	1 to 2				
	. 5x	-1 to -2				
Water level in general-	1.5x	2 to 9				
head boundaries	. 5x	-2 to -9				
Hydraulic conductance in	2x	2 to 7				
general-head boundaries	. 5x	-3 to -10				
Southern boundary	no-flow to					
condition	constant head	1 to 2				
	at 1 ft					

eulations. The results are listed in table 11. Using the linear hydraulie gradient between an isoehlor and the nearest simulated pumped well, it would take about 50 years for water to move from the vieinity of the 250-mg/L isoehlor to the well and about 300 years for water to move from the 500-mg/L isoehlor to the well.

A more immediate concern is that as the high-ehloride water moves northward, more of the aquifer becomes unsuitable as a source of potable water. Because the hydraulie gradient is steepest in the vicinity of a pumped well, ground-water velocity increases as water moves toward the nearest simulated pumped well. For example, the gradient near the simulated pumped well is about 0.013. Assuming an effective porosity of 0.30, the ground-water velocity is about 210 ft/yr, which is about 2-3 times greater than the average linear velocity.

There are several limitations to the above analysis. In the analysis, eonstant hydraulic conductivity was used. If there are zones of higher hydraulic conductivity in the aquifer, high-ehloride water can flow more quickly through these zones and arrive at a pumped well more quickly than calculated. The other limitation is in the value used for effective porosity. This value was not measured directly, but represents a value typical of uniform sands (Bear, 1972, p. 46). Table 11 lists the average linear velocities and travel times using alternative ef-

TABLE II
ESTIMATED TRAVEL TIMES AND AVERAGE LINEAR VELOCITIES OF GROUND WATER UNDER CONDITIONS OF SIMULATED INCREASED PUMPAGE IN THE PRINCESS ANNE AREA [ft/yr = feet per year; mg/L = milligrams per liter]

Effective porosity	Gradient	Average linear velocity (ft/yr)	
	nity of 250-mg/L isochlor	to nearest simulated pu	umped
Mell (about 3,900 feet)		
0.30	0.0051	80	49
.20	.0051	120	32
. 40	.0051	61	64
From vicin	nity of 500-mg/L isochlor	r to nearest simulated po	umped
well (about 13,000 feet)		
0.30	0.0027	43	300
.20	.0027	66	190
. 40	.0027	32	390

fective porosities of 20 and 40 percent. The result of decreasing effective porosity by one-third is to decrease the travel time by one third. Conversely, increasing the effective porosity by one-third increases the travel time by one-third.

In addition to water moving more rapidly from Westover to Princess Anne, water also will migrate more rapidly eastward toward Princess Anne in response to the increased ground-water withdrawals. The Manokin aquifer subcrops beneath the Chesapeake Bay; therefore, brackish water from the bay may migrate toward pumping centers more quickly as the pumpage increases. The simulated hydraulic gradient in the Manokin west of Princess Anne is about 5 ft/mi. Using this gradient and the values from the previous analysis, ground-water velocity in layer 2 west of the pumping centers is about 15 ft/yr.

The water budget generated by the model shows that under the increased-pumpage conditions, 72 percent of the water entering the Manokin aquifer is from general-head boundaries, and 28 percent is from leakage through the confining unit. Almost all of the water, 99 percent, leaves the modeled area by ground-water pumpage. The remaining 1 percent leaves by leakage to the upper layer.

Estimated Effects of Projected Ground-Water Pumpage in the Crisfield Area

The town of Crisfield is the largest user of ground water in Somerset County. Present withdrawals are about 800,000 gal/d and are expected to increase to about 1,100,000 gal/d by the year 2000 (J. Windsor, Somerset County Department of Technical and Community

Services, written eommun., 1986). Water supply is derived from the Paleoeene and Potomae aquifer systems. Few multiple-well aquifer tests have been performed in the Crisfield area because of the cost of drilling deep observation wells, and little is known of the characteristics of the two aquifer systems outside the area. Due to the lack of available data, a detailed analysis using a digital-flow model of the possible effects of projected ground-water pumpage is not feasible. Therefore, an analytical solution was used to give an approximation of the additional water-level decline that may result from increased ground-water pumpage.

The Theis method is an analytical solution to the partial-differential equation that describes unsteady, radial flow in a confined aquifer (Theis, 1935; Freeze and Cherry, 1979, p. 317). The solution, in terms of drawdown, is:

$$h_o - h(r,t) = \frac{Q}{4\pi T} \int_u^{\infty} \frac{e^{-u} du}{u}, \qquad (6)$$

where: $u = r^2S/4Tt$;

r = radial distance from the pumped well, in feet;

t = time since pumping started, in days;

Q = discharge rate of the pumped well, in cubic feet per day;

 h_0 = head at t=0, in feet;

h(r,t) = head at distance r (feet), and time t (days);

T = transmissivity, in feet squared per day; and

S = storage eoefficient (dimensionless).

This investigation utilized a computer program that uses the Theis method to calculate water-level declines at specific distances from pumped wells at specified times (Walton, 1985). The program requires information on aquifer transmissivity, pumped-well locations, well discharges, storage coefficients, and time. Transmissivity was estimated from one multiple-well aquifer test and two single-well aquifer tests in the area. Transmissivity ranges from 1,050 to 2,140 ft²/d, with an average of 1,490 ft²/d. Storage coefficient from the multiple-well aquifer test is 0.0002. The high, low, and average transmissivity values were used in the analysis, resulting in a range of possible water-level declines.

A number of simplifying assumptions were necessary to use the method. These include:

- (1) The aquifer is homogeneous and isotropie;
- (2) ground-water flow is horizontal and radial toward pumped wells;
- (3) vertical leakage to or from the aquifer is negligible;
- (4) the aquifer is infinite in areal extent;
- (5) ground-water discharge is by well withdrawals;
- (6) water removed from storage is discharged instantaneously with declines in water levels; and
- (7) the ground-water-flow system is initially under steady-state conditions.

Several of these assumptions are not met in the aetual ground-water-flow system and may affect the accuracy of the results:

- (1) The flow system at Crisfield actually eonsists of several aquifers and eonfining units. In the method, this eomplex system is simplified to one aquifer;
- (2) leakage to or from the aquifer systems is unknown. If there is significant leakage, the water-level declines obtained from the computer program will be greater than actual declines; and

(3) the aquifer is not of infinite areal extent. The two aquifer systems exhibit complex geology in the area of interest. The aquifers in the systems may pinch out or become finer-grained a short distance from the wells, thereby significantly reducing the aquifer transmissivity. If this is the case, calculated water-level declines will be considerably less than actual declines.

Because the simplifying assumptions do not fully describe the aquifer systems, the values generated by the Theis method should be regarded as rough approximations of actual water-level declines.

Calculated water-level declines resulting from additional ground-water discharge are listed in table 12. These values represent additional declines in response to an additional ground-water discharge of 300,000 gal/d distributed between five pumped wells for a period of 20 years at the rates shown in figure 33. For example, if the calculated water-level decline is 20 ft and the initial water level is 18 ft below land surface, the resulting water level would be 38 ft below land surface. For a given range of water-level declines in table 12 the smaller value generally occurs near the edge of the area of interest, and the larger value occurs near pumped wells (fig. 33). In table 13, the largest declines, 13 to 31 ft, are associated with the smallest transmissivity. Likewise, the smallest declines, 7 to 16 ft, are associated with the largest transmissivity. Considering that water levels could decline about 900 ft before the top of the Paleocene aquifer system is reached, the estimates indicate that little of the available drawdown will be needed to meet the projected additional demand for water. However, because the characteristics of the aquifer systems are not known east of Crisfield, changes in the quality of the water that may result from the additional pumpage are not known.

To more accurately determine the effects of additional pumpage in the Crisfield area, more data are needed on aquifer characteristics, particularly in the area east of Crisfield. Specifically, more information is needed on:

- (1) The degree of hydraulic connection between aquifers and the lateral extent of aquifers;
- (2) hydraulic properties of aquifers and confining beds;
- (3) the quality of water in the aquifers east of Crisfield;
- (4) water levels in the aquifer; and
- (5) rates and locations of ground-water withdrawal with time.

TABLE 12 WATER-LEVEL DECLINES RESULTING FROM INCREASED GROUND-WATER WITHDRAWALS IN THE CRISFIELD AREA, BASED ON AN ANALYTICAL SOLUTION $[ft^2/d\ =\ feet\ squared\ per\ day;\ ft\ =\ feet]$

ransmissivity (ft ² /d) 1.050	Approximate range of water-level declines (ft)					
1,050	13 to 31					
1,490	10 to 22					
2,140	7 to 16					

SUMMARY AND CONCLUSIONS

Somerset County depends on ground water for approximately 84 percent of its water supply. An assessment of the ground-water resources was conducted because development in the county is expected to substantially increase the demand for ground water. In particular, the goals of the investigation were to:

(1) Refine the hydrogeologic framework;

(2) describe the quality of ground water; and

(3) evaluate the effects of projected ground-water withdrawals at Princess Anne and

Crisfield on the ground-water-flow system.

Somerset County is underlain by a wedge of interbedded sands, silts, and clays that dip to the southeast. This wedge of generally unconsolidated sediments forms a series of aquifers and confining units. The aquifers receive water by the infiltration of precipitation, leakage from overlying and underlying aquifers, and by the lateral movement of water from outside the county. The aquifers and aquifer systems that ean supply water to wells in Somerset County are:

- (1) The surficial aquifer system;
- (2) the Pocomoke aquifer;
- (3) the Manokin aquifer;
- (4) the Choptank aquifer;
- (5) the Piney Point aquifer;
- (6) the Paleocene aquifer system; and
- (7) the Potomac aquifer system.

The uppermost water-producing unit is the surficial aquifer system. It is stratigraphically complex and exhibits rapid lithologic changes laterally and vertically. The aquifer system generally is coarser-grained and thicker in the northeastern part of the county and finer-grained and thinner in the remainder of the county. The aquifer system is used primarily to supply domestic wells, since the fine-grained nature and thinness of the system usually preclude its use as a source of water for high-yielding wells. The aquifer system in the northeastern part of the county may yield substantial quantities of water to wells, because of its thickness and coarse-grained nature.

The Pocomoke aquifer is present only in the southeastern part of the county. In some areas, the aquifer directly underlies the surficial aquifer system and receives water directly from it. The Pocomoke aquifer is used for domestic, irrigation, and industrial water supply. Reported specific capacities of wells in the aquifer range from 1 to 50 (gal/min)/ft. The median is 10 (gal/min)/ft.

In subcrop areas, water levels in the Poeomoke aquifer are probably adjusted to nearby surface-water bodies. In confined portions of the aquifer, water levels are less influenced by

surface-water bodies.

The primary aquifer of use in Somerset County is the Manokin aquifer. It supplies water to domestic wells, the town of Princess Anne, the Eastern Correctional Institution, numerous poultry operations, and a plywood-manufacturing plant. The aquifer is not used extensively in the southern part of the county because ehloride concentrations exceed the SMCL of 250 mg/L. Reported specific eapacities range from 0.1 to 75 (gal/min)/ft; the median is 5.2 (gal/min)/ft. Transmissivity of the Manokin aquifer near Princess Anne, ealeulated from four multiple-well aquifer tests and a single-well aquifer test, ranges from 500 to 940 ft²/d. Hydraulic conductivity ranges from 10.9 to 14.7 ft/d, with an average of 13.2 ft/d. Storage coefficient ranges from 0.0002 to 0.001.

Water levels in the Manokin aquifer have declined as much as 45 ft since the 1950's. Water-level altitudes throughout much of the county are below sea level, with the lowest altitudes located near the municipal wells at Princess Anne. Associated with the water-level declines are changes in ground-water-flow directions. Prior to heavy pumping, ground water generally flowed from northeast to southwest. Currently, ground water flows from the borders of the county toward pumped wells at Princess Anne. This change of flow direction could cause poor-quality water from the Chesapeake Bay or from the southern part of the county to migrate toward the pumping centers at Princess Anne.

The Choptank aquifer is capable of supplying large quantities of water to wells, but is not used for water supply because of water-quality concerns. Reported chloride concentrations

exceed 900 mg/L, causing the water to be unacceptable for most uses.

The Piney Point aquifer supplies water to one well on Deal Island, and contributes water to a well at Crisfield that is screened in multiple aquifers. A single analysis indicates that although water in the aquifer has chloride concentrations within acceptable limits, it may contain dissolved solids in excess of 1,000 mg/L, which may make it undesirable as a source of potable water.

The Paleocene aquifer system is used only by the town of Crisfield as a source of public water supply. Two wells are screened entirely in the aquifer system, and the system contributes water to three other wells that are screened in multiple aquifers. The water-yielding capabilities of the Paleocene aquifer system are not well known. Reported specific capacities of two wells screened entirely in the aquifer were nearly 2 (gal/min)/ft. Transmissivity of a well screened in both the Paleocene and Potomac aquifer systems was estimated to be 1,050 ft²/d. East of Crisfield, the extent and quality of water in the aquifer system are not known.

The Potomac aquifer system is used for public water supply by Crisfield, Smith Island, Rumbley, Frenchtown, and Fairmount. Reported specific capacities for four wells range from 1 to 7 (gal/min)/ft. Transmissivity and storage coefficient, calculated from a multiwell aquifer test are 2,140 ft²/d and 0.0002, respectively. Transmissivity estimated from one 24-hour, single-well aquifer test is 1,280 ft²/d.

Water levels in the Potomac aquifer system in the Crisfield area are 8 to 18 ft below land surface. There is more than 950 ft of drawdown available; therefore, greater well yields are possible from the aquifer system. However, the extent and water quality of the producing aquifers are not known east of Crisfield. An electric log from a geothermal test well suggests that one or two additional aquifers may contain potable water.

The ground-water quality in Somerset County differs considerably, both laterally and vertically. Much of the county contains potable ground water, but there are areas where it is not possible to obtain water that meets drinking-water standards. South of Westover the Pocomoke aquifer contains iron concentrations in excess of the SMCL and the Manokin aquifer contains chloride concentrations in excess of the SMCL. Deeper aquifers probably are brackish east of Crisfield.

Chemical analyses of four water samples from the surficial aquifer system indicate the water is soft to moderately hard and slightly acidic. The water quality in the surficial aquifer system is likely to differ from site to site, because of the complex geology and the influence of land-use practices. In sites containing anoxic ground water or where unstable conditions exist, iron concentrations in excess of $300\,\mu\text{g}/\text{L}$ are likely. Nitrate contamination is possible in areas containing oxygenated ground water where the water can come into contact with nitrogen fertilizers or organic wastes.

Ground water from the subcrop areas of the Pocomoke aquifer is generally softer, con-

tains fewer dissolved solids, and is more acidic than water in the confined parts of the aquifer. The confined parts of the aquifer generally contain greater concentrations of calcium and bicarbonate.

The most common water-quality problems in samples from the Pocomoke aquifer are iron and manganese concentrations in excess of the SMCL's. All but 2 of 24 samples exceed the limit for iron, and 18 of 24 exceed the limit for manganese. One sample exceeds the SMCL's for chloride and dissolved solids.

There is a marked areal difference in the quality of water from the Manokin aquifer. North of Westover, water from the aquifer is lower in dissolved solids, softer, and is a sodium bicarbonate type. South of Westover, the water contains more dissolved solids, is harder, and is a sodium chloride type. In the southern part of the county, chloride concentrations in the Manokin aquifer generally exceed the SMCL of 250 mg/L. This may be due to the presence of a transition zone or areas of low permeability that result in incomplete flushing of the aquifer. Dissolved iron is a problem in the northern part of the county. Iron concentrations generally are higher in the northeastern corner and decrease to the south and east.

Few wells produce water from the Choptank or Piney Point aquifers; therefore, little is known about the quality of the water in them. Reported chloride concentrations in the Choptank aquifer exceed 900 mg/L, and one sample from the Pincy Point aquifer contained ehloride concentrations less than 250 mg/L but had dissolved solids in excess of 1,000 mg/L.

Water from the Paleocene and Potomac aquifer systems is soft, has dissolved-solids concentrations ranging from 475 to 1,070 mg/L and has the highest pH of any aquifer in the county. The water from the two systems is of the sodium-bicarbonate type, with sodium accounting for more than 95 percent of the cations. The primary water-quality problem in the Paleocene and Potomac aquifer systems is high concentrations of fluoride. Seven of 10 water-quality samples contain concentrations of fluoride above the SMCL and two of these exceed the MCL.

The effects of projected ground-water pumpage were evaluated for the Princess Anne and Crisfield areas. A digital ground-water-flow model was used under steady-state conditions to evaluate the effects of current and projected pumpage from the Manokin aquifer in the vicinity of Princess Anne. The Theis solution was used to give a rough approximation of the effects of increased pumpage from the Paleocene and Potomac aquifer systems at Crisfield, because data necessary to construct a digital, ground-water-flow model for the Crisfield area were not available.

The aquifers in the Princess Annc area were represented by two model layers. The surficial aquifer system and the Poeomoke aquifer comprised layer 1 and the Manokin aquifer was represented by layer 2. Heads in layer 1 were held constant during the simulation, so that ground-water flow was simulated in layer 2 only. Flow through the confining unit overlying the Manokin aquifer was calculated from leakance values supplied to the model. Laterally, the model boundaries were no-flow on the southern boundary and general-head on the other three sides.

The model was calibrated to water levels measured in the Manokin aquifer during November 1986. The water budget generated by the model indicates that 74 percent of the water entering layer 2 is from general-head boundaries and 26 percent is from leakage through the confining unit. Ground-water withdrawal accounts for 97 percent of the water leaving the model layer.

Simulation of increased ground-water withdrawals from the Manokin aquifer of 600,000

gal/d resulted in water-level declines of 15 to 70 ft from current water levels. Associated with the additional water-level declines are increased ground-water velocities. Using model-derived velocities, it would take high-ehloride water in the vicinity of the 250-mg/L isochlor 50 years to reach the pumped wells near Prineess Anne. The travel time for water to move from the 500-mg/L isochlor to the nearest simulated pumped well is 300 years. These values are only approximations because of uncertainties in aquifer properties and simplifying assumptions. This does not account for possible zones of increased hydraulic conductivity along which the travel time would be shorter.

The water budget generated by the model shows that under increased pumping conditions, 72 percent of the water entering model layer 2 is from general-head boundaries and 28 percent is from vertical leakage. Almost all of the water, 99 percent, leaves the area by ground-water pumping.

Ground-water pumpage in the Crisfield area is expected to increase by 300,000 gal/d. In the solution used to determine the effect of this additional pumpage, the Palcocenc and Potomac aquifer systems were simplified to one aquifer, and the aquifer was assumed to be infinite in areal extent. The storage coefficient of the aquifer was assumed to be 0.0002, a value obtained from a multiple-well aquifer test. Transmissivity values of 1,050, 1,490, and 2,140 ft²/d were used in the analysis. Additional water-level declines under these conditions ranged from 7 to 31 ft. The nature of the aquifer systems are unknown east of Crisfield; therefore, it is not known if high-chloride water would migrate toward the pumping centers at an increased rate because of additional pumpage.

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SUPPLEMENTAL DATA

TABLE 13 CHEMICAL ANALYSES OF GROUND WATER IN SOMERSET COUNTY—explanation of codes

Aquifer codes

112PCPC	Surficial aquifer system
122MOCN	Miocene series (undifferentiated)
122PCMK	Pocomoke aquifer
122MNKN	Manokin aquifer
122CPNK	Choptank aquifer
124PNPN	Piney Point aquifer
125PLCN	Paleocene aquifer system
217PTMC	Potomac aquifer system

TABLE 13
CHEMICAL ANALYSES OF GROUND WATER IN SOMERSET COUNTY—Continued [μ S/cm = microsiemens per centimeter; deg C = degrees Celsius; mg/L = milligrams per liter; < = less than; -- = no data]

Well	no.	Aquifer		Spe- cific con- ductance leboretory (µS/cm)	pH field (stand- erd unite)	Watar temper- ature (deg C)	85	Oxygen, die- solved (mg/L)	Celcium, die- solved (mg/L ee Ce)	Magne- eium, die- eolved (mg/L ee Mg
SO Ad SO Ae SO Ae SO Ae	13 16 17 18 20	122MNKN 122PCMK 112PCPC 122MNKN 122MNKN	09-10-86 08-05-86 09-10-86 09-10-86 08-05-86	172 174 255	6.8 4.6 5.4 6.8 6.6	17.0 15.5 16.0 17.0	88 53 59 59	0.2	5.5 15	11 5.2
SO Bb SO Bb SO Bc SO Bd	15 19 15 33	122MNKN 124PNPN 122MNKN 122MNKN 122MNKN	08-05-86 08-06-86 08-06-86 08-06-86	1,530 1,950 1,890	6.6 6.8 7.6 7.5 6.6		97 220 33 130 34	0 0 0	27 43 6.1 24 7.5	7.2 28 4.2 18 3.8
SO Bd	37	112PCPC	06-26-86		6.6	15.0	140		41	9.1
SO Bd SO Be SO Be SO Be	39 14 49 51	122MNKN 122MNKN 122MNKN 122PCMK 122MNKN	08-07-86 12-08-52 12-08-52 08-28-86		7.6 8.2 5.8 7.8	16.5 16.0	47 52 51 40	.5	9.6 13 12 8.2	5.6 4.7 5.1 4.7
SO Be SO Be SO Be SO Be SO Be	54 56 58 72 77	122MNKN 122MNKN 122MNKN 122PCMK 122PCMK	08-28-86 08-28-86 08-07-86 08-07-86 07-29-86	916 829	8.0 7.8 7.9 6.5 4.5	16.0 16.5 21.0 16.0 16.0	42 47 44 26 49	0 0 0	8.8 9.0 8.4 4.6 8.5	6.0
SO Be SO Be SO Be SO Be		122MNKN 122MNKN 122MNKN 122PCMK 122PCMK	08-05-86 08-05-86 08-12-86 07-29-86 08-06-86	542	7.7 7.4 7.8 5.6 4.9	17.5 18.0 17.0 16.5 16.5	45 32 60 11 56	2.0 0 1.0	10 4.7 12 2.4	4.8 4.9 7.2 1.2 6.8
SO Be SO Be SO Be SO Ba	89 91 92 93 94	122PCMK 122MNKN 122MNKN 122MNKN 122PCMK	08-06-86 08-07-86 09-05-86 09-05-86 09-08-86	737	5.1 7.7 6.6 7.7 6.0	16.5 18.0 18.0 17.5 17.0	16 37 36 2 20	0	2.6 6.9 8.2 .40 3.2	4.7
SO Be SO Bf SO Bf SO Bf SO Bf	95 14 15 17 18	122MNKN 122MNKN 122MNKN 122MNKN 122MNKN	09-08-86 08-07-86 08-07-86 08-07-86 08-08-86	630	7.4 7.9 8.2 7.8 7.6	16.5 18.0 16.5 16.5	46 39 43 78 51	.2	8.5 8.3 9.9 19	5.9 4.3 4.5 7.3 5.0
SO Bf SO Cb SO Cc SO Cc	20 24 6 7	112PCPC 122MNKN 217PTMC 217PTMC	08-07-86 09-04-86 09-04-86 04-03-70	285 1,590 1,160 1,240	5.5 7.1 8.5		79 170 6 6		16 29 1.4 1.4	9.5 24 .60 .70
SO Cd	41	217 PTMC 122 PCMK	05-29-79 05-29-79 05-30-79 09-04-86 08-14-86	1,000	8.3 8.3 8.4 8.7 5.6	27.0 16.5	6 6 6 6	 0 0	1.6 1.7 1.6 1.5	.60 .50 .60
SO Cd SO Cd SO Cd SO Cd SO Cd	45 49 50 51 52	122PCMK 122MNKN 122MNKN 122MNKN 122PCMK	08-14-86 07-29-86 09-03-86 07-31-86 08-14-86	158 2,060 1,570				0 . 7	8.8 25 63 35 8.6	2.6 17 24 25 5.9
SO Ce SO Ce SO Ce	64	122MNKN 122MNKN 122MOCN 122MNKN	08-12-86 08-08-86 08-12-86 08-08-86		7.3 7.8 7.7	20.5 16.5 22.0 17.0	83 230 120		21 41 21 24	7.3 32 17 16

Sodium, dis- solvsd (mg/L ss Ns)	Potss- sium, dis- solved (mg/L aa K)	Alka- linity, csrbonste fisld (mg/L as CsCO ₃)	Sulfate, dis- solved (mg/L as SO ₄)	Chlo- ride, dis- solved (mg/L ss Cl)	Fluoride, dia- solved (mg/L ss F)	Silica, dis- solved (mg/L as SiO ₂	st 180 deg C dis- solved	Nitrogen, ammonis dis- solved (mg/L ss N)	Well	no
43 6.5 5.1 31 25	5.1 2.3 2.0 4.2 4.3	190 5.5 140 158	1.7 25 27 4.4 2.3	15 14 14 7.1 5.8	0.30 < .10 < .10 .30 .20	20 13 12 26 22	220 100 108 173 184	0.220 .020 < .010 .150 .150	SO Ad SO Ae SO Ae SO Ae	13 16 17 18 20
210 400 340 62	19 13 17 6.4	320 772 450 356	12 68 58 3.4	320 180 330 11	.20 2.3 .30 .20	38 23 11 24	846 1,180 1,070 226	2.40 .600 .960 .230	SO Bb SO Bc SO Bd	15 19 15 33
14	2.3	84	79	21	.30	17	236	.150	SO Bd	37
180 150 24 150	8.6 10 4.2 6.8	380 304 22 305	11 18 52 10	74 62 24 47	.30 .40 .20 .40	11 14 26 11	537 460 161 593	.340	SO Be SO Be SO Be	39 14 49 51
150 190 170 56 14	6.8 8.1 7.9 5.7 2.6	305 345 319 152 0	14 10 7.9 7.6	54 93 72 7.7	.40 .40 .40 .40 < .10	11 10 10 39 29	404 524 491 220 160	.280 .380 .330 .180 .240	SO Be SO Be SO Be SO Be	54 56 58 72 77
110 110 210 13 12	5.9 5.6 9.5 1.5 2.1	272 264 410 26 4.0	3.2 3.0 19 14 55	20 20 110 9.3	.40 .30 .40 < .10	12 150 11 27 24	332 329 599 90 150	.250 .220 .370 .060	SO Be SO Be SO Be SO Be	83 84 86 87 88
6.4 150 90 230 20	25 7.5 4.4 1.5 2.2	9.0 405 240 355 61	45 4.1 3.7 19 8.5	10 51 12 120 10	< .10 .40 .50 .40	14 68 16 9.8 38	108 435 276 620 131	.040 .290 .200 .120 .220	SO Be SO Bs SO Be SO Be	93 93 93
180 129 76 53 74	7.5 5.6 5.9 6.1 5.0	350 301 230 200 220	10 5.3 2.1 1.0 1.6	78 28 12 7.0 6.7	.50 .30 .30 .20	11 10 12 14 14	544 378 262 230 250	.370 .220 .180 .200 .190	SO Be SO Bf SO Bf SO Bf	9:
13 270 270 300	8.8 19 6.7 7.0	13 375 660 635	40 6.1 23 24	22 300 6.7 6.2	< .10 .10 4.8 5.2	12 24 12 12	178 847 710 766	.260 2.30 .130	SO Bf SO Cb SO Cc SO Cc	24
260 240 220 230 16	5.8 7.4 6.1 5.5 2.3	510 510 520 530 60	27 27 30 24 59	7.7 8.1 8.0 7.4	2.6 2.6 2.6 2.5	12 11 12 12 76	613 611 629 580 242	 .420 .900	SO Cd	4:
15 370 210 460 49	3.0 16 15 19 2.2	51 395 210 405 60	16 110 47 110 63	15 370 350 490 72	.20 .20 .20 .20	21 10 17 10 33	123 1,190 807 1,430 287	.280 .820 .350 1.00	SO Cd SO Cd SO Cd SO Cd	4: 5: 5: 5:
140 590 240 350	7.6 24 18 16	295 430 400 300	21 140 16 98	63 770 210 340	.30 .20 .10 .20	21 9.9 14 9.8	464 1,860 770 1,100	.390 1.40 .050 .890	SO Cs SO Ce SO Ce	53 56 64

TABLE 13
CHEMICAL ANALYSES OF GROUND WATER IN SOMERSET COUNTY—Continued [deg C = degrees Celsius; mg/L = milligrams per liter; μ g/L = micrograms per liter; < = less than; -- = no data]

Well	no.	Nitrogen, ammonis plus organic dissolved (mg/L ss N)	Nitrogen, NO ₂ + NO ₃ dissolved ³ (mg/L as N	Phos- phorus, dis- solved (mg/L ss P)	Bsrium, dis- solved (µg/L ss Bs)	Beryl- lium, dis- solved (µg/L ss Be)	Boron, dis- solved (µg/L ss B	Cadmium, dis- solved (µg/L ss Cd)	Cobalt, dis- solved (µg/L ss Co)	Copper, dis- solved (µg/L ss Cu)
SO Ad SO Ae SO Ae SO Ae SO Af	13 16 17 18 20	0.40 .40 .70 .40	<0.100 6.30 5.70 < .100 < .100	0.270 .040 .020 .270 .370	23 92 48 20 18	<0.5 < .5 < .5 < .5 < .5	150 20 20 100 110	<1 1 <1 <1 <1	<3 4 <3 <3 <3	<10 10 <10 <10 <10
SO Bb SO Bb SO Bc SO Bd	15 19 15 33	2.3 .70 1.1 .30	< .100 < .100 < .100 < .100	.410 .040 .180 .690	36 10 32 15	< .5 < .5 < .5 < .5	780 3,000 1,200 190	2 <1 4 <1	<3 <3 <3	<10 <10 <10 <10
so Bd	37	.30	< .100	< .010	37	< .5	20	1	<3	<10
SO Bd SO Be SO Be SO Be	39 14 49 51	.60	< .100	.350	18 12	< .5 < .5	560 410	<1 2	<3 <3	<10 40 <10
SO Be SO Be SO Be SO Bs SO Be	54 56 58 72 77	.40 .50 .60 .40	< .100 < .100 < .100 < .100 < .100	.460 .340 .400 1.60	14 13 14 20 160	2 .6 < .5 < .5	450 600 520 120 20	3 2 <1 <1 <1	<3 <3 <3 <3 <3	<10 <10 <10 <10 <10
SO Be SO Be SO Be SO Be SO Be	83 84 86 87 88	.40 .50 .50 < .20 < .20	< .100 < .100 < .100 < .100 1.20	.610 .620 .360 .020	14 16 15 16 75	< .5 < .5 < .5 < .5	260 250 670 20 10	<1 <1 <1 <1 <1	<3 <3 <3 <3 70	10 <10 <10 <10 <10
SO Be SO Be SO Be SO Be SO Be	89 91 92 93 94	.30 .60 < .20 2.2 .30	.150 < .100 < .100 < .100 < .100	< .010 .420 .820 .370 .720	38 12 17 11 27	< .5 < .5 < .5 < .5 < .5	30 480 220 670 40	<1 <1 <1 1 3	<3 <3 <3 <3 <3	<10 <10 <10 <10 <10
SO Be SO Bf SO Bf SO Bf SO Bf	95 14 15 17 18	.60 .50 .30 .40	< .100 < .100 < .100 < .100 < .100	.370 .380 .410 .300	15 14 14 18 16	< .5 < .5 < .5 < .5 < .5	540 350 240 170 200	1 <1 <1 <1 <1	<3 <3 <3 <3 <3	<10 <10 <10 <10 <10
SO Bf SO Cb SO Cc SO Cc	20 24 6 7	.70 2.8 .30	< .100 < .100 < .100	.010 2.10 .240	110 16 36	< .5 < .5 .5	40 840 2,000	<1 <1 <1	<3 <3 <3	10 <10 <10
SO Cd	41	.70	< .100	.290	<100 28	.6	1,100 1,200	 <1	 2 <3	<10
SO Cd SO Cd SO Cd SO Cd SO Cd SO Cd	45 49 50 51 52	1.0 .40 1.0 .60 1.2 < .20	< .100 < .100 < .100 < .100 < .100 < .100 < .100	.130 .010 .140 .010 .120 .020	27 26 27 54 40 48	< .5 < .5 < .5 < .5 <2 < .5	20 1,200 430 1,300 30	2 <1 5 7 7 3	<3 <3 <3 <3 <9 <3	<10 <10 <10 <10 <10 <10 <30 <10
SO Ce SO Ce SO Ce SO Cs	53 56 64 65	.90 1.5 .30	< .100 < .100 .680 < .100	.520 .130 .110 .150	15 43 22 24	< .5 5 < .5 < .5	420 1,600 980 1,100	<1 4 <1 2	<3 <9 <3 <3	<10 <30 40 <10

Iron, dis- solved (µg/L aa Fa)	Lesd, dis- aclved (µg/L aa Pb)	Lithium, dis- solved (µg/L aa Li)	Manga- nese, dis- solved (µg/L as Mn)	Molyb- denum, dis- solved (µg/L aa Mo)	Stron- tium, dis- solved (µg/L as Sr)	Vana- dium, dis- aolved (µg/L ss V)	Zinc, dis- solved (µg/L as Zn)	Wsll	no.
1,400 50 210 3,400 2,500	<10 10 <10 <10 10	15 6 5 15 12	67 34 21 130	<10 <10 <10 <10 <10	210 120 110 140 220	<6 <6 <6 <6	52 42 48 5	SO Ad SO Aa SO Aa SO Ae SO Af	13 16 17 18 20
4,100 370 140 1,700	<10 20 10 <10	18 35 22 12	290 3 42 46	<10 <10 <10 <10	380 180 320 71	<6 <6 <6	50 7 52 160	SO Bb SO Bb SO Bc SO Bd	15 19 15 33
13,000	<10	9	250	<10	250	<6	8	SO Bd	37
130	<10 <10	2,400 600 11	5 10 2,000 9	<10 <10	110 110	<6 <6	100 280 10 8	SO Bd SO Be SO Be	39 14 49 51
57 50 41 2,200 13,000	<10 <10 <10 <10 20	15 11 8 8	12 12 9 32 140	<10 <10 <10 <10 <10	120 120 110 48 110	<6 <6 <6 <6	12 10 100 3 36	SO Be SO Be SO Be SO Be	54 56 58 72 77
71 110 71 7,500 1,300	10 10 <10 <10 <10	8 13 12 6 6	17 17 11 100 30	<10 <10 10 <10 <10	100 100 140 27 180	<6 <6 <6 <6	64 <3 40 17 340	SO Be SO Be SO Be SO Be	83 84 86 87 88
2,200 30 460 13 4,000	20 <10 <10 <10 <10	<4 11 12 13 13	35 8 36 3	<10 <10 <10 <10 <10	41 92 83 5 37	<6 <6 <6 <6	57 <3 17 <3 3,000	SO Be SO Be SO Be SO Be	89 91 92 93
46 19 87 160 89	<10 10 <10 10 <10	13 7 9 7	9 6 9 27 17	<10 <10 <10 <10 <10	120 92 90 170 110	<6 <6 <6 <6	25 17 25 <3 6	SO Ba SO Bf SO Bf SO Bf	95 14 15 17
9 56 21 870	10 <10 <10 40	<4 23 12	11 34 8 	<10 <10 10	190 350 42	<6 <6 <6	5 7 <3	SO Bf SO Cb SO Cc SO Cc	24
360 80 80 10 17,000	20 <10 <10 <10 <10	 <10 11 52	20 <10 20 8 210	 <10 <10	 36 110	 <6 <6	 <3 130	SO Cd	41
11,000 110 3,200 470 41,000	<10 <10 <10 <30 <10	5 24 22 29 19	540 17 310 22 900	<10 <10 <10 <30 <10	55 340 500 520 61	<6 <6 <6 <18 <6	11 26 110 100 79	SO Cd SO Cd SO Cd SO Cd	4 5 4 9 5 0 5 1 5 2
1,300 200 10 140	<10 30 <10 10	13 29 10 14	120 15 11 10	<10 <30 <10 <10	250 690 260 360	<6 <18 <6 <6	610 57 25 32	SO Ce SO Ce SO Ca	53 56 64 65

TABLE 13 CHEMICAL ANALYSES OF GROUND WATER IN SOMERSET COUNTY—Continued [μ S/cm = microsiemens per centimeter; deg C = degrees Celsius; mg/L = milligrams per liter; < = less than; -- = no data]

Wall	no.	Aquifar		Spe- cific con- ductance laboratory (µS/cm)	pH field (atand- ard unita)	Water tempar- ature (deg C)	88	Oxygen, dis- aclvad (mg/L)	Calcium, dia- solved (mg/L aa Ca)	Magne- sium, dia- aolved (mg/L aa Mg
SO Ce SO Ce SO Cf SO Cf	74 81 83 16 20	122MNKN 122PCMK 112PCPC 122MNKN 112PCPC	09-03-86 08-12-86 08-12-86	754 387 1,850	7.8 6.5 5.8 7.5 5.0	16.0 15.0 21.5 16.5 16.5	110 180 45 120 50	0.4	19 53 10 22 11	14 11 4.8 16 5.5
SO Dc	3	217PTMC	$\begin{array}{c} \frac{1}{2}/10-21-66\\ \frac{1}{2}/10-26-66\\ \frac{1}{2}/11-02-66\\ 12-21-70 \end{array}$	1,000 1,220 971 982	8.1 8.6 8.4 8.5	23.5 27.0 28.0 27.5	9 19 4 4		2.6 6.9 1.0 .80	.60 .40 .30
SO Dc SO Dc SO Dd	5 6 29	217 PTMC 122MOCN 112PCPC 122MOCN	08-13-86	988 1,090 8,260	8.5 8.5 7.3 6.5 7.9	27.5 28.0 19.5 16.0	280 2,500 260	0 0	.80 .62 41 800 40	.50 .3: 42 110 39
SO Dd SO Dd SO Dd SO Dd SO Dd	48 51 58 59 60	122MOCN 122MOCN 122PCMK 122MOCN 122MOCN		657 1,820 469 1,100 1,250	7.2 7.3 6.0 7.3 7.3	16.0 15.5 16.5 17.0 17.5	180 360 160 160 150	.2	31 59 47 26 24	25 51 9.4 24 22
SO De SO De SO De SO De SO Df	29 31 33 36 9	122PCMK 122PCMK 122PCMK 122PCMK 122MNKN		384 623 95 463 2,220	6.7 6.4 6.3 6.6 7.6	15.5 16.5 15.5 16.5 18.0	140 180 22 180 140	0 0 0 0	31 60 6.2 55 23	14 7.9 1.6 9.5
SO Df SO Df SO Df SO Df SO Df	13 14 16 20 21	122MNKN 122PCMK 122PCMK 122PCMK 122PCMK	08-15-86 08-15-86 08-15-86 08-13-86 08-13-86	2,110 447 152 743 762	7.6 5.1 5.7 7.2 7.2	16.0 18.0 17.5 16.0 15.5	120 140 40 190 200	0 .6 0 0	21 22 9.1 40 38	17 20 4.1 22 26
SO Df SO Df SO Dg SO D8 SO Ea	26 27 5 7	122PCMK 122PCMK 122MNKN 122PCMK 217PTMC	08-15-86 09-04-86 08-12-86 08-12-86 04-26-48	180 263 2,060 156	4.7 6.4 7.4 6.0 8.4	16.0 15.0 18.0 18.0	24 100 120 48 8	0 0 0	2.8 37 22 12 2.0	4.1 2.4 16 4.4
SO Ea SO Ea	3 4 5	217PTMC 217PTMC 217PTMC	05-06-87 04-26-48 05-06-87 04-12-48	1,480 826	8.6 8.4 8.5	20.5 24.0 24.0	39 9 5 7		7.3 1.5 1.8	5.1 .40
SO Ea SO Ea SO Ec	7 11 1	217 PTMC 217 PTMC 125 PLCN	05-06-87 02-06-70 10-19-51 10-19-66 09-09-86	962 767 1,720 1,740 1,750	8.1 8.6 8.5 7.9 8.5	24.0 22.8 26.5	6 3 8 11 10		1.5 .80 1.8 2.4 1.9	.60 .30 .90 1.1 1.2
SO Ec	3	217PTMC 124PNPN 125PLCN 217PTMC	10-19-51 09-09-86 10-19-51	1,160 1,120 1,170	8.5 8.4 8.5	26.0 26.5 27.0	6 3 5	==	.50 .71 .50	1.2
SO Ec SO Ec SO Ec	48	122CPNK 217PTMC 125PLCN 217PTMC	12-08-52 09-09-86 09-09-86	5,780 1,130 1,240	7.6 8.6 8.7	27.5 28.5	210 3 4	.3	31 .62 .80	
SO Ed SO Ed SO Ed SO Ef	42 43 45	122MNKN 122MNKN 122MNKN 122PCMK	08-27-86 08-27-86 08-27-86 08-29-86	2,510 1,570 1,490 1,800	7.0 7.1 7.4 7.3	18.0 20.0 19.0 17.0	210 250 76 410	.6	34 45 14 91	29 33 10 45

Sodium, dis- aolved (mg/L as Na)	dis- aolved	linity, carbonate field	solved	dis- solved (mg/L	ride, dis- solved (mg/L	Silica, dis- solvad (mg/L as SiO ₂	at 180 deg C dia- solved	Nitrogen, ammonia dis- aolvad (mg/L aa N)	Well no.
350 80 19 370 34	15 7.1 1.7 15 1.6	415 200 56 420 12	86 63 72 86 82	320 110 31 310 54	0.20 .30 <.10 .20 <.10	9.3 51 24 11 42	1,070 444 224 1,060 252	1.00 1.00 .160 .910	SO Ce 74 SO Ce 81 SO Ca 83 SO Cf 16 SO Cf 20
250 240 250 240	6.7 65 6.4 5.0	439 427 447 437	44 56 43 43	55 110 42 41	1.8 1.6 1.2 2.1	12 14 13 12	649 749 620 624		SO Dc 3
240 230 120 870 273	5.0 4.7 22 13 26	437 440 340 220 385	43 47 84 300 162	41 41 120 2,800 278	2.1 2.0 .70 .20	12 12 12 59 15	624 610 590 5,720 1,080	.120 1.00 3.00	SO Dc 4 SO Dc 5 SO Dc 6 SO Dd 28
65 240 31 180 210	17 27 5.5 20 18	295 395 165 350	2.6 200 49 12 26	57 270 40 180 210	< .10 .10 .20 .10	15 12 61 13 13	381 1,100 358 642 745	1.10 1.80 1.10 1.10	SO Dd 48 SO Dd 51 SO Dd 58 SO Dd 59 SO Dd 60
26 55 8.2 27 410	8.9 5.3 1.2 5.3	186 185 36 162 395	2.9 35 12 26 89	17 80 11 36 450	.10 .20 .10 .20	41 56 53 41 10	253 422 116 309 1,260	.510 1.20 .500 .330 1.10	SO De 29 SO De 31 SO De 33 SO De 36 SO Df 9
400 13 13 85 89	17 25 2.5 10 16	380 15 68 240 310	91 120 18 19	420 35 15 95 71	.20 < .10 .20 .20	9.6 13 37 34 32	1,200 281 136 456 464	1.10 .040 .370 .490 .580	SO Df 13 SO Df 14 SO Df 16 SO Df 20 SO Df 21
9.9 12 380 12 241	2.0 1.2 16 2.6	5.0 133 385 76 445	30 2.4 79 22 32	17 13 410 9.4	< .10 .30 .20 .40	19 29 12 36 11	97 177 1,170 132 588	.030 .210 .930 .230	SO Df 26 SO Df 27 SO Dg 5 SO Dg 7 SO Ea 1
300	8.4 4.1	400	47 35 28 30	170 8.0 6.0	2.2 4.4 1.8 2.8	13 13 11	809 475 522	.040	SO Ea 3 SO Ea 4
220 180 440 420 400	5.5 4.0 8.0 8.0 8.9	352 763 711 710	33 39 65 58 62	12 11 100 120 130	2.8 2.0 5.6 4.6 4.4	12 13 13 11 11	551 473 1,090 1,080 1,070	.040	SO Ea 11 SO Ec 1
290 260 290	3.0 5.3 4.4	490 470 487	51 53 56	70 66 72	2.2 2.2 1.8	14 12 14	732 680 730	.500	SO Ec 3
,300 260 280	45 5.0 5.2	984 450 465	62 51 56	1,400 72 110	.70 2.1 2.0	58 12 12	3,440 708 733	.090 .170	SO Ec 33 SO Ec 48 SO Ec 49
440 200 280 190	25 24 16 22	475 380 440 305	53 55 28 48	530 240 220 380	.20 .20 .30 .20	29 19 23 33	1,840 1,180 853 1,440	1.50 1.30 .990 1.00	SO Ed 42 SO Ed 43 SO Ed 45 SO Ef 6

TABLE 13
CHEMICAL ANALYSES OF GROUND WATER IN SOMERSET COUNTY—Continued [deg C = degrees Celsius; mg/L = milligrams per liter; μ g/L = micrograms per liter; < = less than; -- = no data]

Well	no.	Nitrogan, ammonia plus organic dissolved (mg/L as N)	Nitrogen, NO ₂ + NO ₃ dissolved (mg/L as N	Phos- phorus, dis- solved (mg/L as P)	Barium, dis- solved (µg/L as Ba)	Beryl- lium, dis- solved (µg/L as Be)	Boron, dis- solved (µg/L as B	Cadmium, dis- solved (µg/L as Cd)	Cobalt, dis- solvad (µg/L as Co)	Copper dis- solved (µg/L as Cu)
SO Ce SO Ce SO Cf SO Cf	74 81 83 16 20	1.2 1.5 .30 1.1	<0.100 < .100 < .100 < .100 < .100	0.170 .180 < .010 .220 .010	22 37 43 18 56	<0.5 < .5 < .5 < .5 < .5	1,200 140 30 1,100 20	5 <1 4 4 3	<3 <3 <3 <3 <3	<10 <10 <10 <10 <10
SO Dc	3									
SO Dc	4	.30	 < .100	.400	30	< .5	1,300	 <1	<3	<10
SO Dc SO Dc SO Dd	5 6 29	1.0	< .100 < .100	< .100 .100	30 370	< .5	560 80	<10	<3 <3	<10 <10 30
SO Dd SO Dd SO Dd SO Dd SO Dd	48 51 58 59 60	1.4 1.8 1.2 1.3	< .100 < .100 < .100 < .100 < .100	.100 .090 .060 .160	21 38 27 24 30	< .5 < .5 < .5 < .5 < .5	500 1,000 70 800 940	1 <1 2 <1 2	<3 <3 <3 <3 <3	<10 <10 <10 <10 <10
SO De SO De SO De SO De SO Df	29 31 33 36 9	.80 1.4 .50 .60	< .100 < .100 < .100 .700 < .100	.550 .110 .010 .110 .130	27 38 13 58 21	< .5 < .5 < .5 < .5 < .5	100 100 20 70 1,300	<1 1 <1 <1 <1	<3 <3 <3 <3 <3	<10 <10 <10 <10 <10
SO Df SO Df SO Df SO Df	13 14 16 20 21	1.4 .50 .40 .60	< .100 6.50 < .100 < .100 < .100	.160 < .010 .160 .370 .310	20 39 20 19 21	< .5 < .5 < .5 < .5 < .5	1,400 10 20 190 260	<1 2 1 <1 <1	<3 30 <3 <3 <3	<10 <10 <10 <10 <10
SO Df SO Df SO Dg SO Dg	26 27 5 7	< .20 .20 1.1 .40	< .100 < .100 < .100 < .100	< .010 .190 .100 .420	54 33 19 16	< .5 .5 < .5 < .5	20 10 1,300 30	1 <1 6 <1	<3 <3 <3 <3	<10 <10 <10 <10
SO Ea SO Ea	3	.90	< .100 < .100	.230	12 20	< .5 < .5	760 670	 <1 <1	<3 <3	<10 <10
SO Ea SO Ea SO Ea	5 7 11 1	.60 .80	< .100 < .100	.190	20 65	< .5 < .5	930 3,200	 <1 <1	<3 <3	<10 <10
SO Ec	3									
SO Ec SO Ec	4 33 48	.60 .20	< .100 < .100	.350	31 29	< .5 < .5	1,400	1 <1	<3 <3	<10 <10
SO Ec SO Ed SO Ed SO Ed SO Ef	49 42 43 45 6	.30 1.7 1.9 1.5	< .100 < .100 < .100 < .100 < .100	.360 .070 .070 .130	66 14 21 16 31	< .5 <2 < .5 2 < .5	1,600 1,500 960 1,300 310	<1 <1 <1 6 <1	<3 <9 <3 3 <3	<10 <30 <10 <10 <10

 $[\]frac{1}{2}$ sampled interval 1,128 - 1,138 ft

^{2/} sampled interval 1,272 - 1,287 ft

Iron, dis- solvsd (µg/L as Fe)	Lead, dis- solvsd (µg/L as Pb)	Lithium, dis- solved (µg/L as Li)	Manga- ness, dis- solved (µg/L as Mn)	Molyb- denum, dis- solvsd (µg/L as Mo)	Stron- tium, dis- solved (µg/L as Sr)	Vans- dium, dis- solvsd (µg/L as V)	Zinc, dis- solvad (µg/L ss Zn)	Ws11	no.
110 9,700 48,000 140 27,000	<10 <10 <10 <10 <10	13 34 13 11 29	7 180 180 5	<10 <10 <10 <10 <10	330 330 90 370 96	<6 <6 8 <6 <6	100 <3 2,800 210 41	SO Ce SO Ce SO Cf SO Cf	74 81 83 16 20
420 150 190 360 360 20	10	 11	40 10 <5 10 10	 10		 <6	 <3	SO Dc	3
300 55,000 110	<10 70 	17 200 5,000	1,300 10	<10 <10	500 4,500	<6 8 	34 120	SO Dc SO Dc SO Dd	5 6 29
550 500 16,000 180 220	<10 <10 <10 <10 <10	15 29 38 19 21	11 18 300 10 13	<10 <10 <10 <10 <10	360 810 290 320 340	<6 <6 <6 <6	190 91 12 76 14	SO Dd SO Dd SO Dd SO Dd SO Dd	48 51 58 59 60
2,400 4,400 5,500 380 170	<10 <10 <10 <10 <10	12 38 12 17 13	49 120 55 83 17	<10 <10 <10 <10 <10	240 360 31 320 410	<6 <6 <6 <6	140 22 13 7	SO De SO De SO De SO De SO Df	29 31 33 36 9
130 73 13,000 1,000 690	<10 <10 <10 <10 <10	12 5 10 18 19	6 640 250 96 12	<10 <10 <10 <10 <10	400 350 69 350 360	<6 <6 <6 <6	58 120 12 10 17	SO Df SO Df SO Df SO Df SO Df	13 14 16 20 21
7,500 13,000 150 9,500	<10 <10 <10 <10	7 15 11 10	100 270 5 250	<10 <10 <10 <10	47 250 380 90	<6 <6 <6	54 14 210 120	SO Df SO Df SO Dg SO Dg	26 27 5 7
20 14 8	20 10	9	31	<10 <10	180	<6 <6	23 19	SO Ea SO Ea	3 4
20 14 1,800 80 24	<10 <10	6 10 21	 8 4	<10 20	43	<6 <6	<3	SO Ea SO Ea SO Ea	5 7 11 1
120 26 100 170 15	<10 <10	12 14	 7 6	<10 10	29 25	<6 <6	5 6	SO Ec SO Ec SO Ec	3 4 33 48
22 90 210 78 980	<10 <30 <10 <10 <10	12 45 32 31 44	9 <3 7 8 52	10 <30 <10 <10 <10	36 600 540 240 920	<6 <18 <6 <6 <6	<3 22 13 14 29	SO Ec SO EC SO Ed SO Ed SO Ef	49 42 43 45

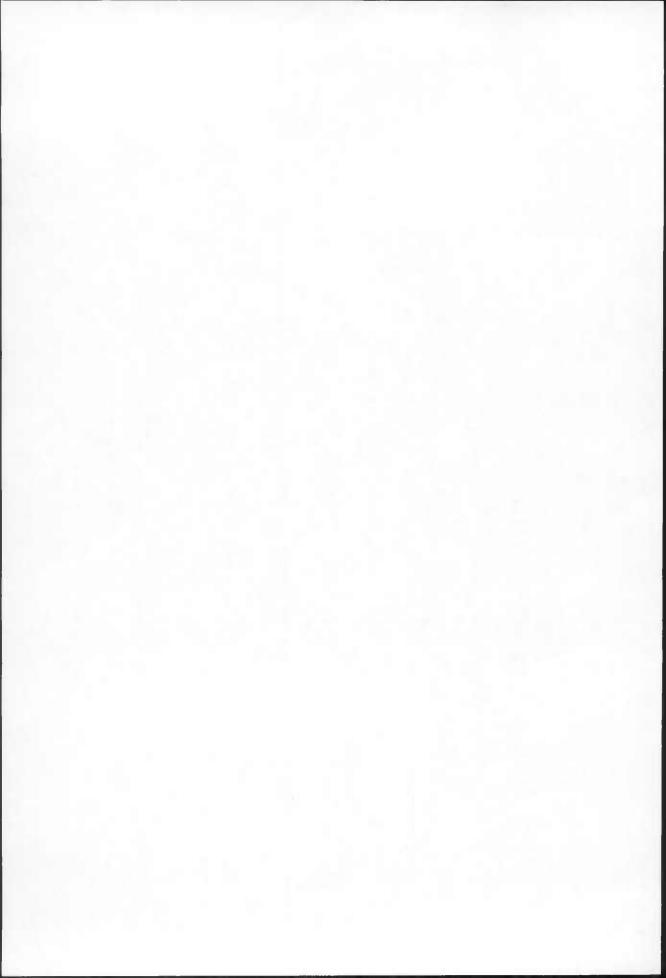


TABLE 14
RECORDS OF WELLS IN SOMERSET COUNTY—explanation of codes

	Aquifer codes	Wat	er-use codes
112PCPC	Surficial aquifer system	С	Commercial
122MOCN	Miocene series	Н	Domestic
122PCMK	Pocomoke aquifer	I	Irrigation
122MNKN	Manokin aquifer	N	Industrial
122CPNK	Choptank aquifer	P	Public supply
124PNPN	Piney Point aquifer	S	Stock
125PLCN	Paleocene aquifer system	U	Unused
217PTMC	Potomac aquifer system	Z	Other

TABLE 14 RECORDS OF WELLS IN SOMERSET COUNTY—Continued

[ft = foot; in. = inch; gsl/min = gallons per minute; (gsl/min)/ft = gsllons per minuts per foot; -- = no dats; m = meesured; + = above land surface; f = flowing well; SWD = Somerset Well Drilling; IWD = Ideal Well Drillers; USGS = U.S. Geological Survey; MD. WRA = Merylend Water Resourcas Administration]

Loce well no.		Permit number	Owner	Contractor	Dsts well con- structed		Depth drillsd (ft)	Diam- ster of cssing (in.)
SO Ad SO Ad SO Ad SO Ad SO Ad	1 2 3 4 5	==	LYON, W D LONG, THOMAS J PINKETT, BRISCOE, JR BLOODSWORTH, FLOYD LARIMER, SUSY M	MUIR WHITE WHITE MUIR	00-00-44 00-00-32 00-00-51 00-00-49 00-00-51	10 10 10 10	40 16 106 53 26	1.5 36 1.5 1.5
SO Ad SO Ad SO Ad SO Ad SO Ad	6 7 8 9	SO-73-1240	BOUNDS, CLAUD MCGRATH, HAROLD STREET, R B REDDEN, J ELMER SOMERSET CO. BOARD OF ED	WHITE WHITE WALLACE	00-00-42 00-00-20 00-00-10 00-00-45 07-25-78	5 10 5 15	14 114 20 130 130	1.5 2 1.75 2
SO Ad SO Ad SO Ad SO Ad SO Ae	12 13 14 15	SO-73-1735 SO-81-0467 SO-81-0707 SO-00-5877	WHITE, TOM PETERS, ARBY FISHER, CHARLES WILES, THOMAS CLARK, M K	KAUFFMAN SWD SWD WHITE	07-17-80 07-30-84 08-26-85 05-00-50	10 10 3 13	66 150 49 150 123	4 2 2 2 2
SO Ae SO Ae SO Ae SO Ae SO Ae	2 3 4 5 6	=======================================	GALE, C STURGIS, CURTIS BENEPEE, ROBERT BEMEPEE, ROBERT BELL, ALDA	WHITE WHITE	00-00-37 00-00-47 00-00-37 00-00-37 00-00-46	5 20 5 5	20 30 130 90 25	1.5 1.5 2 2 1.25
SO Ae SO Ae SO Ae SO Aa SO As	7 8 9 10		GRIFFIN, H L JONES, LESTER CROSWELL, VIRGINIA MUIR, EDNA D SMITH, NORMA	WHITE	00-00-48 00-00-77 00-00-02 05-00-51	15 15 20 20 20	30 147 25 35 50	1.25 1.5 1.5 1.5
SO Ae SO Ae SO Ae SO Ae SO Ae	12 13 14 15 16	SO-00-9644 SO-81-0088	INGERSOL, WALTER BARKLEY, MARTIN HITCH, THORNTON ADKINS, EPHRAIM ADKINS, JEFF	WHITE SWD	00-00-21 00-00-21 04-00-52 07-01-82	10 20 20 10 20	27 27 121 50 65	1.5 1.5 2 4
SO Ae SO Ae SO Ae SO Ae	17 18 19 20 21	SO-81-0266 SO-81-1000 SO-81-0611	WHITE, MILLER SPENCE, HUGH U.S. GEOLOGICAL SURVEY U.S. GEOLOGICAL SURVEY POTTS, TERRY	SWD USGS MD. WRA LARSON WELLS	08-00-83 00-00-78 12-11-86 03-11-87 05-22-85	10 25 20 20 20	35 160 14 140 360	2 2 2 2 2
SO Af SO Af SO Af SO Af SO Af	1 2 3 4 5		TAYLOR, S IRVING CHRISTOPHER, E WHITE, M POLLITT, L E HARMON, PIERCE	TAYLOR OWNER SHOCKLEY SHOCKLEY	00-00-50 00-00-51 00-00-49 00-00-43	30 40 45 40	34 37 32	1.25 1.25 1.25 1.5
SO Af SO Af SO Af SO Af SO Af	6 7 8 9	==	POLLITT, ELMER WILLEY, J E ARMSTRONG, EDDIE MERCER, S A WILLEY, HERBERT	CAMPBELL CAMPBELL MURRAY	00-00-41 00-00-47 00-00-47 00-00-51 02-00-52	45 35 40 30 25	24.0 32 30 30 35.0	1.5 1.5 1.5 1.5
SO Af SO Af SO Af SO Af SO Af	11 12 13 14 15		BARKLEY, A C JONES, MELVIN JONES, OSCAR PRYOR, DAVID SNELLING, PEARL	CAMPBELL	00-00-20 00-00-44 00-00-24 00-00-50 00-00-47	25 40 45 30 30	33 16 32 35 33	1.5 1.5 1.5 1.75 1.75
SO Af	16 17		U.S. GEOLOGICAL SURVEY U.S. GEOLOGICAL SURVEY	HARDIN ASSOC HARDIN ASSOC	11-17-80 11-17-80	30 30	90 78	2 2
SO Af SO Af SO Af	18 19 20	SO-81-0129 SO-73-0899 SO-81-0091	CHESAPEAKE LUMBER CO TUCKER, HOWARD EDEN MARKET	DASHIELL DRLNG IWD LARSON DRLNG	08-17-82 05-23-77 06-23-82	30 30 30	70 50 180	4 2
SO Af SO Af SO Af SO Af SO Af	23 24	SO-81-0520 SO-73-0586 SO-71-0069 SO-81-0302 SO-81-0324	EASTERN SHORE OIL EDEN MOBILE HOME VILLAGE EDEN MOBILE HOME VILLAGE EDEN MOBILE HOME VILLAGE EDEN MOBILE HOME VILLAGE	LARSON WELLS IWD IWD DASHIELL DRLNG DASHIELL DRLNG		30 30 30 30 30	190 195 45 255 255	2 4 4 4 4

Bottom		Bottom										
of caeing or cesed inter- val (ft)	Diameter of screen (in,)	of ecraen or screened inter- vel (ft)	Aquifer code	Water level (ft)	Dete weter level mees- urad	Drew- down (ft)	Die- cherge (gel/ min)	Pumpins period (hours)	Spe- cific cepe- city [(gel/ min)/ ft]		Loce well no.	
			112PCPC							H	SO Ad	1
16			112PCPC							H	SO Ad	2
			122MNKN 112PCPC							H	SO Ad	3
			112PCPC							H	SO Ad	5
			112PCPC							н	SO Ad	6
			122MNKN							H,S	SO Ad	7
			112PCPC							H,S	SO Ad	8
130 105		125	122MNKN 122MNKN	8	07-25-78		60	2	15.0	H,S	SO Ad	9 11
	,											
36 130	2	66 150	112PCPC	8 11	07-17-80 07-30-84	8 7	60 17	1.5	7.5	H	SO Ad	12
130	2	130	122MNKN 112PCPC	. 4m	04-14-87				4	U	SO Ad	14
130	2	150	122MNKN	14	08-26-85	1	50	1	50.0	H	SO Ad	15
115	2	123	122MNKN	3	05-00-50		12	4		H,S	SO Ae	1
			112PCPC							H	SO Ae	2
			112PCPC 122MNKN	+10f	02-18-52					H,S	SO Am	3
			122MNKN	+101	02-18-52					IJ	SO As	5
			112PCPC							H	SO Ae	6
			112PCPC							H	SO Ae	7
137		147	122MNKN							H	SO Ae	8
			112PCPC							H	SO Ae	9
			112PCPC 112PCPC							H,S	SO Ae	10
			112PCPC 112PCPC	3	00-00-52					H	SO Ae	12
113	2	121	122MNKN	1	04-00-52					H	SO A	14
27	4	50	122PCMK	15	07-01-82	5	60	2	12.0	Ï	SO Ae	15
			122PCMK	11.7m	04-14-87					U	SO Ae	16
25	2	35	112PCPC	12	08-00-83	3	10	1	3.3	I	SO As	17
			122MNKN							H	SO Ae	18
9 130	2 2	14 140	112PCPC 122MNKN	10.2m 8.3m	04-14-87 04-17-87					U	SO Ae	19
340	2	360	122CPNK	7	05-22-85	20	20	3	1.0	H	SO A	21
			112PCPC							Н	SO Af	1
34			112PCPC	8.7m	02-06-52					H,S	SO Af	2
			112PCPC							H,S	SO Af	3
			112PCPC 112PCPC	2	00-00-49					H,S	SO Af	5
										H,S		3
			112PCPC	3.5m	02-04-54					H,S	SO Af	6
			112PCPC 112PCPC							H	SO Af	7
			112PCPC							H,S	SO Af	8
			112PCPC	5	02-00-52					H,S	SO Af	10
			112PCPC							Н	SO Af	11
			112PCPC							H,S	SO Af	12
			112PCPC							H,S	SO Af	13
			112PCPC							H,S	SO Af	14
			112PCPC							H,C	SO Af	15
36.5 38	5 2 2	16.5- 26.5 18 - 28	122PCPC 112PCPC	3.4m 5.9m	06-10-81 06-10-81					U	SO Af	16 17
0.0	,	7.0	1000000	5.2m	04-14-87		4.05					4 -
60	4	70	122PCMK	8	08-17-82	5	100	1	20.0	C	SO Af	18
40 170	2	50 180	112PCPC 122MNKN	15	05-23-77 06-23-82	16	80 50	3	5.0 16.7	S,Z	SO Af	19 20
180	2	190	122MNKN	17	10-29-84	3	45	1	15.0	C	SO Af	21
185	4	195	122MNKN	7	04-16-76	18	100	1	15.0 5.66	C P	SO Af	22
35	4	4.5	112PCPC	5	05-25-71	20	70	3	3.5	P	SO Af	
235	4	255	122MNKN	14	08-17-83	6	35	1	5.8	P	SO Af	24
235	4	255	122MNKN	14	09-07-83	7	60	1	8.6	P	SO Af	25

TABLE 14 RECORDS OF WELLS IN SOMERSET COUNTY—Continued

[ft = foot; in. = inch; gal/min = gallons per minute; (gal/min)/ft = gallons per minute per foot; -- = no data; m = measured; + = above land surface; f = flowing well; SWD = Somerset Wall Drilling]

Loca well no.		Parmit numbar	Ownar	Contractor	Date well con- structed	Alti- tuda of land sur- face (ft)	Depth drilled (ft)	Diam- eter of casing (in.)
SO Bb	1	SO-00-1591	SOMERSET SEAFOOD CO	CUSICK	09-13-47	5	814	2.5
SO Bb SO Bb SO Bb SO Bb	2 3 4 5	SO-00-0903 SO-00-0382 SO-00-0904	PRICE, THOMAS WHITE, HAROLD BEECHAM, J T CORBIN, G M	TODD WHITE TODD	10-17-46 05-15-46 10-20-46	5 5 5	144 122 140 145	1.5 2 1.5 3
SO Bb SO Bb SO Bb SO Bb	6 7 8 9 10	SO-73-0603 SO-81-0024	PARKS, W S JONES, ELDRIDGE BROWN, CLARENCE UNITED METHODIST CHURCH PHILLIPS 66-EAST SHORE OIL	 DASHIELL DRLNG LARSON DRLNG	00-00-42 00-00-42 00-00-44 11-06-75 12-01-81	5 5 10 10	10 9 12 140 115	1.25 1.25 1.5 2
SO Bb SO Bb SO Bb SO Bb SO Bb	11 12 13 14 15	SO-73-0802 SO-73-1088 SO-67-0020 SO-73-1880 SO-73-1490	ROCK CREEK METHODIST CH TAYLOR, ALBERT WILLING, ELDON, SR LAST CHANCE MARINA ISLAND SEAFOOD INC	FORD HALL MARSHALL SWD KELLY	09-10-76 10-28-77 08-03-66 03-30-81 06-26-79	10 10 10 10	135 15 111 150 130	2 2 2 4
SO Bb SO Bb SO Bb	16 17 18 19	SO-67-0070 SO-73-0296 SO-81-0586	SOMERSET SEAFOOD CO ISLAND SEAFOOD INC EASTERN SHORE OIL CO SOMERSET SEAFOOD CO	FORD FORD SWD S. SHANNAHAN	10-26-66 05-28-74 06-19-85 12-01-58	5 10 10 5	126 133 140 720	2 2 2 4 2.5
SO Bb	20	SO-81-0675	RUSSELL, SYLVAN	SWD	08-10-85	8	150	2
SO Bb SO Bb SO Bb SO Bb SO Bc	21 22 23 24 1	SO-81-0743 SO-81-0749 SO-81-0900 SO-81-0820 SO-00-8344	WINDSOR, ROY SUHR, FRED LOU, FRANK SCOTIS COVE MARINA DASHIELL, ELMER	SWD SWD SWD SWD WHITE	11-01-85 10-25-85 06-17-86 04-04-86 08-00-51	8 8 8 5	140 160 160 150	2 2 2 2 2
SO Bc SO Bc SO Bc SO Bc SO Bc	2 3 4 5 6	SO-00-5242 SO-00-5186 SO-00-5243 SO-00-0842 SO-00-3753	BOZMAN, HARRY L STARK, W G WEBSTER, BAIN D WALLACE, VAUGHN MCDORMAN, WALTER F	WHITE WHITE TODD WHITE	02-15-50 02-13-50 02-20-50 09-23-46 03-30-49	5 5 5 5	105 105 135 122 127	2 2 2 1.5 2
SO Bc	7	SO-00-1583	MESSICK, HENRY	FARLOW	06-28-47	5	139	2
SO Bc SO Bc SO Bc	8 9 10 11	SO-00-6147 SO-00-5985 SO-00-2098	CAREW, BROOKS WEBSTER, MASON WEBSTER, MASON HOPKINS, FORD	CUSICK WHITE WHITE WHITE WHITE	06-30-50 00-00-50 06-00-50 12-30-47	5 5 5 5	147 132 95 90	1.5 2 2 2
SO Bc SO Bc SO Bc	12 13 14	SO-00-3207 SO-01-3230	WALLACE, HARWOOD JONES, MONROE SOMERSET CO. BOARD OF ED	WHITE	10-07-48 1882 09-00-53	5 5 5	135 6.5 132	2 24 2
SO Bc SO Bc	15 16	SO-73-0366 SO-81-0818	MACEDONIA UNITED METHODIST O HOWARD, JOE	CH FORD SWD	11-05-74 06-12-86	5	140 120	2 4
SO Bc SO Bc SO Bc SO Bd SO Bd	17 18 19 1 2	SO-81-0799 SO-81-0719 SO-81-0690 SO-00-0383 SO-00-0384	DUNN, GREG JONES, ROGER CULVER, DONALD MCINTYRE, ROSS MCINTYRE, NEARY	SWD SWD SWD WHITE WHITE	02-28-86 09-09-85 08-15-85 05-11-46 05-10-46	5 5 5 10 10	120 140 125 136 136	2 2 4 2 2
SO Bd SO Bd SO Bd SO Bd SO Bd	3 4 5 6 7	SO-00-0385 SO-00-8132 SO-00-8345 SO-00-6980 SO-00-0734	DASHIELL, JENNIE H EISNOR, EDNA DUNTON, EDGAR L HORNER, JOHN W PARKS, JAMES	WHITE WHITE WHITE WHITE WHITE	05-07-46 07-00-51 09-00-51 11-00-50 09-21-46	10 5 15 5 5	139 107 150 101 94	2 2 2 2 2
SO Bd SO Bd SO Bd SO Bd	8 9 10 11	SO-00-3208 SO-00-0735 SO-00-6060 SO-00-9181	CAUSEY, HARRY SIMMS, CLARK I DASHIELL, HERMAN KOHLHEIM, R J	WHITE WHITE WHITE CUSICK	10-08-48 09-24-46 04-00-50 12-03-51	5 5 5	104 98 122 178	2 2 2 1.5
SO Bd	12	SO-00-6337	KOHLHEIM, R J	CUSICK	07-29-50	10	179	1.5

Bottom of cssing or cassd inter- val (ft)	Diameter of screen (in.)	Bottom of screen or screened inter- val (ft)	Aquifer coda	Wster lavel (ft)	Dete weter levs1 mess- ured	Draw- down (ft)	Dis- charge (gal/ min)	Pumping period (hours)		Usa of water	Locsl well no.	
0-546 546-609 659-671		609 -659	124PNPN	+.5f 1	09-13-47 06-27-50		30	14		С	SO Bb	1
144 122 140	1.5	136 -142 110.5-120.5 132 -138	122MNKN 122MNKN 122MNKN 122MNKN	8 3 8	10-17-46 05-15-46 10-20-46	==	10 40 10	8 2 8		H U H H	SO Bb SO Bb SO Bb	2 3 4 5
132 105	2 2	140 115	112PCPC 112PCPC 112PCPC 122MNKN 122MNKN	 11 7	11-06-75 12-01-81	6 5	30 25	1 3	5.0	H H T	SO Bb SO Bb SO Bb SO Bb	6 7 8 9
120 12 105 130 105	2 2 2 4	135 15 111 150 125	122MNKN 112PCPC 122MNKN 122MNKN 122MNKN	5 3 7 10 6	09-10-76 10-28-77 08-03-66 03-30-81 06-26-79	2 2 11 4 19	10 4 14 8 15	2 1 1 1 1 3	5.0 2.0 1.3 2.0	T H P C C	SO Bb SO Bb SO Bb SO Bb	11 12 13 14 15
112 113 120 151 524	2 2 2	126 133 140	122MNKN 122MNKN 122MNKN 124PNPN	4.1m 6 5	04-16-87 05-28-74 06-19-85	2 5	10 15 30	2 3 1	7.5 6.0	C C,N C	SO Bb SO Bb SO Bb	16 17 18 19
130	2	150	122MNKN	5	08-10-85	5	25	1	5.0	Н	SO Bb	20
120 140 140 130 123	2 2 2 2 2	140 160 160 150 131	122MNKN 122MNKN 122MNKN 122MNKN 122MNKN	5 6 10 3	11-01-85 10-25-85 06-17-86 04-04-86 08-00-51	5 2 2 1	30 35 60 50	1 1 1 1 5	6.0 18 30 50	H H C H	SO Bb SO Bb SO Bb SO Bd SO Bc	21 22 23 24
105 105 135 116 115	2 2 2 2 2	95 -103 95 -103 125 -133 122 126	122MNKN 122MNKN 122MNKN 122MNKN 122MNKN	3 3 4 8	02-15-50 02-13-50 02-20-50 09-23-46 03-30-49		10 10 8 10 15	2 3 6 8 2		H H H H	SO Bc SO Bc SO Bc SO Bc SO Bc	2 3 4 5 6
139 137 87 78	2 2 2 2	129.5-137.5 147 95 90	122MNKN 122MNKN 122MNKN 122MNKN 122MNKN	1 5.5m 2 8 1.5	06-28-47 01-06-52 06-30-50 00-00-50 06-00-50	1	15 16 20	2 2 2 1	15	U H P H H	SO Bc SO Bc SO Bc SO Bc	7 8 9 10 11
123	2	135 133	122MNKN 112PCPC 122MNKN	.5 2.0m 1.5 1.8m	10-07-48 03-13-52 09-00-53 01-19-54		17 12	2 3	==	H H,S	SO Bc SO Bc SO Bc	12 13 14
130 100	2 4	140 120	122MNKN 122MNKN	5 7 5,5m	11-05-74 06-12-86 12-11-86	5	10 50	2	5.0	H	SO Bc	15 16
100 120 100 126 136	2 2 2 2	120 140 125 136 126.5-134.5	122MNKN 122MNKN 122MNKN 122MNKN 122MNKN	3 8 7 2 4	02-28-86 09-09-85 08-15-85 05-11-46 05-10-46	1 1 3 	40 45 40 50	1 1 1 2	40 45 13	H H H H	SO Bc SO Bc SO Bc SO Bd SO Bd	17 18 19 1 2
139 99 138 93 86	2 2 2 2 2	127.5-137.5 107 150 101 94	122MNKN 122MNKN 122MNKN 122MNKN 122MNKN	3 +1.5 3 6 2	05-07-46 07-00-51 09-00-51 11-00-50 09-21-46		40 12 10 15	2 3 1.5		H U H	SO Bd SO Bd SO Bd SO Bd SO Bd	3 4 5 6 7
92 92 112 158	2 2 2 1.5	104 98 122 178	122MNKN 122MNKN 122MNKN 122MNKN	3	10-08-48 09-24-46 04-00-50 12-03-51	 2	18 12 15 22	2 2 3 2	 11	H H H U	SO Bd SO Bd SO Bd SO Bd	8 9 10 11
164	1.5	179	122MNKN	2.8m 1.5	01-16-52 07-29-50	1	20	2	20	Н	SO Bd	12

[ft = foot; in. = inch; gal/min = gallons per minute; (gel/min)/ft = gallona per minute per foot; -- = no data; m = measured; + = above land aurfaca; SWD = Someraat Well Drilling; CD&P = Coastal Drilling and Pump; SAW = Shannahan Art. Well; USGS = U.S. Geological Survay; MD. WRA = Maryland Weter Reaourcea Administration]

Loce well no.		Permit number	Owner	Contractor	Data wall con- structed		Depth drilled (ft)	Diam- eter of caaing (in,)
SO Bd SO Bd SO Bd SO Bd	13 14 15 16	SO-00-2097 SO-00-0457 SO-00-0456 SO-00-4988	ST. PETERS CHURCH TRUSTEES MELSON, MATT REESE, J B FARLOW, JAMES M	WHITE FARLOW WHITE WHITE	11-22-47 05-17-41 05-22-46 02-00-50	5 5 5 10	145 143 144 136	2 2 2 2 2
SO Bd	17		FARLOW, JAMES M NOBLE, TOM	WHITE	03-00-52	10	148	2
SO Bd	18		WILLING, DENWOOD		00-00-41	5	22	1.25
SO Bd	19		SEARS, LEANORD MADDOX, OSCAR		00-00-49 12-00-51	10	51 25	1.5
SO Bd	21		TILGHMAN, OTHO		00-00-42	5	150	2
SO Bd	22		BOZMAN, WESLEY		00-00-44	5	38.5	1.5
SO Bd	23		BOZMAN, JAMES	TODD	00-00-40	5	160	1.5
SO Bd	24 25		DIZE, THOMAS HALL, GEORGE	JARRETT	00-00-44	5	160 160	1.5
SO Bd	26		DASHIEL, NATT		00-00-30	5	20	1.25
SO Bd	27		FITZGERALD, HARRY		00-00-12	5	25	1.5
SO Bd	28	SO-00-1789	WHITE. CLARENCE E	CUSICK	09-18-47	5	136	1.5
SO Bd	29		WHITE, CLARENCE E NOBLE, HARRY		00-00-44	5	165	2
SO Bd	30	SO-67-0065	JONES, EDGAR MONIE CREEK GUN CLUB	FORD	00-00-03 10-21-66	5	80 117	1.5
SO Bd	33	SO-81-0343	MT. VERNON FIRE CO	SWD	06-05-84	10	140	2
SO Bd	34	SO-71-0011	ASBURY METHODIST CHURCH	FORD	08-26-70	10	105	2
SO Bd	35	SO-73-1362	MT. VERNON PACKING CO	SWD	10-08-78	5	120	2
SO Bd	36	SO-81-0178	WHITE, TOM	CD&P	03-17-83	10	150	8
SO Bd	37		WHITE, TOM	CD&P	00-00-83	10	75	8
SO Bd	38	SO-73-0985	MUIR, JENNINGS	DASHIELL DRLNG	04-13-77	10	180	2
SO Bd SO Bd SO Bd SO Bd	39 40 41 42 43	SO-71-0030 SO-67-0056 SO-81-0999 SO-81-0832	VENTON METHODIST CHURCH ST. PETERS METHODIST CH U.S. GEOLOGICAL SURVEY U.S. GEOLOGICAL SURVEY BOUNDS, MARY	FORD FORD USGS MD. WRA DASHIELL	09-23-70 10-22-66 11-01-86 03-13-87 04-15-86	10 5 5 5 5	148 105 15 138 150	2 2 2 15 2
SO Bd SO Bd SO Be SO Be	44 45 46 1 2	SO-81-0814 SO-81-0912 SO-81-0854	LAWRENCE, MORRIS ANDERSON, LAWKFORD BEDSWORTH, JERRY U.S. GEOLOGICAL SURVEY TOWN OF PRINCESS ANNE	SWD SWD SWD USGS SAW	04-01-86 07-20-86 05-06-86 08-16-49 00-00-45	8 5 8 20 20	170 160 160 22.5	2 2 2 2 1.25
SO Be	3	SO-00-8260	REYNOLDS, ROBERT	cusick	07-21-51	20	200	1.5
SO Be SO Be SO Be	5 6 7	SO-00-2498 SO-00-7351 SO-00-0917 SO-00-1064	SMITH, ROY W LONG, EARL BENSON, FRED E BOZMAN, HERMAN	WHITE WHITE CUSICK WHITE	07-26-48 05-00-51 11-08-46 12-28-46	20 10 20 15	187 151 196 183.5	2 2 2 3 2 2
SO Ba	8	SO-00-0945	KEAN, DAVID B	CUSICK	11-15-46	15	196	2
SO Ba SO Be	10	SO-00-7350 SO-00-0665	POLLÍTT, EDWARD CARTER, HARRY	WHITE WHITE	02-00-51 09-14-46	10 10	203 194	2 3 2
SO Be SO Be	11 12	SO-00-6522 SO-00-5566	STROBLE, ERVIN E PORTER, JAMES	CUSICK WHITE	08-21-50 05-00-50	15 15	188 170	1.5
SO Be	13	SO-00-8870	GORDY, FRED O	CUSICK	10-24-51	20	208	1.5
SO Be	14 15	SO-00-5410	PUSEY, VADOR MRS	CUSICK	04-05-50	17	212	1.5
SO Be	16	SO-00-7209 SO-00-7595	PUSEY, ELLA MRS CARROW, T LESTER MRS	CUSICK CUSICK	01-03-51 05-01-51	15 15	198 188	1.5
SO Be	17	SO-00-5567	SIMPKINS, DOUGLAS	WHITE	05-05-50	20	180	2
SO Be	18	SO-00-8876	PINTO, ROBERT	CUSICK	11-21-51	15	200	1.5
SO Be	19	SO-00-5916	RUSSELL, HARVEY	WHITE	05-00-50	15	166	2
SO Be	20		BLACK, MRS DRYDEN, ALTON	MUIR	00-00-51	15 20	15 30	1.5
SO Be	22		MEREDITH, C E		00-00-47	10	22	1.5

ces ces	er-	Diam- eter of ecreen (in.)	Bottom of ecreen or screened inter- vel (ft)	Aquifer code	Weter level (ft)	Dete weter level mees- ured	Drew- down (ft)	Dis- cherge (gel/ min)	Pumping period (hours)	Spe- cific cepe- city [(gel/ min)/ ft]	Use of weter	Locel well no.	
	145 143 138 126	2 2 2 2	135.5-143.5 131.5-141.5 144 134	122MNKN 122MNKN 122MNKN 122MNKN 122MNKN	+.5 3 4 2	11-22-47 05-17-46 05-22-46 02-00-50		15 30 30 10	1.5 2 2 2		T H H H	SO Bd SO Bd SO Bd SO Bd SO Bd	13 14 15 16 17
				112PCPC 122PCMK 112PCPC 122MNKN 112PCPC		=======================================					н н н н	SO Bd SO Bd SO Bd SO Bd SO Bd	18 19 20 21 22
				122MNKN 122MNKN 122MNKN 112PCPC 112PCPC	+1	00-00-40					H,S H H H	SO Bd SO Bd SO Bd SO Bd SO Bd	23 24 25 26 27
	126 109 105	1.5	136 117 140	122MNKN 122MNKN 122MOCN 122MNKN 122MNKN	2 +.5 +1.5 1	09-18-47 01-19-54 00-00-03 10-21-66 06-05-84	3	28 10 10	2 2 1	9.3	H U Z T	SO Bd SO Bd SO Bd SO Bd SO Bd	28 29 30 32 33
+1 35	97 100 - 15 - 90	2 8 8	105 120 15 - 35 90 -150	122MNKN 122MNKN 112PCPC 122MNKN	7 8 8	08-26-70 10-08-78 03-17-83	3 4 22	10 11 700	2 1 4	3.3 2.8 32	T C I	SO Bd SO Bd SO Bd	34 35 36
	172	2	180	112PCPC 122MNKN 122MNKN	4.0m	06-26-86	8	15	1	1.9	U C	SO Bd	37
	140 97 10 80 140	2 2 2 2 2	148 105 15 90 150	122MNKN 122MNKN 112PCPC 122MNKN 122MNKN	7 1 2.2m 5.1m	09-23-70 10-22-66 04-14-87 04-14-87 04-15-86	3	10 10 20	2 2 1	3.3	T Z U U H	SO Bd SO Bd SO Bd SO Bd SO Bd	39 40 41 42 43
	150 140 140 22.5	2 2 2	170 160 160	122MNKN 122MNKN 122MNKN 112PCPC 122PCMK	6 9 15 3.9m	04-01-86 07-20-86 05-06-86 08-16-49 00-00-45	1 3 13	60 75 30 	1 1 1	60 75 10 8.1	H S H U P	SO Bd SO Bd SO Be SO Be	44 45 46 1 2
0 105	180 187 138 168 -105 -170	1.5 2 2 2 2	200 173.5-185.5 150 196	122MNKN 122MNKN 122MNKN 122MNKN	2 4 1 3 7	07-21-51 07-26-48 05-00-51 11-08-46 12-28-46	7 1	30 8 16 20 18	3 4 3 4	20	H H,S H,S H,S	SO Be SO Be SO Be SO Be	3 4 5 6 7
182	-183.5 174 190		196 202	122MNKN 122MNKN	3	11-15-46 02-00-51	.5	20 10	1 4	40	H H,S	SO Be	8
84	- 84 -180.5 5-194		180.5-192.5	122MNKN	3	09-14-46		12	3		H,S	SO Be	10
	178 158	1.5	188 170	122MNKN 122MNKN	6.0	08-21-50 05-00-50	2	20 12	2	10	H	SO Be	11 12
	193 197 183 173 168	1.5 1.5 1.5 1.5	208 212 198 188 180	122MNKN 122MNKN 122MNKN 122MNKN 122MNKN	4 3 2	10-24-51 04-05-50 01-03-51 05-01-51	4 5	22 25 25 12	3 2 2 6	22 6.3 5.0	H H,S H C,H	SO Be SO Be SO Be SO Be	13 14 15 16 17
	180 154 	1.5	200 166 	122MNKN 122MNKN 112PCPC 112PCPC 112PCPC	5 4	11-21-51 05-00-50 	2	30 10 	3 3 	15	H,S H H,S H,S	SO Be SO Be SO Be SO Be	18 19 20 21 22

[ft = foot; in. = inch; gal/min = gallons per minute; (gal/min)/ft = gallons per minute per foot; -- = no data; m = meaaured; f = flowing well; SWD = Someraet Well Drilling; IWD = Ideal Well Drillera]

Local well no.		Permit number	Owner Co	ontrector	Date well con- atructed	Alti- tude of land sur- face (ft)	Depth drilled (ft)	Diam- eter of casing (in.)
SO Be SO Be SO Be SO Be SO Be	23 24 25 26 27		ADAMS, M H LONG, E W MILLER, J B BRIDDELL, A E ADAMS, M	WHITE 	00-00-48 00-00-12 00-00-47 00-00-12 00-00-51	15 15 10 10 15	210 20 20 20 20	1.5 1.25 1.5 36 1.5
SO Be SO Be SO Be SO Be	28 29 30 31 32		TOWN OF PRINCESS ANNE TOWN OF PRINCESS ANNE TOWN OF PRINCESS ANNE TOWN OF PRINCESS ANNE	CANNON CANNON CANNON CANNON	00-00-18 00-00-18 00-00-18 00-00-18		20 32.3 26.3 27.9 23.5	1.5
SO Be SO Be SO Be SO Be SO Be	33 34 35 36 37		TOWN OF PRINCESS ANNE	CANNON CANNON CANNON CANNON CANNON	00-00-18 00-00-18 00-00-18 00-00-18 00-00-18	20 20 20	34.5 33.0 34.5 34.9 26.8	
SO Be SO Be SO Be SO Be SO Ba	38 39 40 41 42		TOWN OF PRINCESS ANNE UNIV. OF MD EASTERN SHORE UNIV. OF MD EASTERN SHORE SUPPLEE, WILLS, JONES SMITH, E MACE	CANNON CUSTIS CUSTIS PENTZ WHITE	00-00-18 00-00-29 00-00-29 04-00-42 00-00-29	10	25.1 196 204 60 184	2 2 2 3 2
SO Be SO Be SO Be SO Be	43 44 45 46 47	SO-01-0639 SO-00-9787 SO-00-9913 SO-00-9848 SO-00-9849	FITZGERALD, JOHN H DYKES, HERMAN LAYFIELD, PHILLIP MRS POWELL, ELMER CARTER, WILLIAM	CUSICK CUSICK WHITE WHITE WHITE	10-01-52 04-30-52 06-00-52 05-00-52 05-00-52	20 20 15 15	203 189 213 209 204	1.5 1.5 2 2
SO Be SO Be SO Be	48 49 50 51	SO-01-1188 SO-01-1924 SO-01-1924	TAYLOR, EUGENE TOWN OF FRINCESS ANNE TOWN OF FRINCESS ANNE TOWN OF FRINCESS ANNE	WHITE KELLY WELL SYDNOR P&W SYDNOR P&W	11-00-52 00-00-28 03-01-53 08-14-53	10 20 20 20	188 64 419 214	2 24 -18 16 10
SO Be	52	SO-01-2429	TOWN OF PRINCESS ANNE	SYDNOR P&W	05-15-53	20	77	8
SO Be SO Be SO Ba SO Be	53 54 55 56 57	SO-65-0054 SO-67-0126 SO-81-0558 SO-81-0471 SO-81-0630	CUSTOM PET FOOD PACKERS, II SOMERSET CO. SANITARY COMM STATE OF MARYLAND SOMERSET CO. SANITARY COMM STATE OF MARYLAND	NC LAYNE-ATLNTC IWD LAYNE-ATLNTC DELMARVA DRLNG LAYNE-ATLNTC	11-00-64 06-08-67 04-18-85 05-14-85 05-30-85	5 12 15	262 260 250 255 220	6 10 12 4
SO Be SO Be SO Ba SO Ba SO Be	58 59 60 61 62	SO-73-0628 SO-73-1453 SO-73-0690 SO-73-0090 SO-73-0991	SOMERSET ANIMAL HOSPITAL PERDUE FARMS INC DYKES FRUIT AND PRODUCE CHOPTANK ELECTRIC COOP CHELTONS'S WELDING SERVICE	FORD SWD FORD IWD FORD	01-21-76 03-26-79 04-21-76 07-26-73 07-08-77	20 20	200 210 200 85 101	2 2 2 4 2
SO Be SO Be SO Be SO Be SO Be	63 64 65 66 67	SO-73-0804 SO-73-0883 SO-73-1115 SO-68-0009 SO-70-0065	CORBETT BREEDERS, INC CORBETT BREEDERS, INC LAYFIELD AUTO PARTS SMITH, WILLARD CUSTOM PET FOOD PACKERS, IN	DASHIELL DRLNG DASHIELL DRLNG FORD FORD TORD		20 20	194 194 180 173 176	2 2 2 2 2
SO Be SO Be SO Be	68 69 70 71	SO-70-0066 SO-70-0081 SO-66-0055 SO-81-0607	CUSTOM PET FOOD PACKERS, II CUSTOM PET FOOD PACKERS, II FIRST BAPTIST CHURCH MOUNTAIRE HATCHERY	FORD FORD SWD	03-02-70 04-28-70 12-29-65 05-17-85	10 10 20	171 180 188 230	2 4 2 4 3
SO Be	72	SO-73-1948	NICHOLS, THOMAS	SWD	08-30-81	20	90	2
SO Be SO Be SO Be SO Be	73 76 77 78 79	SO-73-0018 SO-73-0763 SO-81-0760 SO-73-0645 SO-73-0837	TOWN OF PRINCESS ANNE UNIV. OF MD EASTERN SHORE UNIV. OF MD EASTERN SHORE SOMERSET CO. SANITARY COMM WHITTINGTON FARMS	IWD IWD DELMARVA DRLNG IWD KAUFFMAN	09-07-73 08-30-76 03-04-86 02-27-76 04-26-78		51 200 75 195 63	4 6 6 8 15

Bottom of casing or cesed inter- vel (ft)	Diam- eter of screen (in,)	Bottom of screen or screened inter- vel (ft)	Aquifer code	Weter level (ft)	Dete weter level mees- ured	Drew- down (ft)	Dis- cherge (gel/ min)	Pumping period (hours)	Spe- cific cepe- city [(gel/ min)/ ft]	Use of weter	Loce: well no.	1
195	1.5	210	112PCPC 112PCPC 112PCPC 112PCPC 112PCPC							H H,S H,S	SO Be SO Be SO Be SO Be	23 24 25 26 27
		==	112PCPC 112PCPC 112PCPC 112PCPC 112PCPC	13.0m 11.3m 11.8m 15.5m	01-15-52 01-15-52 01-15-52 01-15-52					H U U U	SO Be SO Be SO Be SO Be	28 29 30 31 32
			112PCPC 112PCPC 112PCPC 112PCPC 112PCPC	15.4m 15.7m 15.7m 15.7m 14.4m	01-15-52 01-15-52 01-15-52 01-15-52 01-15-52					U U U U	SO Be SO Be SO Be SO Be	33 34 35 36 37
192 200	2 2 	196 204	112PCPC 122MNKN 122MNKN 122PCMK 122PCMK	13.9m 11 11 42.6m	01-15-52 00-00-29 00-00-29 04-15-87		30			U U U N	SO Be SO Be SO Be SO Be	38 39 40 41 42
188 174 203 201 191	2 2 2 2	203 189 213 209 204	122MNKN 122MNKN 122MNKN 122MNKN 122MNKN	6 5 6 2.5 6	10-01-52 04-30-52 06-00-52 05-00-52 05-00-52	6	28 20 12 10	2 2 3 5 2	28	H,S H,S H H	SO Be SO Be SO Be SO Be	43 44 45 46
172 34 .5-163	2	188 	122MNKN 122PCMK 122CPNK	3 10 7	11-00-52 00-00-28 08-14-53	23	10 435 350	4 24	19	H P U P	SO Be SO Be SO Be	48 49 50 51
0 -160.5 54.5	10	160.5-213.5	122MNKN 122PCMK	13	05-15-53	40	80	12	2.0	P	SO Be	52
160 165 190 170	6 10 12 4	205 240 250 220	122MNKN 122MNKN 122MNKN 122MNKN 122MNKN	27 32 28 25 29.7m	06-08-67 06-11-85 05-29-85 05-30-85 04-15-87	78 111 117 50	25 135 128 500 50	5 30 48 24 2.5	1.7 1.2 4.3 1.0	U U P U	SO Be SO Be SO Be SO Be	53 54 55 58 57
180 180 180 75 86	2 2 2 4 2	200 210 200 85 101	122MNKN 122MNKN 122MNKN 122PCMK 122PCMK	16 16 16 18	01-21-76 03-26-79 04-21-76 07-26-73 07-08-77	6 194 4 7 4	10 12 10 20 10	2 2 3 4 3	1.7 .06 2.5 2.9 2.5	00000	SO Be SO Be SO Be SO Be SO Be	58 59 60 61 82
186 186 160 157 156	2 2 2 2	194 194 180 173 176	122MNKN 122MNKN 122MNKN 122MNKN 122MNKN	9 17 12 35	07-13-76 12-21-77 11-04-77 08-21-67 03-02-70	5 5 6 25	20 20 10 10	1 1 2 2 2	4.0 4.0 1.7 	CCC	SO Be SO Be SO Be SO Be	63
151 157 172 170 190	2	171 177 188	122MNKN 122MNKN 122MNKN 122MNKN	35 20 22 25	03-02-70 04-28-70 12-29-65 05-17-85	25 25 10	10 20 14 100	2 6 2 24	. 40 . 80	N N T N	SO Be SO Be SO Be	61 71 7
60	2	90	122PCMK	10	08-30-81	2	30	1	15	I	SO Be	72
41 - 42 151 63	6 6 8 15	51 68 191 53	122 PCMK 122MNKN 122 PCMK 122MNKN 122 PCMK	12 9.8m 18 5	09-07-73 04-16-87 02-27-76 04-26-78	18 10.4 77 33	100 4 33 120 800	3 1 12 38	5.6 3.2 6 24	Z I I P I	SO Be SO Be SO Be SO Be	73 76 73 78

TABLE 14
RECORDS OF WELLS IN SOMERSET COUNTY—Continued

[ft = foot; in. = inch; gal/min = gallons per minute; (gal/min)/ft = gallons per minute per foot; -- = no data; m = measured; + = above land surface; SWD = Somerset Well Drilling; CD&P = Coastal Drilling and Pumping; USGS = U.S. Geological Survey; IWD = Ideal Well Drillera}

	Loca Well no.	L	Permit number	Owner	Contractor	Date well con- structed	Alti- tude of land sur- face (ft)	Depth drilled (ft)	Diam- eter of casing (in,)
SO	Be Be	80 81 82	SO-73-0839 SO-73-1893 SO-67-0052	WHITTINGTON FARMS HOUSE OF JACOBS	KAUFFMAN SWD	04-28-77 04-21-81	20 20	63 165	15 2
SO	Be Be	83 84	SO-81-0451 SO-71-0089	LEE, IRVING S LEE, IRVING S LEE, IRVING S	FORD SWD FORD	10-07-66 08-07-84 06-02-71	10 10 10	179 200 187	2 2 2
SO	Be Be Be	85 86 87	SO-81-0216 SO-73-0897 SO-73-1872	BOWEN, CHARLES TWINING, CARL EAST, CRAIG	CD&P IWD LARSON DRLNG	03-28-83 04-27-77 04-20-81	20 10 10	43 182 83	2 4 12
SO	Be	88	SO-73-0945	W P HEARNE PRODUCE	WOOD	04-06-77	20	73	8
so	Be	89	SO-73-1180	W P HEARNE PRODUCE	WOOD	04-05-78	20	80	8
	Be	90	SO-73-1891	BOUNTIFUL RIDGE NURSERY	BUNDICK	06-01-81	10	55	4
	Be Be	91 92		NICHOLS, THOMAS	HALL		20		
	Be	93	SO-73-0743	JACOB, GEORGE SMITH, EARL	FORD FORD	00-00-78 07-03-76	10 10	160 190	2
SO	Be	94	SO-81-0827	POLLITT, ROBERT	SWD	06-20-86	15	85	2
	Be Be	95 96	SO-81-0746	POLLITT, ROBERT POLLITT, ROBERT	HALL HALL	10-30-85	20 10	225 210	2
SO	Ba	97	SO-81-0994	U.S. GEOLOGICAL SURVEY	USGS	12-11-86	2.0	15	2
	Be Be	98 99	SO-81-0993 SO-81-0996	U.S. GEOLOGICAL SURVEY U.S. GEOLOGICAL SURVEY	USGS USGS	12-12-86 12-12-86	15 20	15 15	2 2
	Be Be		SO-81-0997	U.S. GEOLOGICAL SURVEY	USGS	12-15-86	15	20	2
	Be		SO-81-0995 SO-81-0992	U.S. GEOLOGICAL SURVEY U.S. GEOLOGICAL SURVEY	USGS USGS	12-15-86 12-16-86	15 15	16 14	2
SO	Be Be	103	SO-81-0998 SO-81-0945	U.S. GEOLOGICAL SURVEY KING, J COMILLOUS	USGS DASHIELL DRLNG	12-16-86 08-05-86	15 10	22 190	2 2
	Ba Bf	110	SO-81-0470 SO-00-5594	SOMERSET CO. SANITARY COMM		09-19-84	20	280	4
	Bf	2	30-00-3394	PUSEY, R B POWELL, RUSSELL	CUSICK	04-12-50	30 25	232	2 2 4
	Bf Bf	3		DOODY, R T LONG, F	BEAUCHAMP	00-00-20 00-00-49	3.5 4.0	35.3	1.25 1.25
	Bf Bf	5		PUSEY, S F DYKES, R WADDY, W A	PUSEY		40	10	1.25
	Bf	6 7		WADDY, W A	DYKES BEAUCHAMP	00-00-50	35	46 30	1.25
	Bf Bf	8		WARWICK, L ORVIS, C M	BEAUCHAMP ORVIS	00-00-49	30	35 26	1.25
	Bf Bf	10 11	n	MILES, A	MILES	00-00-43	30	21.5	1.25
	Bf	12		JENKINS, W CANNON, MARGARET	JENKINS	00-00-52	3.5 2.0	36.5	1.25
	Bf	13	SO-73-1076	ST. MARKS UNITED METHODIST		09-13-77	30	220	2
50) Bf	14	SO-81-0135	WOOLFORD, STEPHEN	SWD	08-17-82	20	245	3
SO	Bf	15	SO-73-0885	SMULLEN, GRACE	IWD	01-28-77	40	250	4
	Bf Bf	16 17	SO-73-1527 SO-81-0595	BROWN, PAUL BROWN, PAUL	DASHIELL DRLNG SWD	07-10-79 05-03-85	40	240 230	4
so	Bf	18	SO-81~0207	ROPER, VIRGIL	SWD	02-28-83	20	200	3
SO	Bf	19	SO-73-1361	STEWART, WM E	WM BURNS	07-25-79	30	210	3
00	70.0	0.0	50 74 000°	DELIVER DE LA COMP					
	Bf Bf	20 21	SO-71-0005 SO-81-1016	REYNOLDS, ALBERT U.S. GEOLOGICAL SURVEY	IWD USGS	07-16-70 01-06-87	30 35	27	4
SO	Bf	22	SO-81-0962	ENNIS, CARROLL	SWD	01-06-87	30	12 230	2
	Bf	23	SO-81-0837	BOSTON, JERRY	SWD	04-18-86	20	240	2
20	Bf	24	SO-81-0862	BRENT, SCOTT	LARSON WELLS	05-08-86	25	178	2

c	ottom of esing or esed nter- val ft)	Diam- eter of screen	5	Bottom of screen or creened inter- val (ft)	Aquifer code	Weter level (ft)	Dete water level meas- ured	Drew- down (ft)	Dis- charge (8al/ min)	Pumpins period (hours)	Spe- cific capa- city [(8e1/ min)/ ft]	Use of weter	Loca well no.	
	33	15		63	122PCMK	6	04-28-77	29	800	3	28	I	SO Be	80
	140	2		165	122MNKN	12	04-21-81	4	12	1	3.0	T	SO Be	81
	155	2		179	122MNKN	14	10-07-66		10	2		Z	SO Be	82
	170 167	2		200 187	122MNKN 122MNKN	10 22	08-07-84	7	10 10	1 2	2.5	C	SO Be	83
	107			10/	122PINKN	22	06-02-71	/	10	2	1.4	C	SU De	04
	33	2		43	122PCMK	1	03-28-83	5	100	1	20	I	SO Be	8.
	172	2		182	122MNKN	23	04-27-77	27	20	4	.74	H	SO Be	86
	20	12		60	122PCMK	6	04-20-81	25	500	5	20	I	SO Be	8
	50			7.0	1000000	2.3m	04-16-87		4.0		0.0	-	00.8	
	53	8		73	122PCMK	3.9m	04-06-77 04-16-87	3 4.2	10 73	8	3.3	I	SO Be	81
	40	8		80	122PCMK	7	04-05-78	18	150	8	8.3	I	SO Be	8
				-	22210141	6.1m	04-16-87	10	200	0	0,0	_	00 00	
	3.5	4		55	122PCMK	6	06-01-81	36	25	6	. 69	I	SO Be	90
					122MNKN	3 . 8m	04-16-87					н	SO Be	9
					122MNKN							H	SO Be	9
	170	2		190	122MNKN	12	07-03-76	3	10	2	3.3	H	SO Be	9
	65	2		8.5	122PCMK	12	06-20-86	3	40	1	13.3	I	SO Be	9
					122MNKN							Z	SO Be	9
	190	2		170	122MNKN	29	10-30-85	2	40	1	20	Ü	SO Be	9
	200	_		210	12211111	30,3m	04-16-87		70	_	20		DO 20	
	10	2		15	112PCPC	8.3m	04-16-87					U	SO Be	9
	10	2		15	112PCPC	7.7m	04-16-87					U	SO Be	9
	10	2		15	112PCPC	4 . 8m	04-16-87					U	SO Be	9
	15	2		20	112PCPC	6.4m	04-16-87					U	SO Be	10
	11	2		16	112PCPC	8 . 4m	04-16-87					U	SO Be	
	9	2		14	112PCPC	4.8m	04-16-87					U	SO Be	
	17	2		22	112PCPC	8.0m	04-16-87					U	SO Be	
	170	2		190	122MNKN	25	08-05-86	3	20	1	6.6	H	SO Be	10
	180	4		230	122MNKN	28	09-19-84	7	60	3	8.6	U	SO Be	11
	217			232	122MNKN	14	04-12-50	9	25	5	2.8	H,S	SO Bf	
					112PCPC	2.3m	02-06-52					H,S	SO Bf	
					112PCPC 112PCPC	3.4m 1.3m	02-06-52 02-06-52					H,S	SO Bf	
					112FCFC	1.30	02-00-32					D	SO DI	
					112PCPC							H	SO Bf	
					112PCPC							H,S	SO Bf	
					112PCPC							H	SO Bf	
					112PCPC 112PCPC							H,S	SO Bf	
					TIZPCPC							а	SO DI	
					112PCPC	1.2m	02-08-52					H	SO Bf	1
					112PCPC	2.9m	02-08-52					H,S	SO Bf	1
					112PCPC							H	SO Bf	1
+1	200	2		220	122MNKN	17 20	09-13-77	7	10 70	2 2	1.4	T H	SO Bf	1
85		3	205	-245	122MNKN	25.0m	04-14-87	10	70	2	7.0	п	SO DI	Ţ
	240	2		250	122MNKN	30	01-28-77	30	20	4	. 67	H	SO Bf	1
	220	4		240	122MNKN	33.2m 22	04-14-87	7	20	1	2.9	H	SO Bf	1
	220	-		240	12241111	26.8m	04-14-87	,	20	1	2.5	41	OO DI	1
+1						25	05-03-85	5	30	2	6.0	I	SO Bf	1
20		3	200	-230	122MNKN	20	02-20 02	2	2.0	2	1.5	13	20 20	4
80		3	175	-200	122MNKN	21,6m	02-28-83	2	30	3	15	H	SO Bf	1
	190	2	2.0	210	122MNKN	3.5	07-25-79	115	100	2	.87	H	SO BE	1
						31,5m	04-14-87							
	20	4		30	1120000	12	07-10-70	10	20	4	0.0	т.	20 20	
	7	2		12	112PCPC 112PCPC	12 2.6	07-16-70 04-15-87	13	30	1	2.3	U	SO Bf	2
	205	2		230	122MNKN	30	08-27-86	3	60	1	20	H	SO BE	2
	215	2		240	122MNKN	22	04-18-86	1	65	i	65	S	SO BE	2
	168	2		178	122MNKN	20	05-08-86	4	15	2	3.8	Н	SO Bf	2

[ft = foot; in. = inch; gal/min = gallona per minute; (gal/min)/ft = gallons per minute per foot; -- = no data; m = measured; + = above land aurface; f = flowing well; SWD = Somerset Well Drilling; SP&W = Sydnor Pump and Well]

Loc wel no	1	Permit number	Owner	Contractor	Date well con- atructed	Alti- tude of land aur- face (ft)	Depth	Diam- eter of caaing (in,)
SO Bf SO Bg SO Bg SO Bg SO Bg	1 2 3	SO-81-0922 SO-73-1390 SO-81-0016	FORD, GREGORY ALDER, G BROWN, H GRUY FEDERAL INC PERRY HAWKIN CHRIST CHURCH	SWD BEAUCHAMP GLASCOCK SWD	07-10-86 00-00-48 00-00-52 11-14-78 11-06-81	43 30 25 30 20	235 14.8 1,050 275	2 1.5 4.5
SO Bg SO Ca	5	SO-81-0963 SO-00-8350	BEAUCHAMP, ROYCE BERWICK DEVELOPMENT CO	SWD CUSICK	09-03-86 10-15-51	25	260 871	2
SO Cb SO Cb SO Cb	2	SO-00-6230 SO-00-6338 SO-00-1893	BRITTINGHAM, BOYD THOMAS, WILLIAM C WEBSTER, JOHN W	CUSICK CUSICK CUSICK	07-08-50 07-26-50 11-25-47	5 5 5	157 142 142	1 1.5 1.5 1.5
SO Cb SO Cb SO Cb SO Cb SO Cb	5 6 7	SO-00-0905 SO-00-0586	JOHN BENNETT & PARTNERS HARRISON, STANFORD WILSON SEAFOOD CO BRITTINGHAM, BOYD DANIELS, MR	TODD CUSICK ROBBINS CUSICK	10-25-46 12-05-46 00-00-15 00-00-50 00-00-42	5 5 5 5	693 140 500-700(140 23	1.5 1.5 ?) 4 1.5 1.5
SO Cb SO Cb SO Cb SO Cb	10 11 12		WHITE, STANFORD WHITE, MR JONES, ROBERT S WALTERS, ADOLPHUS ABBOTT, OSCAR	CUSICK	00-00-45 00-00-47 00-00-49 00-00-52 00-00-43	5 5 10 5 5	23 23 19 7.5	1.5 1.5 1.25 1.5
SO Cb			BAKER, WALTER SOMERSET CO. BOARD OF ED	WHITE	00-00-45 09-00-53	5	8 147	1.5
SO Cb SO Cb	17	S0-73-0784	J. H. BURTON & SONS EDWARDS HENRY S FAITH SEAFOOD CO	OWNER FORD	00-00-34 05-30-80 07-30-76	5 6 5	313 21.4 150	3 2
SO Cb SO Cb SO Cb	20	SO-70-0040 SO-73-1326 SO-65-0070	ST. JAMES AME CHURCH SOMERSET CO. BOARD OF ED SUTTER, EVERETT C	FORD WALLACE MARSHALL	11-07-69 08-04-78 03-17-65	5 10 10	142 161 135	2 6 3
SO Cb		SO-68-0103 SO-68-0101	WEBSTER, ROY SOMERSET CO. BOARD OF ED	MARSHALL FORD	07-28-68 09-25-68	10 10	135 144	2 2 2
SO Cb SO Cb SO Cb SO Cb SO Cc	25 26 27	SO-81-0040 SO-81-0901 SO-81-0763 SO-81-0755 SO-00-0849	COLLIER, CHUCK HORNER, FRANK EDWARDS, HENRY WALTERS, CLYDE WHALEY, T B	HALL HALL HALL HALL WHITE	03-15-82 07-17-86 12-12-85 11-13-85 04-10-52	10 5 5 5 5	140 140 140 140 840	2 2 2 2 3 2 1.25
SO Cc SO Cc SO Cc	3	SO-68-0040	HOLLAND, AUBREY MEREDITH, CALVERT MRS U.S. GEOLOGICAL SURVEY RUMBLEY-FRENCHTOWN WATER CO	REVEL DEL GEO SURV SP&W	00-00-46 00-00-40 11-18-52 03-25-68	5 5 5	8 8 95 1,084	36 36 4
								4
SO Cc	6	SO-68-0041	RUMBLEY-FRENCHTOWN WATER CO	SP&W	05-27-68	5	1,090	4 8 6 6 6
50.0	7	20 70 0010	DIWEG DODING	VANADD				6
SO Cc	7	50-70-0043	DYKES, ROBERT E	KANARR	03-13-70	5	1,140	2
SO Cd SO Cd SO Cd	2	SO-00-6983 SO-00-8346 SO-00-1535	BLOODSWORTH, BEAUCHAMP KAUFFMAN, CAPT F.B. CHRISTENSEN, KOREN	WHITE WHITE FARLOW	11-00-50 09-00-51 06-14-47	5 5 5	151 166 150	2 2 2 2
SO Cd	4	SO-00-8351	BAUGHER, MARGARET	CUSICK	08-01-51	10	193	1.5

ca	of of or or eacd iter val	8	Diam- eter of screen (in.)	scr in	ttom of raen or eaned ter- al	Aquifar code	Watar level (ft)	Date watar level meaa- ured	Draw- down (ft)	Dis- charge (gal/ min)	Pumping period (hours)	Spe- cific capa- city [(8a1/ min)/ ft]	Uaa of water	Local well no.	
	1	210 ,039 245	2		235 275	122MNKN 112PCPC 112PCPC 124PNPN 122MNKN	28 2.6m 10	07-10-86 02-08-52 11-06-81	2 2	60 10	1 1	30 5.0	H H H,S U	SO Bf SO Bg SO Bg SO Bg SO Bg	25 1 2 3 4
0	-	230 540	2		260	122MNKN	16	09-03-86	_1	60 10	1	60	H	SO Bg SO Ca	5
540	-	865 137 122 132			157 142 142	217 PTMC 122MNKN 122MNKN 122MNKN	3 2 4	07-08-50 07-26-50 11-25-47	8 2 3	8 15 30	4 4 1	1.0 7.5	H H C,H	SO Cb SO Cb SO Cb	1 2 3
		693 120 120			140	124PNPN 122MNKN 124PNPN 122MNKN 112PCPC	8 4	10-25-46 12-05-46 	2 	10 25 	8 2 	3	C,H C,H C H	SO Cb SO Cb SO Cb SO Cb	5 6 7 8
						112PCPC 112PCPC 112PCPC 112PCPC 122MNKN	==						H H H,S H	SO Cb SO Cb SO Cb SO Cb	9 10 11 12 13
0	-	42				112PCPC	3	09-00-53		12			H T	SO Cb	14 15
0	Ī	133 18 130	3 2	133	- 147 21.4 150	122MNKN 122CPNK 112PCPC 122MNKN	5.1m 6 4.5m	05-30-80 07-30-76 04-16-87	1	15	2	 15	C	SO Cb SO Cb	16 17 18
+1	_	130 117 42	6		142 138	122MNKN 122MNKN	6 8	11-07-69 08-04-78	3 16	10 20 18	2 3 2	3.3	T N C	SO Cb SO Cb	19 20 21
+1	-	129 129 124	2 2 2	129	- 135 135 144	122MNKN 122MNKN 122MNKN	12 8	07-28-68 09-25-68	10	9 10	1 2	.90	C	SO Cb	22
+2	- 5-	120 120 120 120 107 588	2 2 2 2		140 140 140 140	122MNKN 122MNKN 122MNKN 122MNKN	10 7 8 5 +5	03-15-82 07-17-86 12-12-85 11-13-85 04-10-52	3 3 1 	9 50 40 4	1.5	3.0 17 40	Н Н Н Н	SO Cb SO Cb SO Cb SO Cc	24 25 26 27
510	-	720	1.25	720	- 740	124PNPN									
+1.	6-	8 8 202				112PCPC 112PCPC 122PCMK	 +10.5m	05-03-68		100	 12		H H U P	SO Cc SO Cc SO Cc	3 4 5
202 1,043 1,063 1,078	-1 -1	,038 ,058 ,068 ,084	4 4	1,038 1,058 1,068	-1,043 -1,063 -1,078	217PTMC 217PTMC 217PTMC									
+1. 209 1,026 1,057 1,078	5- -1 -1 -1	209 ,021 ,052 ,073 ,083	6 6 6	1,021 1,052 1,073	-1,028 -1,057 -1,078	217PTMC 217PTMC 217PTMC	+1.5f	05-27-68		200	12	~=	P	SO Cc	6
+2 1,055 1,128	-1	,035 ,114 ,140	4 2 2	1,035 1,045 1,114	-1,050 -1,055 -1,128	217PTMC 217PTMC 217PTMC	+8 3.2m	03-09-70 04-15-87	138	30	2	. 22	С	SO Cc	7
148.		141 156 140.	2 2 5	140.	151 166 5- 148,5	122MNKN 122MNKN 122MNKN	+1 1.5 +1	11-00-50 09-00-51 06-14-47		12 12 30	6 2		H,S H,S	SO Cd SO Cd SO Cd	1 2 3
140.	5 -	150 183			193	122MNKN	0	08-01-51	1	30	2	30	H,S	SO Cd	4

[ft = foot; in. = inch; gsl/min = gsllons per minuts; (gsl/min)/ft = gsllons per minute per foot; -- = no data; m = measured; + = sbove lsnd surfsce; f = flowing well; SWD = Somersst Well Drilling]

Loca well no.		Permit number	Owner	Contractor	Date well con- structed	fscs	Depth	
O Cd	5		WALSTON, ARZIE			5	21	1.25
30 Cd	6		WILSON, EVELYN		00-00-40	5	8	36
SO Cd	8		STEVENSON, ETHEL MRS SLAGLE, ELMER F		00-00-27		30 20	1.5
SO Cd	9		SOMERSET CO. BOARD OF ED		00-00-36	_	350	2
0 Cd	10		ROSS, E P		1875	5	25	36
SO Cd	11	~ -	FAIRMOUNT PARSONAGE		00-00-02			36
O Cd	12		DAVIS, ELWOOD	DAVIS WATERS	00-00-50		60 25	1.5
O Cd	14		WATERS, CARRIE MCLEAN, WILLIAM	AVIEV2	00-00-13		18	1.5
0 Cd	15		MCLEAN, WILLIAM		00-00-12	5	20	24
	16		WARWICK, JAMES		1890	5	169	2
O Cd	17		CARPENTER, EDWARD		00-00-39		92	2
O Cd	18		CARPENTER, EDWARD CARPENTER, EDWARD	ENNIS BROS	10-00-44	20	15.7 196	36 6
SO Cd	20							4
				ENNIS BROS	03-25-41	20	200	2.5
00 Cd	21		LONG, M T		00-00-12		20	1.25
O Cd			REICHARD, CHAS REICHARD, CHAS		00-00-35		15 15	1.75
	24		BEECHUM, ROBT		10-00-12		20	1.5
0 Cd	25		JONES. ROY		00-00-37	5	21	1.25
o cd	26		GROVER, WM M		00-00-47		15	2
SO Cd			GROVER, WM M		00-00-47		15	2
80 Cd	28		GROVER, WM M FONTAINE, CHAS MRS	CUSICK	00-00-47	20 10	15 190	1.5
o Cd	30		FONTAINE, W W	FONTAINE	00-00-40		15	1.5
SO Cd	31		JOYNES, J P	JOYNES	00-00-40		15	1.25
O Cd	32			JOYNES	00-00-49	5	19	1.25
0 Cq	33		GREEN, SAMUEL ROBINSON, MARGARET	GREEN	00-00-42		20 15	1.5
0 Cd	35		MADDOX, RANDOLPH		04-17-52		4.5	24
SO Cd	36		BOARD, CHESTER P		1877	10	7.6	24
50 Cd	37 38		RUARK, GARLAND	RUARK	00-00-50		90	1.5
30 Cd	39		MAALOE, F W HAYMAN, E G	MAALOE WHEATLEY	00-00-27		23 160	1.5
30 Cd		SO-73-1425 SO-70-0052	SOMERSET CO. SANITARY COMM BOZMAN, HAROLD	DELMARVA DRLNG FORD	03-21-79 12-11-69		1,145 53	8
SO Cd	43	SO-73-0661	BOZMAN, HAROLD	FORD	04-01-76	10	5.5	2
	44	SO-81-0252	SAMUEL WESLEY UM CH	SWD	06-01-83		50	2
SO Cd	45	SO-73-1738	WHITE, TOM	KAUFFMAN	07-16-80	5	70	4
SO Cd	46	SO-73-1742	WHITE, TOM	KAUFFMAN	07-18-80	5	155	4
SO Cd	47	SO-73-1737	WHITE, TOM	KAUFFMAN	07-16-80		70	4
SO Cd	48	SO-73-0563 SO-81-0800	BLOODSWORTH, DOUGLAS	DASHIELL	09-22-75	5	160	2
			MONICK, STEPHEN	SWD	03-07-86		165	4
SO Cd	50	SO-81-0086	AINSWORTH, FREEDOM H	SWD	07-03-82	5	162	3
SO Cd	51	SO-73-0529	CARPENTER, MURTON D	FORD	02-05-76	20	200	4
SO Cd	52		FORD, CECIL				57	
SO Cd	53	SO-81-0840 SO-81-0824	WALLER, DAVID LEE SCHERBACK, JOHN	DASHIELL DRLNG HALL	04-18-86 06-14-86	5	180 160	2
SO Ce			DUNCAN, C K	WHITE	05-02-46			3
								2
								2
O Ce	2	SO-00-3479	DORSEY, THOMAS	CUSICK	12-31-48		246	
O Ce	3	SO-00-4145 SO-00-8871		CUSICK CUSICK	06-28-48 10-31-51		228 225	1.5
O Cs	5	SO-00-5741	RICHARDS J R	CUSICK	05-11-50		240	1.5
	- 40	20 00 0011	MASSEY, HAROLD E	distriction of the second	11-20-46		90	

Bott				Bottom of							Spa-			
cas: o: css: into vs: (ft	ing r ad er- 1	Diam- ater of screen (in.)		screen or creened inter- val (ft)	Aquifer coda	Water level (ft)	Date water level meaa- ured	Draw- down (ft)	Dia- charge (gal/ min)	Pumping pariod (hours)	cific caps- city	Use of watar	Loca well no.	
					112PCPC							H	SO Cd	
					112PCPC							H	SO Cd	- 1
					112PCPC 112PCPC							H H,S	SO Cd	
					122CPNK	+1f	06-00-44					U	SO Cd	
					*******							Н	SO Cd	1
					112PCPC 112PCPC							H	SO Cd	1
					122PCMK							H	SO Cd	1.
					112PCPC							H	SO Cd	1
					112PCPC							п	SU Ca	1
					112PCPC							H	SO Cd	1
					122MNKN							H	SO Cd	1
					122PCMK 112PCPC	7.7m	02-06-52					H,S	SO Cd	1
	170	4		196	122MNKN	18	10-00-44					H,S	SO Cd	1
) -	170											H,S	so cd	2
0 -	183		183	- 193	122MNKN							n,s	SO Cu	-
					112PCPC							H,S	so Cd	2
					112PCPC							H,S	SO Cd	1
					112PCPC 112PCPC							S H,S	SO Cd	1
					1121010									
					112PCPC	. 5	02-16-52					H,S	SO Cd	4
					112PCPC 112PCPC							H,S	SO Cd	1
					112PCPC							H,S	SO Cd	1
					122MNKN							H	SO Cd	-
					112PCPC	5	04-17-52					s	SO Cd	3
					112PCPC		04-17-32					H	SO Cd	3
	15			19	112PCPC	0	00-00-49					H	SO Cd	3
					112PCPC 112PCPC							H	SO Cd	200
					1125050							2.1	SU Ca	
					112PCPC	1.9m	04-17-52					H	SO Cd	3
					112PCPC 122PCMK	3.7m	04-17-52					H	SO Cd	
					112PCPC							H	SO Cd	
					122MNKN	*						H	SO Cd	
1	1,110	6		1,140	217 PTMC	+7.8	05-28-79	99.4	100	26	1.0	P	SO Cd	
) -	21	J		1,140	21/FIRC		03-26-79	39.4	15	2		ć	SO Cd	
2 -	45	2	45	- 53	122PCMK		01 61 -				4.5			
	40	2		55 50	122PCMK 122PCMK	5 7	04-01-76 06-01-83	1 3	10 15	2	10 5.0	C	SO Cd	
	3.5	4		65	122PCMK	4.5	07-16-80	7	60	2.5	8.6	Ī	SO Cd	
						2.0m	04-15-87							
	100	4		155	122MNKN	12.2	07-18-80	5.8	3 40	3	6.9	I	SO Cd	
	37	4		67	122PCMK	9	07-16-80	9	60	2.5	6.7	Ī	SO Cd	
	152	2		160	122MNKN	20	09-22-75	7	15	1	2.1	C,H	SO Cd	
	145	4		165	122MNKN	10	03-07-86	1	50	1	50	H	SO Cd	
1 -	- 87					9.0m 7	04-14-87	3	40	1	13	N	so cd	
7 -	142	3	142	- 162	122MNKN	9.2m	04-14-87	9	70	_	10	44	50 04	
	100			0.00					0.7				60 6:	
	180	4		200	122MNKN	9 19.7m	02-05-76 04-14-87	18	25	2	1.4	I	SO Cd	
					122PCMK							-	SO Cd	
	170	2		180	122MNKN	16	04-18-86	7	16	1		H	SO Cd	
0 -	140	2		160	122MNKN	4	06-14-86 05-02-46	_1	47 40	1 2	47	H	SO Cd	
0 -	68.		68	.5- 76.	5 122PCMK	4	VJ VZ-40		40	4			55 00	
6.5-	78													
	228	1.5		246	122MNKN	2	12-31-48	h	22	2	5 5	н	SO Ce	
	213	1.5		228			06-28-48	4	15	3	3.8	H,S		
	210	1.5		225	122MNKN	3	10-31-51	1.5	30 24 30	2	20	H	SO Ce	
	225	1.5		240		2	05-11-50	2	24	2	12	H	SO Ce	
	78	2		90	122PCMK	3	11-20-46	1	30	3	30	5	SO Ce	

[ft = foot; in. = inch; gel/min = gallone per minute; (gel/min)/ft = gallons per minute per foot; -- = no date; m = measured; + = above land eurface; f = flowing; SAW = Shannahan Art. Well; SWD = Somereet Well Drilling}

Local well no.	Permit number	Owner	Contractor	Date well con- structed	Alti- tude of lend sur- face (ft)	Depth drilled (ft)	Diam- eter of casing (in.)
SO Ce 7 SO Ce 8 SO Ce 9 SO Ce 10 SO Ce 11	SO-00-3435 SO-00-1020 SO-00-1019 SO-00-4272 SO-00-3618	WIDDOWSON, N D CHAMBERLIN, JOHN A JAMES, WM T, JR RUARK, DONALD M MCDORMAN, ROBERT H	WHITE CUSICK CUSICK CUSICK WHITE	12-21-48 12-24-46 12-28-46 07-21-49 09-00-49	10 15 15 15 10	198.5 235 238 222 192	2 1.5 1.5 1.5
SO Ce 12 SO Ce 13 SO Ce 14 SO Ce 15 SO Ce 16	SO-00-2002 SO-00-2433 SO-00-1105 SO-00-1102	BLEVINS, SOMERS LONG, DENET LONG BROS RING, ROY J JOYNES, J F	CUSICK CUSICK CUSICK CUSICK WHITE	11-19-47 04-21-48 08-22-47 09-27-47 00-00-47	10 15 15 10	210 237 233 190 190	1.5 1.5 2.5 1.5
SO Ce 17 SO Ce 18 SO Ce 19 SO Ce 20 SO Ce 21		RICHARDSON, PHILLIP MULCAHY, D J WHITE, WILLIAM MRS SUMMER LABOR CAMP	 KOHL BROS	00-00-47 00-00-51 00-00-50 	10 15 15 5	18 20 40 	1.5 1.5 1.5
SO Ce 22 SO Ce 23 SO Ce 24 SO Ce 25 SO Ce 26		THOMPSON, FRANCES BRUIN, ELLA COLLINS, DAVID WHITE, ISAAC	TAYLOR	00-00-51 00-00-49 00-00-42 00-00-46	15 15 15 15 15	100 25 19.7 18 20	1.5 1.5 1.5 1.25 1.25
SO Ce 27 SO Ce 28 SO Ce 29 SO Ce 30 SO Ce 31		SIGRIST, JOE DORSEY, WALTER MCLENDON, COL E L RUE, WILLIAM		08-00-48 00-00-47 1795	20 20 5 5	25 25 39.1 9.4 24	1.25 1.5 1.25 24 1.25
SO Ce 32 SO Ce 33 SO Ce 34 SO Ce 35 SO Ce 36		WILLIAMS, GEO FORD, WILLIAM POOLE, CHARLIE MRS HUFFMAN, D F	WILLIAMS BUFFMAN	00-00-51 00-00-42 00-00-40 00-00-46	10 15 15 15 15	30 24 41.2 11.1 32	1.25 1.25 1.25 1.25 1.25
SO Ce 37 SO Ce 38 SO Ce 39 SO Ce 41	SO-01-0638 SO-81-0390	PERRY, W W KEENAN, HARRY BARNES, G HOWETH EASTERN CORRECTIONAL INST	CUSTIS SCOTT CUSICK SAW	00-00-17 00-00-50 08-27-52 06-11-84	5 5 20 10	190 226.5 232 254	3 2 1.5
SO Ce 42	SO-81-0394	EASTERN CORRECTIONAL INST	SAW	06-15-84	15	215	4
SO Ce 43	SO-81-0393	EASTERN CORRECTIONAL INST		06-19-84	20	246	2
SO Ce 44	SO-81-0556 SO-81-0389	EASTERN CORRECTIONAL INST		03-22-85	10	250	10
SO Ce 46	SO-81-0391	EASTERN CORRECTIONAL INST EASTERN CORRECTIONAL INST		06-12-84 06-13-84	10	227	2
SO Ce 47	SO-81-0435	EASTERN CORRECTIONAL INST		06-28-84	10	229	2
SO Ce 48	SO-81-0557	EASTERN CORRECTIONAL INST	LAYNE ATLNTC	04-09-85	10	235	10
SO Ce 49	SO-81-0433	EASTERN CORRECTIONAL INST	SAW	07-03-84	10	244	6
SO Ce 50	SO-81-0454	EASTERN CORRECTIONAL INST	SAW	07-05-84	10	222	2
SO Ce 51	SO-81-0434	EASTERN CORRECTIONAL INST	SAW	06-29-84	10	240	2
SO Ce 52	SO-81-0392	EASTERN CORRECTIONAL INST	SAW	06-15-84	10	205	2
SO Ce 53 SO Ce 54 SO Ce 55 SO Ce 56 SO Ce 57	SO-73-1252 SO-73-0633 SO-73-0490 SO-81-0733 SO-81-0120	SOMERSET CO. TOURIST COMMUS. POSTAL SERVICE ENGLISH'S GRILL SOMERSET WELL DRLG. CO. KATO INC	SWD FORD DASHIELL DRLNG SWD SWD	06-30-78 04-28-76 05-29-75 12-20-85 07-30-82	15 20 20 10 20	220 50 230 230 240	2 2 2 4 2

Bottom of cesing or cesed inter- val (ft)	Diem- eter of ecreen (in,)	Bottom of ecreen or screened inter- vel (ft)	Aquifer code	Weter lavel (ft)	Dete weter level mees- urad	Drew- down (ft)	Die- cherge (gal/ min)	Pumping pariod (hours)		Uee of water	Loce well no.	
182. 220 223	1.5 1.5	198.5 235 238	122MNKN 122MNKN 122MNKN	4 4 4	12-21-48 12-24-46 12-28-46	1.5	15 30 25	3 3 3	20 17	H,S H H	SO Ce SO Ce	7 8 9
207 179	1.5	222 191	122MNKN 122MNKN	1.5	07-21-49 09-00-49	4.5	24 15	3	5.3	H,S	SO Ce	10
195 222	1.5	210	122MNKN	1	11-19-47	3	20	2	6.7	H,S	SO Ce	12
215	1.5 2.5	237 233	122MNKN 122MNKN	3	04-21-48 08-22-47	3	30 50	16 3	7.5	H N	SO Ce	13
175	1.5	190	122MNKN	+.7	09-27-47	2.7	28	2	10	H,S	SO Ce	15
***			122MNKN							H,S	SO Ce	16
			112PCPC							Н	SO Ce	17
			112PCPC 112PCPC							H,S	SO Ce	18
			112PCPC				12			P	SO Ca	20
			122MNKN							P	SO Ce	2:
	40. 40.		122PCMK							C,H	SO Ca	22
			112PCPC		0/ 1/ 50					H	SO Ce	23
			112PCPC 112PCPC	4.1m	04-14-52					H,S	SO Ca	24
			112PCPC							Н	SO Ce	26
22		25	112PCPC				4			H	SO Ce	27
			112PCPC							H,S	SO Ce	28
			122PCMK 112PCPC	2.3m 4.9m	04-18-52					U	SO Ce	29
			112PCPC	4,90	04-18-32					H,S	SO Ce	31
			112PCPC							Н	SO Ce	32
			112PCPC							H.S	SO Ce	33
			122PCMK	4.6m	04-18-52					H,S	SO Ce	34
			112PCPC 122PCMK	2.1m	04-18-52					U	SO Ce	35
			122MNKN	f			4.5			H.S	SO Ca	37
214.		226.5	122MNKN	+.5	00-00-50		4,5			H,S	SO Ce	38
217		232	122MNKN	6	08-29-52	4	25	2	6.3	H	SO Ce	3
196	4	246	122MNKN	24 53.4m	06-11-84 04-14-87	38	70	41	1.8	U	SO Ce	40
185	4	215	122MNKN	21 24.5m	06-15-84 04-14-87	142	45	29	.32	U	SO Ce	42
211	2	226	122MNKN	25 24.3m	06-19-84 04-14-87	15	17	1	1.1	บ	SO Ce	43
190	10	240	122MNKN	24	03-22-85	119	125	48	1.1	T	SO Ce	41
203	2	218	122MNKN	23.8	07-01-85 06-12-84	144	180 15	48	1.3	U	SO Ce	4:
				31.1m	04-14-87							
203	2	218	122MNKN	26 37.0m	06-13-84 04-14-87	8	17	1	2.1	U	SO Ce	46
204	2	219	122MNKN	25 30,6m	06-28-84 04-14-87	26	19	1	. 73	U	SO Ce	47
190	10	230	122MNKN	22	04-09-85 05-28-85	124 146	180 125	48	1.5	T	SO Ce	48
197	6	228	122MNKN	24	07-03-84	155	152	72	.98	U	SO Ce	45
207	2	222	122MNKN	34.1m 22	04-14-87 07-05-84	25	20	1	. 80	U	SO Ce	50
				32.0m	04-14-87	23		1			20 Ce	30
206	2	221	122MNKN	25 31.2m	06-29-84 04-14-87	25	20	1	.80	U	SO Ce	51
190	2	205	122MNKN	24 25.6m	06-15-84 04-14-87	41	19	1	.46	ប	SO Ce	52
200	2	220	122MNKN	8	06-30-78	2	10	1	5.0	C	SO Ce	53
40	2	50	122PCMK	5	04-28-76	1	10	3	5.0	C	SO Ce	54
220	2	230	122MNKN	21	05-29-75	7	15	1	2.1	C	SO Ce	55
140	3	230	122MNKN	20	12-20-85	3	65	5	22	C	SO Ce	56
210	2	240	122MNKN	12	07-30-82	2	10	1	5.0	C	SO Ce	5

[ft = foot; in. = inch; gal/min = gallons per minute; (gal/min)/ft = gallons per minute per foot; -- = no data; m = measured; + = above land aurface; SWD = Somerset Well Drilling; IWD = Ideal Well Drillere, USGS = U.S. Geological Survay; CD&P = Coaatal Drilling and Pump]

Local well no.	Permit number	Owner	Contractor	Data well con- atructad	Alti- tude of land aur- face (ft)	Depth drilled (ft)	Diam- etar of cesing (in,)
SO Ce 58 SO Ce 59 SO Ce 60	SO-81-0166 SO-73-1709 SO-73-1079	SOMERSET WELL DRILLING CO RICHARD'S EXXON SERVICE FORD, MARION	SWD SWD FORD	11-10-82 06-30-80 10-01-77	10 20 10	240 240 100	2 2 2
SO Ce 61 SO Ce 62	SO-73-0091 SO-73-0742	MD STATE HIGHWAY ADMIN. FORD, MARION	KELLEY FORD	05-09-73 09-13-76	20 10	63 150	6 2
SO Ce 63 SO Ce 64	SO-73-0117 SO-73-0639	LONG BROS. SOMERSET CO. HEALTH CENTER	FORD IWD	05-11-73 03-20-76	10 10	230 128	2
SO Ce 65	SO-73-0796	LONG BROS.	FORD	08-02-76	10	240	2
SO Ce 66 SO Ce 67	SO-73-0725 SO-65-0076	CRUSADE EVANGEL CHURCH LONG, EDWIN D	HALL FORD	05-18-76 03-23-65	20 10	20 231	3
SO Ce 68 SO Ce 69	SO-73-0402 SO-81-0159	MAPLE GRILL DELMARVA POWER & LIGHT	FORD SWD	12-27-74 10-14-82	20 10	230 240	2
SO Ce 70 SO Ce 71 SO Ce 72	SO-73-0638 SO-69-0039 SO-71-0094	J MILLARD TAWES VOCAT'L SCH DUNCAN, CLINTON K ST. JAMES METHODIST CHURCH	IWD IWD FORD	03-25-76 10-14-68 05-07-71	10 20 10	129 230 220	3 6 4 2
SO Ce 73 SO Ce 74 SO Ce 75 SO Ce 76	S0-67-0102 S0-73-0303 S0-81-0092	DUNCAN, CLINTON K LAKE SOMERSET CAMP WISE, THOMAS JEHOVAH'S WITNESSES	FORD FORD SWD	03-15-67 09-04-74 07-01-82	20 20 20 20	236 250 240	2 2 2 2
SO Ce 77 SO Ce 78 SO Ce 79 SO Ce 80 SO Ce 81	SO-81-0374 SO-81-0555 SO-73-1896 SO-81-0498 SO-81-0036	SCHROCK, RANDY MILFORD FERTILIZER HONEST BOB'S SMALL MALL ST. ELIZABETH CATHOLIC CH DARWENT, BASIL DR	DELMARVA DRLNG SWD SWD SWD CD&P	03-23-84 01-24-85 04-23-81 09-12-84 03-30-82	10 20 10 10	230 240 230 80	2 2 2 2 4
SO Ce 82	SO-73-0181	SOMERSET CO. BOARD OF ED	BUNDICK	03-13-73	10	230	4
SO Ce 83 SO Ce 84	SO-73-1515 SO-81-0597	GEORGE W. ENNIS' SONS WIDDOWSON FARMS	IWD SWD	06-14-79 04-22-85	20 10	30 210	2 4
SO Ce 85 SO Ce 86 SO Ce 87	SO-81-0625 SO-81-0624 SO-81-1014	PERDUE, INC PERDUE, INC. U.S. GEOLOGICAL SURVEY	SYDNOR HYDRO SYDNOR HYDRO USGS	08-16-85 08-16-85 12-10-86	20 20 15	240 260 13	3 6 6 2
SO Ce 88 SO Ce 90 SO Ce 92 SO Ce 94	SO-81-1015 SO-81-0679 SO-73-1894 SO-81-0977	U.S. GEOLOGICAL SURVEY PUGH, WILLIE NELSON, FREDERICK HALL, MICHAEL	USGS SWD SWD SWD	01-05-87 08-14-85 04-22-81 10-02-86	5 13 10 8	12 230 190 183	2 2 2 4
SO Cf 1	SO-00-2038	GREEN, FULTON	WHITE	10-31-47	15	210	2 3 2 2
SO Cf 2 SO Cf 3 SO Cf 4	SO-00-7353 SO-00-2039	MD STATE HIGHWAY ADMIN. PAYNE, MAURICE BENSON, GEORGE	USGS WHITE WHITE	08-16-49 03-00-51 11-29-47	20 10 15	15 216 243	1.25 2 3 2
50 Cf 5	SO-00-1272	PUSEY, W F WEIDEMA, WM	CUSICK	00-00-45 06-24-47	25 20	39 256	2 1.25 1.5
SO Cf 7 SO Cf 8 SO Cf 9 SO Cf 10 SO Cf 11		LONG, W H MD GAME & INLAND FISH COMM COTMAN, LAURA MRS. BEAUCHAMP, T B		00-00-42 00-00-12 00-00-37 1852	20 20 20 20 20	19 15.1 16 23 21.2	36 1.5 1.25 1.25
SO Cf 12 SO Cf 13 SO Cf 14 SO Cf 15 SO Cf 16	SO-73-1291 SO-81-0276 SO-73-1635	TAYLOR, M V CREASY, FRED WOLFRAM PAUL DEPT NAT RES WILDLIFE ADMIN OLSEN'S FURNITURE	SWD DRLNG SWD	00-00-49 00-00-40 07-29-78 07-28-83 12-28-79	20 20 22 20 20	22.5 35 90 244 240	1.25 1.5 2 2

B	of			ttom							Spe-			
c	aeing or cased inter- vel (ft)	Diam- eter of ecreen (in,)	ecr	reen or eened ter- al ft)	Aquifer code	Water level (ft)	Date water level meae- ured	Draw- down (ft)	Die- cherge (gel/ min)	Pumping period (hours)	cific capa- city	Use of water	Loca well no.	
	200	2		240	122MNKN	18	11-10-82	4	30	2	7.5	С	SO Ce	5
	210 70	2		240 100	122MNKN 122PCMK	10 7 5.0m	06-30-80 10-01-77 11-24-86	3	10 10	3	3.3	C	SO Ce	5
	53 125	6 2		58 150	122PCMK 122MNKN	2 8	05-09-73 09-13-76	10 2	60 12	3 2	6.0	Z C	SO Ce	6
	214 121	2		230 127	122MNKN 122MOCN	10	05-11-73 03-20-76	91	10 10	3 24	2.5	C	SO Ce	6
	210	2		240	122MNKN	16.9m	04-16-87 08-02-76	2	20	2	10	C	SO Ce	6
	16			20	1100000	22.6m	04-15-87					_		
	16 84	2		20 231	112PCPC 122MNKN	6	05-18-76 03-23-65	3	14	1 2	1.3	C	SO Ce	6
	210 105	2		230	122MNKN	12 18	12-27-74 10-14-82	2 2	10 30	2 2	5.0 15	C	SO Ce	6
	210 118	3 6		240 127	122MNKN 122PCMK	4	03-25-76	91	20	24	22		SO Ce	7
	214	3		230	122PCMK 122MNKN	18	10-14-68		20	6	.22	T	SO Ce	7
	204			220	122MNKN	10	05-07-71	2	10	3	5.0	T	SO Ce	7
	216 220	2 2		236 250	122MNKN 122MNKN	16 10	03-15-67	2	10 10	2	5.0	C	SO Ce	7
					112PCPC		09-04-74					C	SO Ce	7
	220 30	2		240 75	122MNKN 122PCMK	4.5	07-01-82 03-23-84	3	10 50	1.5	3.3	I	SO Ce	1
	205	2		230	122MNKN	11	01-24-85	4	11	1	2.8	С	SO Ce	
	220 205	2 2		240 230	122MNKN 122MNKN	12 10	04-23-81	4	10 10	1	2.5	C	SO Ce	1
	40	4		80	122 PCMK	2	03-30-82	5	200	î	40	Î	SO Ce	1
	200	4		230	122MNKN	0.5m 18	04-16-87 03-13-73	57	25	6	. 44	T	SO Ce	
+1	25 1 -120	2		30	112PCPC	12 18	06-14-79 04-22-85	3 7	80 50	1 3	27	CS	SO Ce	
120	0 -180	3	180	-210	122MNKN									
	198 198	6		228 228	122MNKN 122MNKN	26	08-16-85 08-16-85	102	80 80	2	.78	U	SO Ce	1
	8	2		13	112PCPC	28 2.4m	04-16-87	102		2	.78	S	SO Ce	8
	7 210	2 2		12	112PCPC	4.7m	11-24-86					U	SO Ce	8
	170	2		230 190	122MNKN 122MNKN	5	08-14-85	5	22 15	1	5.0	H	SO Ce	-
+1	-158	2	158	-183	122MNKN	12	10-02-86	1	60	1	60	H	SO Ce	(
0		<i>5</i> 2	100	5 000 5	100,000	0	10-31-47		18	3		H	SO Cf	
	0 -196. 8.5-210	5 2	196.	5-208.5	122MNKN									
	12	1		15	112PCPC	1.5m	04-14-87					υ	SO Cf	
(203	2		215	122MNKN	4.5	03-00-51 11-29-47		10 18	6 2.5		H,S H,S	SO Cf	
- (5 2	225.	5-237.5	122MNKN	4.5	11-29-47		10	2.3		п,ъ	50 CI	
	241	1.5		256	112PCPC		06-24-47				10	H	SO Cf	
	241	1.5		256	122MNKN 112PCPC	6	06-24-47	2	20	3	10	H	SO Cf	
					112PCPC	3.1m	03-27-52					H,S	SO Cf	
					112PCPC							H,S	SO Cf	
					112PCPC 112PCPC	1.8m	03-27-52					H,S U	SO CI	
					112PCPC							H	SO Cf	1
	75	2		90	112PCPC 122PCMK	12	07-29-78	3	8	1	2.7	H	SO Cf	1
	224	2		244	122PCMK 122MNKN	26	07-29-78	14	15	1	1.1	Z	SO Cf	1
	220	2		240	122MNKN	12	12-28-79	3	10	2	3.3	C	SO Cf	1

[ft = foot; in. = inch; gal/min = gallona per minute; (gal/min)/ft = gallona par minuta per foot;
-- = no data; m = measured; + = above land aurface; f = flowing well; SWD = Somarset Well
Drilling; IWD = Ideal Well Drillera; SAW = Shannahan Art. Well; USGS = U.S. Geological Survey)

Loca well no.		Permit number	Owner	Contractor	Date well con- atructed	Alti- tude of land sur- face (ft)	Depth drilled (ft)	Diam- atar of caaing (in.)
SO Cf SO Cf SO Cf	18 19 20	SO-66-0080 SO-81-1017 SO-73-1985	GREENHILL CH OF THE BRETH U.S. GEOLOGICAL SURVEY MICHELS, WALTER	FORD USGS SWD	05-19-66 01-07-87 11-08-81	20 20 20	234 12 35	2 2 2
SO Cf SO Cf	21 22	SO-81-0925 SO-73-1979	MATTHEWS, HELEN UN. CH. OF JESUS CHRIST	DASHIELL BLAKE & CO	07-11-86	15 20	93	2
SO Cg SO Cg SO Cg SO Cg SO Cg	1 2 3 4 5	SO-81-0851	BEAUCHAMP, R POPE, J MRS SOMERSET CO. BOARD OF ED OWEN, MELVIN THOMPSON, WILLIAM	BEAUCHAMP SWD	00-00-49 00-00-48 00-00-48 04-23-86	20 25 10 20 5	24 19 11.8 14.5 220	1.25 1.25 1.25 1.25 2
SO Cg SO Db	6	SO-65-0030 SO-04-7688	CHESAPEAKE EGG CO U.S. FISH & WILDLIFE SERV	SAW AARON	09-04-64 07-24-62		115 860	2.5
SO Dc SO Dc SO Dc	1 2 3	SO-01-0513 SO-67-0007	U.S. GEOLOGICAL SURVEY WARD, NOAH U.S. GEOLOGICAL SURVEY	DEL GEO SURV CUSICK SYDNOR HYDRO	11-17-52 07-30-52 11-11-66	5	64 139 1,514	1.5 4 1.5 6 4
								2 2
SO Dc	4	SO-71-0064	CITY OF CRISFIELD	DELMARVA DRLNG	12-05-70	5	1,200	8
SO Dc	5	SO-73-1722	BRAMBLEWOOD ASSOCIATION	IWD	08-22-80	5	122	6
SO Dc SO Dd	6	SO-00-0162	KATO-CROCKETT SHIPYARD BLAKE, ALAN F	 WHITE	04-18-46	5	201	2
SO Dd	2	SO-00-2036	GALE, G	CUSICK	11-06-47	10	170	1.5
SO Dd SO Dd SO Dd SO Dd SO Dd	3 4 5 6 7	SO-00-5169 SO-00-6589 SO-00-7607 SO-00-8342 SO-00-8133	DAUGHERTY, O L TULL, HONISS COONS, A J SOMERS, GROVER S HAISLIP, F C	CUSICK WHITE WHITE WHITE WHITE WHITE	01-06-50 07-00-50 04-00-51 09-00-51 07-00-51	5 10 10	152 33 86 86 57	1.5 2 2 2 2
SO Dd SO Dd SO Dd SO Dd SO Dd	8 9 10 11 12	SO-00-8134 SO-00-1919 SO-00-2006 SO-00-3755 SO-00-6442	HAISLIP, F C WARD, PHILLIP MRS COULBOURN, DR GEORGE C. MARION FIRE CO. PALMER, HAROLD	WHITE CUSICK CUSICK WHITE CUSICK	07-00-51 10-21-47 10-25-47 04-01-49 08-12-50	10 10 10	100 91 131 81 192	1.5 1.5 2 1.5
SO Dd	13	SO-00-3815	WHITTINGTON, R BRICE	WHITE	07-00-49	10	86	2
SO Dd	14	SO-00-3814	SOMMERS, G	WHITE	07-00-49	10	84.5	2 2
SO Dd SO Dd SO Dd	15 16 17	SO-00-3816 SO-00-3434 SO-00-3817	PRICE, HOWARD PUSEY, ROY BUTLER, EARL	WHITE WHITE WHITE	07-00-49 12-10-48 07-00-49	5	72 88.5 75	2 2 2 2 2
SO Dd SO Dd SO Dd SO Dd SO Dd	18 19 20 21 22	SO-00-3058 SO-00-3209 SO-00-5984 SO-00-1674 SO-00-1651	HALL, REGINALD POWELL, HOWARD WHITINGTON, NORMAN CHAFFEY, W T CHAFFEY, W T	WHITE WHITE WHITE CUSICK CUSICK	09-10-48 10-22-48 07-00-50 08-14-47 08-09-47	10 10	87 78 55 214 220	2 2 2 1.5 1.5
SO Dd SO Dd SO Dd SO Dd SO Dd	24 25 26	SO-00-3059 SO-00-1604 SO-00-2037	POWELL, L Q GREEN, GEO A BRADSHAW, WM COULBOURN, N R STEVENSON, IRA E MRS	WHITE WHITE CUSICK CUSICK CUSICK	07-00-50 09-08-48 07-23-47 11-01-47 06-04-49	10 5 10	86 95 155 152 100	2 2 1.5 1.5
SO Dd SO Dd SO Dd	29	SO-00-2118 SO-00-7193 SO-00-5897	HAISLIP, F C ROSS, DR ALEXANDER CRISFIELD AIRPORT	WILSON CUSICK CUSICK	12-05-47 12-26-50 06-16-50	5	84 165 163	1.5 2 2.5
SO Dd SO Dd		SO-00-7663 SO-00-0762	CRISFIELD DEHYDRATING CO	CUSICK CUSICK	04-24-51 09-06-46		140 165	1.5 1.5 3.5 2.5 1.5

of casin or case inte val (ft)	ng d r-	Diam- eter of acreen (in.)	acre int	tom of aen or eened ter- al		Aquifer code	Water leval (ft)	Data water level meas- urad	Draw- down (ft)	Dia- charga (gal/ min)	Pumping period (houra)	Spe- cific capa- city [(gal/ min)/ ft]		WE	ocal
	216 7 25	2 2 2			234 12 35	122MNKN 112PCPC 112PCPC	6 1.1m 6	05-19-66 04-15-87 11-08-81	 2	14 10	2	5.0	T U C	SO CI	19
	83	2			93	122PCMK	4.5m 9	04-15-87 07-11-86	5	15	1	3.0	H	SO CI	
						112PCPC 112PCPC							H,S	SO CE	
						112PCPC	.75	02-08-52					T	SO CE	3
	200	2			220	112PCPC 122MNKN	12.5	04-23-86	1.5	43	1	29	H,S	SO C	
0	94	3			104	122PCMK	13	09-04-64	20	30	4	1.5	N	SO C	
210 -	210 810					217PTMC	+15f	07-24-62		43	8		H	SO DE	
	124	1.5			139	122PCMK 122MOCN	4	07-30-52	2	30	1	15	U H,S	SO Do	
+1.5- 250 -		4	1,128	-1,		217PTMC	+11	11-03-66	72.3	200	5	2.8	U	SO D	3
,138 -	1,262 1,272	2	1,272			217PTMC									
,287 -	1,308														
	199 1,116	6	1 116		106	0.1 7 7 77 40	+4	12-05-70	71	150	12	2.1	P	SO D	c 4
199	117	4	1,116		122	217 PTMC 122MOCN	5 3,3m	08-22-80 04-15-87	65	35	1	. 54	P	SO D	c :
+1.5-	189	.5 2	189	. 5-	199.5	112PCPC 122MNKN	8	04-18-46		15	3-		Z H	SO De	
199.5-		1.5			170	122MOCN	2	11-06-47	5	10	2	2.0	Н	SO D	
	128	1.5			152	122MOCN	1	01-06-50		21	3		Н	SO D	
	25	2			33	112PCPC	4	07-00-50		12	2		S	SO D	d
	78 75	2 2			86 86	122PCMK	7	04-00-51		10	3		H	SO D	
	46	2			56	122PCMK 122PCMK	10 2.7m	09-00-51 01-04-52		10	2 3		H S	SO D	
	90	2			100	122MOCN	3	07-00-51		8	4		H	SO D	d
	76	1.5			91	122PCMK	2	10-21-47	1	30	1	30	H	SO D	
	91 68	1.5			116 81	122MOCN 122PCMK	2 2	10-25-47	4	30 15	2 3	7.5	H P	SO D	
	182	1.5			192	122MNKN	3	08-12-50	2	28	2	14	H,S	SO D	
0 - 83 -	71 84		71	-	83	122PCMK	4	07-00-49		15	3		H	SO D	d 1
0 -	73 82	2	73	-	81	122PCMK	3	07-00-49		14	2		С	SO D	d 1
-	58	2			72	122PCMK	3	07-00-49		12	2		H,S	SO D	
0 -	76 60	2	60	-	88.5 72	122PCMK 122PCMK	5 2	12-10-48 07-00-49		10 12	2		H	SO D	
72 -	73														
	68	2			78	122PCMK 122PCMK				17 10	2 2		H	SO D	
	47	2			55	122PCMK	8	07-00-50		12	4		C	SO D	d 2
	199 205	1.5			214 220	122MNKN 122MNKN	5	08-14-47 08-09-47	3	20 25	2	5.0	S H,S	SO D	
	78				86	122PCMK	7	07-00-50		6	3		S	SO D	d 2
	85	2			95	122PCMK	6	09-08-48		10				SO D	d 2
	130				155 152	122MOCN 122MOCN		07-23-47 11-01-47		20 15	2	3.0	H,S	SO D	d 2
	84	1.5			99	122PCMK	3	06-04-49		37.5	2	19	H	SO D	
	69				84	122PCMK	2	12-05-47	10	12	4	1.2	U	SO D	
0 -	145	2			165	122MOCN	3	12-26-50 06-16-50	6	18 25	3 6		H	SO D	
84 -	148				163	122MOCN									
+1.5-	125 50				140	122MOCN	3	04-24-51 09-06-46	3	30 25	1 3	10 50	C H,S	SO D	
50 -															

TABLE 14
RECORDS OF WELLS IN SOMERSET COUNTY—Continued

[ft = foot; in. = inch; gel/min = gellona per minute; (gel/min)/ft = gellona per minute per foot; -- = no data; m = measured; + = above land aurface; SWD = Someraet Well Drilling; IWD = Idael Well Drillera]

	Loca well no.		Permit number	Owner	Contractor	Date well con- atructad	Alti- tude of land aur- face (ft)		Diam- ater of casing (in,)
	SO Dd	33		COULBOURN, N R COULBOURN, N R	CUSICK CUSICK	04-03-49 11-01-47	5	147 152	1.5
5	O Dd	35 36	SO-00-4269 SO-00-0489	RUEBEN, HERMAN CHAS D BRIDDELL INC	CUSICK	07-10-49 08-14-46	5	150 155	1.5 3.5 2.5 1.5
52	O Dd	37	SO-00-1608	NELSON, J FRANK	CUSICK	07-22-47	5	166.5	1.5
0.0	SO Dd	38	SO-00-0488	STERLING, HAROLD E, JR	CUSICK	06-20-46	5	152	3.5 2.5 1.5
0 0	SO Dd SO Dd SO Dd	39 40 41 42	SO-00-0816 SO-00-4573 SO-00-1673 SO-00-4146	LONG, JESSE L STERLING, RINGGOLD DAUGHERTY, WILLIAM B LAWSON, ALFRED J	CUSICK CUSICK CUSICK	09-26-46 09-28-49 08-01-47 06-20-49	5 5 5 5	160 152 149 150	1.5 1.5 1.5 1.5
02 02	SO Dd SO Dd SO Dd	43 44 45 46	SO-00-3789 SO-00-3790 SO-01-0067 SO-73-1289	NELSON, BENJAMIN F COULBOURN, NELSON R REVELLE, SAMUEL J GRUY FEDERAL, INC	CUSICK CUSICK CUSICK GLASCOCK	04-07-49 04-02-49 05-23-52 08-29-78	5 5 5 5	149 147 163 1,020	1.5 1.5 1.5 8.63
0.0	SO Dd	47	79-GEA-001	GRUY FEDERAL, INC	ROWAN	07-18-79	3.5	5,562	4.5 16 10.75 7.63 7.63 7.63 7.63
00 00 00	SO Dd SO Dd SO Dd SO Dd	48 49 50 51 52	SO-73-1299 SO-73-0347 SO-73-1721 SO-73-0385 SO-73-1587	DAUGHERTY, WILLIAM B MROHS METERED GAS SERVICE MT. PEER CHURCH FIME RIDGE BAPTIST CH TRINITY METHODIST CHURCH	FORD SWD SWD BUNDICK SWD	10-12-78 08-20-74 06-27-80 06-10-75 11-07-79	10 10 10 5	155 148 62 148 150	2 2 2 2 2
00 00 00	SO Dd SO Dd SO Dd SO Dd SO Dd	53 54 55 56 57	SO-73-0663 SO-71-0073 SO-73-0269 SO-73-1175	MARION FIRE CO HIGHWAY HOLINESS CHURCH SOMERSET CO. BOARD OF ED CRISFIELD ELKS CLUB CRISFIELD ELKS CLUB	SWD FORD BUNDICK FORD FORD	04-07-76 07-19-71 03-25-74 03-30-78	10 10 10 10	145 75 142 100	2 2 4 2 2
6	SO Dd	58	SO-73-0576	WHITTINGTON, THOMAS	DELMARVA DRLNG	11-07-75	10	95	4
	SO Dd	59 60	SO-73-1961 SO-73-1989	DAUGHERTY, WILLIAM ABBOTT, M E	SWD SWD	09-09-81 09-30-81	5	155 160	4
	SO Dd	61 62	SO-71-0053 SO-70-0026	EDER, G H BYRD, MELVIN	IWD FORD	11-28-70 10-20-69	10 10	154 145	3 4 2
	SO Dd SO Dd SO Dd SO Dd SO Dd	63 64 65 66 67	SO-81-0689 SO-81-0767 SO-81-0694 SO-81-0673 SO-81-0762	SCHARZ, HENRY BRITTINGHAM, JOHN FOXWELL, MACE LINTON, WILLIAM GOLDSBOROUGH, CHARLES	SWD BUNDICK SWD SWD SWD	11-06-85 12-15-85 08-18-85 07-15-85 12-05-85	8 5 5 8	160 185 160 160	2 2 2 2 2
	SO Dd SO Dd SO Dd SO Dd	68 69 70 71 72	SO-81-0684 SO-81-0782 SO-81-0897 SO-81-1073 SO-81-1088	MATARAZZO, GEORGE WHITELOCK, LELAND EVANS, WILLIAM BERKEYPILE, JIM SIEGMANN, RAY	LARSON WELLS SWD SWD SWD SWD	07-30-85 02-14-86 06-13-86 05-27-87 06-04-87	5 8 5 5 5	160 155 155 152 145	2 2 2 2 2
00 00	SO De SO Da SO De	73 1 2 3	SO-81-1009 SO-00-9081 SO-00-7608 SO-00-8343	DORMAN, LADELL GREEN, CARL PRICE, EARL WHITIINGTON, NORMAN T	SWD WHITE WHITE WHITE	11-12-86 03-00-52 04-00-51 10-00-51	5 10 10 10	165 420 81 120	2 2 2
0.0	SO De	4	SO-00-8877	CHAMBERLIN, ROBERT L	cusick	11-07-51	10	93	2
	SO De	5	SO-00-1010 SO-00-4760	BOWLAND, J E BONNEVILLE, MITCHELL	CUSICK WHITE	11-29-46 10-23-49	10 10	90 116	1.5
0.0	SO De	7	SO-00-3754	GREEN, J B	WHITE	03-23-49	10	118	2
	SO Da	8	SO-00-5501 SO-00-5500	COULBOURN, GEO DR HALL, ROGER	WHITE WHITE	04-17-50 04-24-50	10	117 91	2 2 2

ca in	of of or or ised iter (a)	8	Diam- eter of screen (in.)	scr in	ttom of reen or eened ter- eal	Aquifer code	Water level (ft)	Date weter level meas- urad	Draw- down (ft)	Dia- charge (gal/ min)	Pumping period (hours)	Spe- cific capa- city [(gal/ min)/ ft]		Local well no.	
+1. 50	5-	132 131 125 50 90	1.5 1.5 1.5		147 152 150	122MOCN 122MOCN 122MOCN	3 22	07-10-49 08-14-46	 8 4	 10	3 3	2.5	H,S H,S H,S	SO Dd SO Dd SO Dd SO Dd	33 34 35 36
90	5-	150 151.	5 1.5	150 151.		122MOCN 122MOCN	4	07-22-47	2	20	2	10	H,S	SO Dd	37
÷.	5-	50 90					3.5	06-20-46		10	3		H	SO Dd	38
90	-	134 150 142 124 125	1.5 1.5 1.5 1.5	134	- 152 160 152 149 150	122MOCN 122MOCN 122MOCN 122MOCN 122MOCN	4 2 3.5 3	09-26-46 09-28-49 08-01-47 06-20-49	.5 14 3.5 2	15 28 12 12	2 3 3 3	30 2.0 3.4 6.0	H,S H,S H	SO Dd SO Dd SO Dd SO Dd	39 40 41 42
0	_	124 132 148 136	1.5 1.5 1.5		149 147 163	122MOCN 122MOCN 122MOCN	1.5	04-07-49 04-02-49 05-23-52	4 2.5 3	18 25 30	2 2 2	4.5 10 10	H.S H U	SO Dd SO Dd SO Dd SO Dd	43 44 45
136	-	,020 165				217PTMC							U	SO Dd	47
165 1,715 3,846 4,032 4,223	-3 -3 -4	,715 ,798 ,901 ,148 ,625	7.63 7.63 7.63	3,798 3,901 4,148	-3,846 -4,032 -4,223	217PTMC 217PTMC 217PTMC									
		135 138 52 138 125	2 2 2 2 2		155 148 62 148 150	122MOCN 122MOCN 122PCMK 122MOCN 122MOCN	8 7 10 5	10-12-78 08-20-74 06-27-80 06-10-75 11-07-79	31 2 41 3	10 20 9 12 10	1 4 1 3 2	2.5 .65 4.5 .29 3.3	T	SO Dd SO Dd SO Dd SO Dd SO Dd	48 49 50 51 52
		130 65 130 70	2 4 2 2		145 75 142 100	122MNKN 122MOCN 122PCMK 122MNKN 122MOCN	6 4 18 8	04-07-76 07-19-71 03-25-74 03-30-78	2 16 4 2	10 40 10 15	2 4 2 3	5.0 2.5 2.5 7.5	TCCC	SO Dd SO Dd SO Dd SO Dd SO Dd	53 54 55 56 57
		70	4		85	122PCMK	6 3.0m	11-07-75 04-15-87	19	61	1	3.3	I	SO Dd	58
+2	_	130	4		155	122MOCN	12	09-09-81 09-30-81	3 5	10 10	1	3.3	H H	SO Dd	59 60
80	-	140 154 129	3	140	- 160 154 145	122MOCN 122MOCN 122MOCN	3.4m 10 6		20	15 20	4 2	.75		SO Dd SO Dd	61 62
		140 175 140 140 140	2 2 2 2 2		160 185 160 160 160	122MOCN 122MNKN 122MOCN 122MNKN 122MOCN	20 5 7 6 6	11-06-85 12-15-85 08-18-85 07-15-85 12-05-85	1 40 3 6	45 20 30 40	1 3 1 1	45 .50 5.0	H H H	SO Dd SO Dd SO Dd SO Dd SO Dd	63 64 65 66 67
		140 135 135 122 115	2 2 2 2		160 155 155 152 145	122MOCN 122MOCN 122MOCN 122MOCN 122MOCN	10 5 15 15 5	07-30-85 02-14-86 06-13-86 05-27-87 06-04-87	5 1 10 3 30	40 45 20 38 10	3 1 1 1	8.0 45 2.0 13	H H H H	SO Dd SO Dd SO Dd SO Dd	68 69 70 71 72
		145 70	 2			122MOCN 122CPNK 122PCMK		11-12-86 04-00-51	1	60 10	1 3	60	H U H,S	SO Dd SO Da SO De	1
		108	2		120	122PCMK	2			12	2		ΰ	SO De	3
		8.5	2		93	122PCMK	4	11-07-51	1	25	2	25	H,S	SO Da	4
0 114	-	75 106 116	1.5	106	- 114	122PCMK 122PCMK	3	11-29-46	.5	25 12	1 4	50	S H	SO De	5
0	-	106 118	2	106	- 116	122PCMK	2	03-23-49		12	2		Н	SO De	7
		107	2			122PCMK 122PCMK	3	04-17-50 04-24-50		10	3		H,S	SO Da	

[ft = foot; in. = inch; gsl/min = gallona per minute; (gal/min)/ft = gsllons per minuta per foot; -- = no dsta; m = messursd; + = above land aurface; SWD = Someraet Well Drilling; CD&F = Coastal Drilling and Pump; IWD = Ides1 Well Drillara; USGS = U.S. Geological Survsy]

Locs well no.		Permit number	Owner	Contractor	Data well con- atructad	Alti- tuds of land sur- faca (ft)	Dapth	Diam- etar of casing (in.)
SO De SO Da SO De SO De SO De	10 11 12 13 14	SO-00-5569 SO-00-5568 SO-00-1892 SO-00-0869 SO-00-3210	ADAMS, MITCHELL W LANDON, EDWARD GREEN, CARL, JR WILLIAMS, EDWARD S HALL, B J	WHITE WHITE CUSICK CUSICK CUSICK	03-30-50 04-03-50 10-14-47 10-11-46 12-20-48	5 10 10 5 5	88 112 150 108 75	2 2 1.5 1.5
SO De SO De SO De SO De SO De	15 16 17 18 19	SO-00-0837 SO-00-6191 	CHELTON, GUY WILKINS, PAUL DAVIS, H A ADAMS, F MATHEWS LUMBER & CANNING CO	CUSICK WHITE DAVIS 	10-31-46 07-00-50 00-00-49 00-00-40 00-00-48	10 10 10 10	100 87 9 48.5 75	2 2 1.25 1.25 2
SO De SO De SO De SO De	20 21 22 23 24		DASHIELL, A T MILLER, C HAYMAN, C GERALD, J TAYLOR, L	DASHIELL CUSICK	00-00-49 00-00-25 00-00-32 00-00-40	10 15 10 5 5	23 13 35 10.8	1.25 1.25 1.25 36 1.25
SO De SO De SO De SO De SO De	25 26 27 28 29	SO-00-1742 SO-73-1389 SO-73-1394 SO-81-0211	SCHUMACHER, WILLIAM C HARTMAN, CLARENCE E GRUY FEDERAL INC GRUY FEDERAL INC EBENEZER METHODIST CH	SCHUMACHER CUSICK GLASCOCK GLASCOCK SWD	00-00-51 10-06-47 11-16-78 11-18-78 04-06-83	10 10 5 5	34 120 1,037 1,051 100	1.5 1.5 4.5 4.5
SO De SO De SO Da	30 31 32 33	SO-68-0075 SO-73-0604 SO-73-1858 SO-73-1857	REHOBETH METHODIST CHURCH DASHIELL, A T WHITTINGTON, PAUL T WHITTINGTON, PAUL T	FORD FORD DELMARVA DRLNG DELMARVA DRLNG	04-30-68 01-05-76 05-06-81 04-15-81	10 10 10 10	90 100 120 130	2 2 16 4
SO Da	34	SO-81-0526	BUTLER, DENNETT	DELMARVA DRLNG	11-27-84	10	105	2
SO De SO Ds	35 36	SO-73-1866 SO-81-0560	BUTLER, DENNETT BUTLER, DENNETT	DELMARVA DRLNG DELMARVA DRLNG	04-13-81 03-04-85	10 10	110 110	4
SO De SO Da	38	SO-81-0037 SO-81-0172	HOWARD, PETE SIX L'S ENNIS FARM	CD&P CD&P	03-30-82 02-22-82	10 10	90 120	8
SO Da	40	SO-81-1019	U.S. GEOLOGICAL SURVEY	USGS	01-05-87	5	12	8 2
SO De SO De SO Df SO Df SO Df	41 42 1 2 3	S0-81-0751 S0-81-0712 S0-00-6229 S0-00-7089	HOOD, SHERMAN FULTON, WILLIAM BELL, GEO W UNDERHILL, ELIZABETH S KURTZ, JOHN	SWD SWD WHITE CUSICK	10-00-85 09-05-85 07-00-50 12-09-51 00-00-51	5 5 5 5	80 210 82 440 34	2 2 2 1.5 1.25
SO Df	5		MARRINER, L CLUFF, F	MARRINER		10 10	15 27	1.25
SO Df SO Df SO Df	6 7 8		POWELL, G DRYDEN, R L MCCREADY, S T	POWELL BEAUCHAMP	00-00-47	15 15 20	20 8.8 23	1.25 24 1.25
SO Df	9	SO-73-0375	STEVENSON EQUIP. COMPANY	BUNDICK	11-04-74	20	300	4
SO Df	10	SO-73-1975	CHESAPEAKE BAY PLYWOOD	BUNDICK	09-20-81	10	210	4
SO Df	11	SO-66-0067	CHESAPEAKE BAY PLYWOOD	SYDNOR HYDRO	03-01-66	10	178	8 6 6
SO Df SO Df	12 13	SO-73-1190 SO-81-0027	CHESAPEAKE BAY PLYWOOD LANKFORD FOODS	DELMARVA DRLNG SWD	04-12-78 12-29-81	10 20	130 280	6 10 2
	15	SO-67-0012 SO-81-0098 SO-81-0065 SO-73-1694 SO-73-1890	LANKFORD, ARTHUR W, JR HOLLY GROVE MENNONITE CH. HOLLY GROVE MENNONITE CH. COWGER, PAUL W BUTLER, DENNETT	MARSHALL SWD SWD L E COWGER DELMARVA DRLNG	08-09-66 07-05-82 05-20-82 05-30-80 04-10-81	20 20 20 20 10	69 85 80 20 90	2 2 2 2 4
	20	SO-81-0785	BUTLER, DENNETT BUTLER, DENNETT BUTLER, DENNETT	DELMARVA DRLNG DELMARVA DRLNG	03-27-85 00-00-81	10 10 5	130 105 120	4
SO Df	22	SO-73-1877	VESSEY'S ORCHARD	DELMARVA DRING		10	130	4
SO Df	23	SO-81-0061	VESSEY, WILLIAM	COASTAL DRLNG	04-29-82	10	100	6

Bottom of caains		Bottom of screen			Date				Spe- cific			
or cased inter- val (ft)	Diam- eter of screan (in.)	or ecreened inter- val (ft)	Aquifer code	Water lavel (ft)	water leval meas- urad	Draw- down (ft)		Pumping period (hours)	capa- city [(gal/ min)/ ft]	Uae of water	Local well no.	
10		86 112 150 108 75	122 PCMK 122 PCMK 122 MOCN 122 PCMK 122 PCMK	2 3.5 5 .5	03-30-50 04-03-50 10-14-47 12-20-48	5 .2	8 8 8 25 35	4 5 4 1	1.6 125 5.8	H H H,S H,S	SO Da SO De SO De SO De SO De	10 11 12 13 14
	38 2 31 2 	100 87 	122PCMK 122PCMK 112PCPC 112PCPC 122PCMK	2 4	10-31-46 07-00-50 	.5	30 10 	1 3	60	H,S H,S H,N	SO Da SO Da SO De SO De SO De	15 16 17 18 19
			112PCPC 112PCPC 112PCPC 112PCPC 122PCMK	3.0m -6m	03-26-52 03-26-52					P H H,S H,S	SO De SO Da SO Da SO Da	20 21 22 23 24
1,00 1,00		120 100	112PCPC 122PCMK 125PLCN 125PLCN 122PCMK	3 5	10-06-47 04-06-83	2 5	20 20	2	10 4.0	H,S H,S U U	SO De SO De SO De SO De SO De	25 26 27 28 29
	82 2 85 2 38 16 38 4	90 100 118 118	122PCMK 122PCMK 122PCMK 122PCMK	7 5 27 5 3,9m	04-30-68 01-05-76 05-06-81 04-15-81 04-15-87	2 35 8	10 10 400 60	2 3 2 2	5.0 11 7.5	H I U	SO De SO De SO De	30 31 32 33
	37 4	97	122PCMK	4	11-27-84	5	35	3	7.0	I	SO De	34
	40 4 40 4	100 100	122 PCMK 122 PCMK	3	04-13-81 03-04-85	3 13	60 74	2.5	20 5.7	I	SO Da	35 36
0 -	50 4 20 8 60 8 7 2	20- 40 60- 120 12	122PCMK 122PCMK 122PCMK 112PCPC	2.8m 5 10 2.5m 0.8m	12-10-87 03-30-82 02-22-82 04-15-87 04-15-87	5 10	200 500	5	40 50	I I	SO De SO De	38 39 40
1	60 2 85 2 70 2 20 1.5	80 210 82 440	112PCPC 122MNKN 122PCMK 122CPNK 112PCPC	5 10 3 1.5	10-00-85 09-05-85 07-00-50 12-09-51	3 1 6.5	35 35 8 30	1 1 4 6	12 35 4.6	H H,S H,S	SO De SO De SO Df SO Df SO Df	41 42 1 2 3
			112PCPC 112PCPC 112PCPC 112PCPC 112PCPC	1.32m	03-25-52					H,S H,S H S	SO Df SO Df SO Df SO Df	5 6 7 8
	50 4	290	122MNKN	18 20.2m	11-04-74 04-15-87	54	10	5	.19		SO Df	9
100- 1 0- 91- 109- 1	00 90 2 81 6 99 6 16 6	190- 210 81- 91 99- 109 16- 126	122MNKN 122PCMK 122PCMK 122PCMK	18	09-20-81	53	307	2 24	5.8	N	SO Df	10
1	10 10 50 2	130 280	122PCMK 122MNKN	15 12	04-12-78 12-29-81	13	300 10	6	23	N	SO Df	12 13
	65 2 70 2 70 2 16 2 10 4	69 85 80 20 78	122 PCMK 122 PCMK 122 PCMK 112 PCPC 122 PCMK	7 10 8 3	07-05-82 05-20-82 05-30-80 04-10-81	5 5 2	12 15 10 5	1 1 1 2.5	3.0 2.0 25	H T C I	SO Df SO Df SO Df SO Df SO Df	16 17
	59 4 25 4 	119 76 120 118	122 PCMK 122 PCMK 122 PCMK 122 PCMK	3 1 5.7m	04-09-81 03-27-85 04-15-87 04-08-81	19	60 150 39 60	2.5	30 21 2.1 12	I U I	SO Df SO Df SO Df	21
	80 6	100	122PCMK	3.0m 15	04-15-87 04-29-82	5	200	1	40	I	SO D£	23

[ft = foot; in. = inch; gal/min = gallons per minute; (gal/min)/ft = gallona per minute par foot; -- = no data; m = meaaured; + = above land aurface; f = flowing well; SWD = Somerset Well Drilling; IWD = Ideal Well Drillera; SAW = Shannahan Art. Well; PI&CS = The Packers Ice and Gold Storaga Company]

Local well no.		Permit number	Owner	Contractor	Date well con- structed	Alti- tude of land sur- face (ft)	Depth	Diam- atar of caaing (in,)
SO Df SO Df SO Df	24 25 26	SO-73-1876	VESSEY, WILLIAM DRYDEN, ELMO DRYDEN, ELMO	COASTAL DRLNG DELMARVA DRLNG DELMARVA DRLNG		20	95 110 100	8 4 4
SO Df SO Df	27 28	SO-81-0593 SO-81-0859	MALIN, GEORGE BIG APPLE MARKET	SWD SWD	04-16-85 06-21-86		80 240	2 4 2
SO D8 SO D8 SO Dg	1 2 4	SO-66-0066	JOHNSON MEAT PRODUCTS INC JOHNSON MEAT PRODUCTS INC CHESAPEAKE BAY PLYWOOD		00-00-42 00-00-39 03-15-66	5	95 95 154	2 2 8 6 6
SO D8 SO D8	5		CORBIN MEAT PACKING CO CHOPTANK ELECTRIC COOP	BUNDICK IWD	08-08-73 07-26-73		200 85	6 2 4
SO Dg SO Dg SO Dg SO Dg SO Ea	7 8 9 10 1	SO-73-0137 SO-67-0100 SO-66-0087 SO-81-0834 SO-00-0023	SOMERSET CO. LIQUOR BOARD RIVERVIEW MARKET SOMERSET PACKING COMPANY MARSHALL, DORSEY LORA C & BEN WHITELOCK	FORD FORD FORD SWD CUSICK	07-17-73 02-03-67 07-22-66 04-14-86 09-20-45	10 10 10	100 78 98 220 841	2 2 2 2 3.5 1.5
SO Ea	2	SO-00-0047	BRADSHAW, HARVEY	CUSICK	10-11-45	2	840	3.5
SO Ea	3	SO-00-0076	EVANS, WILLIE	CUSICK	10-31-45	2	865	1.5 2.5
SO Ea	4	SO-00-0077	EVANS, CHARLTON	CUSICK	02-09-46	2	852	1.5
SO Ea	5	SO-00-0133	MIDDLETON, CLAYTON	CUSICK	03-27-46	2	850	1.5 2.5
SO Ea	6	SO-00-0295	EVANS, MARY W. MRS	CUSICK	05-10-46	2	860	1.5 2.5 1.5
SO Ea	7	SO-00-2950	HOFFMAN, ROLAND MRS	CUSICK	09-23-48	2	848	2.5
SO Ea	8	SO-00-2951	EVANS, MILTON	CUSICK	09-02-48	2	871	1.5
SO Ea	9	SO-00-3120	MARSH, ARCHIE MRS	CUSICK	11-11-48	2	915	1.5
SO Ea		SO-70-0030	TYLER, SCHULTZ FOWLER, JOHN	FORD	00-00-15 03-17-70		7.2 1,060	1.5 24 2
SO Ea SO Ea SO Ec	12 13 1	SO-70-0042 SO-72-0047	MARSH, CHARLIE EWELL WATER WORKS CITY OF CRISFIELD	FORD DELMARVA DRLNG SAW	03-12-70 11-15-71 12-00-37	2	940 960 994	4 8
SO Ec	2		CITY OF CRISFIELD	SAW	00-00-38	5	997	6 8
SO Ec	3		CITY OF CRISFIELD	SAW	00-00-28	5	1,076	10
								6
SO Ec	4	SO-00-2205	CITY OF CRISFIELD	LAYNE ATLNTC	04-24-48	5	1,302	18 8 6
								6
SO Ec SO Ec SO Ec	5 6 7	SO-00-1652 SO-00-5379	GEO A CHRISTY & SON MASSEY CHEV SALES PI&CS	SAW CUSICK SAW	00-00-10 08-06-47 05-12-50	5	1,011 81 1,042	6 1.5 8 6
SO Ec	8		PI&CS	SAW	1895	2	1,018	5 8

Case		tom		Bott								C			
35 4 100 122PCMK 10 04-06-81 5 50 2 510 I SOD MADE NO. 100 122PCMK 10 04-06-81 5 50 2 101 I SOD MADE NO. 100 122PCMK 10 04-06-81 5 50 2 101 I SOD MADE NO. 100 122PCMK 10 04-06-85 3 10 1 3.3 H SOD MADE NO. 100 122PCMK 10 04-16-85 3 10 1 3.3 H SOD MADE NO. 100 122PCMK 10 04-16-85 3 10 1 3.3 H SOD MADE NO. 100 122PCMK 10 04-16-85 3 10 1 3.3 H SOD MADE NO. 100 122PCMK 10 04-16-85 3 10 1 3.3 H SOD MADE NO. 100 122PCMK 10 04-16-85 3 10 1 3.3 H SOD MADE NO. 100 122PCMK 10 04-16-85 3 10 1 3.3 H SOD MADE NO. 100 122PCMK 10 04-16-85 3 10 1 3.3 H SOD MADE NO. 100 122PCMK 10 04-16-85 3 10 1 3.3 H SOD MADE NO. 100 122PCMK 10 04-16-85 5 3 10 1 2.0 C SOD MADE NO. 100 122PCMK 10 04-16-85 5 3 10 1 3.3 H SOD MADE NO. 100 122PCMK 10 04-16-85 122PCMK 10 100 122PCMK 10 04-16-85 122PCMK 10 04-16-85 122PCMK 10 100 122PCMK 10 04-16-85 122PCMK 10 100 122PCMK 10 04-16-85 122PCMK	caa o cas int	ing r ed er-	of acreen	acree inte	en ned r-		level	watar lavel maaa-	down	charga (gal/	period	capa- city [(gal/ min)/	of	Local wall no.	L
70		3.5	4		100	122PCMK	10 10	04-07-81 04-06-81	5	50	2.5	10	I	SO Df SO Df SO Df	2 2 2
200 2 2 200 2 4 122MNN	+1-		2		80	122PCMK	10	04-16-85						SO Df	2
122PCMK	20-	200	2	200-	240	122MNKN									
0- 80 6 80- 95 122PCMK 15 03-15-66 59 205 24 3.5 N SO D 95- 100 6 100- 110 122PCMK 16 116- 126 122PCMK 18 08-08-73 30 30 4 1.0 C SO D 12 PC SO														SO Dg SO Dg	
- 180	95-	100 116	6	100-	110	122PCMK 122PCMK	15	03-15-66	59	205	24	3.5		SO Dg	
70	-	180												SO Dg SO Dg	
0- 60		70 89 195 60	2 2		78 98	122PCMK 122PCMK	8 5 8	02-03-67 07-22-66 04-14-86	1	10 18 35	2 2 1	35	C C H	SO Dg SO Dg SO Dg SO Dg SO Ea	
0- 567 67- 819 1 567- 840 217PTMC 0- 576 76- 837 0- 567 67- 830 1.5 830- 850 217PTMC 0- 567 67- 830 1.5 830- 850 217PTMC 0- 567 67- 830 1.5 830- 850 217PTMC 0- 567 67- 800 1.5 800- 860 217PTMC 0- 567 0- 567 0- 567 0- 567 0- 567 0- 567 0- 567 0- 567 0- 567 0- 567 0- 568 0- 567 0- 568 0- 567 0- 568 0- 567 0- 568 0- 567 0- 568 0- 568 0- 567 0- 568 0- 568 0- 567 0- 568 0- 568 0- 568 0- 567 0- 568 0- 568 0- 567 0- 568 0- 568 0- 568 0- 567 0- 568 0- 568 0- 567 0- 568 0- 568 0- 567 0- 568 0- 568 0- 567 0- 568 0- 568 0- 567 0- 568 0- 568 0- 567 0- 568 0- 568 0- 567 0- 568 0- 568 0- 567 0- 568 0- 56			PP 40			217PTMC									
0- 576 76- 837 0- 567 0- 567 0- 567 0- 567 67- 830 1.5 830- 850 217PTMC f 03-27-46	-						+12f	11-17-53					P	SO Ea	
0- 567 67- 830	0 -	576	1	567-	840		+15f	11-17-53	11				P	SO Ea	
0- 567 67- 830	0 -	567					f	02-09-46	1	w w			P	SO Ea	
0- 567 67- 800	0-	567					£	03-27-46					P	SO Ea	
0- 588 88- 808 0- 567 67- 840 0- 638 38- 889 1 889- 915 217PTMC 7.2 112PCPC 3.9m 11-17-53 25 1 P SO I 990 1,010 217PTMC +8 03-17-70 8 20 2 2.5 H SO I 990 940 217PTMC +8 03-17-70 8 20 2 2.5 H SO I 920 940 217PTMC +8 03-17-70 8 20 2 2.5 H SO I 920 940 217PTMC +3 03-12-70 3 50 2 16 H SO I 920 940 217PTMC f 11-15-71 25 40 P SO I 0- 245 45- 944 6 944- 994 125PLCN f 11-15-71 25 40 P SO I 45- 944 6 944- 994 125PLCN f 00-00-38 300 24 P SO I 45- 941 6 941- 995 125PLCN f 00-00-38 300 24 P SO I 45- 941 6 941- 995 125PLCN f 00-00-28 210 P SO I 50-1,042 4.5 1,042-1,076 217PTMC 0- 85 50-1,042 4.5 1,042-1,076 217PTMC 0- 86 81 819- 829 125PLCN 20 00-00-28 210 P SO I 17.7 04-15-87 17.7 04-15-87 17.7 04-15-87 18.89 6 819- 829 125PLCN 20 00-00-28 210 N SO I 46 81 122PCMK 3 08-06-47 3 20 30 6.7 C SO I 0- 283	0-	567					+15f	11-17-53					P	SO Ea	
88- 808			2.5	000	000	21/1110	+15f	11-17-53	NO. 100	25	1	ato ato	P	SO Ea	
67- 840 0- 638 38- 889	88-	808				217 PTMC					_			SO Ea	
38-889		840				217PTMC								SO Ea	
990 1,010 217FTMC +8 03-17-70 8 20 2 2.5 H SO 1 920 940 217FTMC +3 03-12-70 3 50 2 16 H SO 1 900 4 922 217FTMC f 11-15-71 25 40 P SO 1 0- 245 3 12-00-37 54 100 1.9 P SO 1 45- 944 6 944- 994 125FLCN f 00-00-38 300 24 P SO 1 45- 941 6 941- 995 125FLCN f 00-00-38 300 24 P SO 1 45- 941 6 941- 995 125FLCN 20 00-00-28 210 P SO 1 25- 457 57- 950 50-1,042 4.5 1,042-1,076 217FTMC 0- 85		889		889-										SO Ea	
900 4 922 217PTMC f 11-15-71 25 40 P SO I					1,010		+8		8	20	2	2.5		SO Ea	
45- 944 6 944- 994 125PLCN f 00-0-38 300 24 P SO I 45- 941 6 941- 995 125PLCN g 00-00-28 210 P SO I 25- 457 57- 950 50-1,042 4.5 1,042-1,076 217PTMC 0- 85		900					f	11-15-71					P	SO Ea	
45- 941 6 941- 995 125PLCN 20 00-028 210 P SO I 25- 457 57- 950 50-1,042 4.5 1,042-1,076 217PTMC	45-	944	6	944-	994	125PLCN								SO Ec	
25- 457 57- 950 50-1,042	45-	941	6	941-	995	125PLCN								SO Ec	
50-1,042 4.5 1,042-1,076 217PTMC 0- 85	25-	457					20	00-00-28		210			P	SO Ec	
85- 375 75- 726			4.5	1,042-	1,076	217PTMC									
29-1,136 6 1,136-1,146 217FTMC N SOI 46 81 122PCMK 3 08-06-47 3 20 .30 6.7 C SOI 0- 283 +4 05-12-50 56 200 48 3.6 N SOI	85- 75-	375 726		726-	731				104	300	48	2.9	P	SO Ec	
46 81 122PCMK 3 08-06-47 3 20 .30 6.7 C SO 1 0- 283 +4 05-12-50 56 200 48 3.6 N SO 1		1,136	6	1,136-	1,146	217 PTMC							NY.	00.0	
83- 951 217PTMC	0-	46 283				122PCMK	3	08-06-47	3	20	.30	6.7	C	SO Ec SO Ec	
38-1,002 6 1,002-1,027			6	1,002-		217PTMC	10 4-	01-29-54		20			11	SO Ec	

[ft = foot; in, = inch; gal/min = gallons per minute; (gal/min) ft = gallons per minute per foot; -- = no data; m = measursd; + = abovs land surface; SWD = Somerset Well Drilling; PI&CS = The Packsrs Ice and Cold Storage Company]

Local wsl1 no.	-14	Psrmit number	Owner	Contractor	Date well con- structed		Depth drilled (ft)	Diameter of casing (in.)
SO Ec	9		PI&CS		1892	2	1,060	8
SO Ec	10	SO-00-0585 SO-00-2220	BOZMAN MOTOR CO MCCREADY MEM. HOSPITAL	CUSICK	08-30-46 05-05-48	5	58 384	1.5 2.5 1.5
SO Ec	12	SO-00-0663	HINMAN, WILMER	WHITE	04-00-48	5	150	2
SO Ec	13	SO-00-2482	EMELY, STEWART	CUSICK	05-22-48	5	196	1.5
O Ec	14 15 16 17 18	SO-00-2481 SO-00-1920 SO-00-7192 SO-00-1605 SO-00-0870	BYRD, WILLIAM R, MRS WARD, M L RIGGIN, OTIS J R LAIRD & E BELL BYRD, BENNETT	CUSICK CUSICK CUSICK CUSICK	05-17-48 10-23-47 12-30-50 07-07-47 10-08-46	5 5 5 5 5	186 196 188 198 192	1.5 1.5 1.5 1.5 2.5 1.5
30 Ec	19	SO-00-0589	AYRES, PRESTON	CUSICK	08-17-46	5	183	3.5
SO Ec	20	SO-00-2568	PARKS, NELLIE	CUSICK	05-12-48	5	193	1.5
SO Ec	21 22	SO-00-2431 SO-00-0588	LOWE, W H, JR REESE, JAMES B	CUSICK	04-23-48 06-29-46	5	193 189	1.5 3.5 2.5
O Ec	23	SO-00-0761	THORNTON, MILFORD J	CUSICK	08-28-46	5	179	1.5
O Ec	24	SO-00-6099	OWENS, EDWARD	CUSICK	06-21-50	5	195	1.5
O Ec	25	SO-00-2949	RYLE, WILLIAM, JR	CUSICK	08-06-48	5	211	1.5
SO Ec SO Ec	26 27 28	SO-00-1617 SO-00-1582 SO-00-1609	JONES, WALTER MCINTOSH, JOHN LAIRD, CHARLES T	CUSICK CUSICK	07-17-47 07-01-47 07-11-47	5 5 5	214 198 210	1.5 1.5 1.5
	29	SO-00-4270	JONES, WALTER	CUSICK	07-16-49	5	205	1.5
SO Ec SO Ec	30 31	SO-00-2432 SO-00-0664	BYRD, MERRILL O SR DAUGHERTY, C HUBBARD	CUSICK WHITE	04-08-48 08-15-46	5 5	362 95.5	1.5
SO Ec	32 33	SO-00-3431 SO-00-2984	WARD, OTIS WARD, LEROY	CUSICK CUSICK	12-24-48 08-14-48	5 3	360 362	1.5
SO Ec SO Ec	34 35 36 37 38	SO-00-2350 SO-00-2747 SO-00-2705	MORGAN, WILLIAM MOSHER, EARL J DIZE, EARL H STANT, ALVIN HENDERSON, EARL	CUSICK CUSICK CUSICK	05-27-48 06-05-48 05-29-48 00-00-49 00-00-27	5 5 5 5	151 152 152 62.5 4.25	1.5 1.5 1.5 1.5
SO Ec SO Ec SO Ec SO Ec	39 40 41 42	SO-01-0636 SO-04-6301	MORGAN, BARNEY CULLEN, LENA M REVELLE, G BRYCE CITY OF CRISFIELD	CUSICK LAYNE ATLNTC	00-00-32 1802 09-16-52 05-25-62	5 5 5 5	7.5 8.2 189 1,207	30 30 1.5 8 8
SO Ec	43	SO-02-1762	HIS NIBBS SHIRT CORP	S. SHANNAHAN	CO 06-28-5	6	91	6
SO Ec SO Ec	46 47 48	SO-05-0273 SO-72-0054	SEARS, FRANK W CITY OF CRISFIELD CITY OF CRISFIELD	FORD LAYNE-ATLNTC LAYNE-ATLNTC	01-16-63 09-13-72	5 5 5	133 1,113 1,216	2 8
SO Ec	49	SO-81-0413	CITY OF CRISFIELD	SYDNOR HYDRO	07-21-84	5	1,365	8 12 8
SO Ec	50		JOHN T HANDY, INC			5		8 8 8 8
		50-72-1006		chio	00-20-02		200	
SO Ec SO Ec SO Ec SO Ec	51 53 54 55	SO-73-1986 SO-81-0970 SO-81-0803 SO-81-1158 SO-00-0163	DRYDEN CAROL CO INC PARKS, RALPH NELSON, DEBRA REVELLE, BRICE MORRIS, RALPH	SWD SWD SWD WHITE	09-22-81 09-16-86 03-20-86 08-10-87 11-15-45	5 5 5 5	200 200 200 200 144	2 2 2 2 2

Bottom of casing or caaed inter- val (ft)		Diam- eter of acreen (in,)	Bottom of acreen or acreened inter- val (ft)		Aquifer code	Water level (ft)	Date water level meas- ured	Draw- down (ft)	Dis- charge (gal/ min)	Pumping period (hours)	Spe- cific capa- city [(8al/ min)/ ft]	Use of water	Local well no.	
0- 296-	43 340 359		359	 58 - 384	217 PTMC 122 PCMK 122 CPNK	+12 5 +1.5	1892 08-30-46 05-05-48	5.5	130 20 40	2 8	7.3	U C U	SO Ec SO Ec	9 10 11
0-	130.	5 2	130				04-00-48		3			Н	SO Éc	12
	186			196	122MNKN	2	05-22-48	2	25	1	13	Н	SO Ec	13
0-	176 186 178 188 120			186 196 188 198	122MNKN 122MNKN 122MNKN 122MNKN	+1.5 .5 5	10-23-47 12-30-50 07-07-47 10-08-46	3 1.5 5.5 1.5	30 26 28 20 20	1 2 2 2 2	10 17 5.1 13 40	H H H H	SO Ec SO Ec SO Ec SO Ec	14 15 16 17 18
120-	182		182	- 192	122MNKN									
50-	50 90		470			1.8	08-17-46	. 2	10	3	50	Н	SO Ec	19
90- 0- 50-	170 181 183 50 90		170	- 183 193 193	122MNKN 122MNKN 122MNKN	2 1.2 2.2	05-12-48 04-23-48 06-29-46	2 . 8	25 30 10	10 2 2	13 38	H H H	SO Ec SO Ec	20 21 22
90~	171 169		171	- 189 179	122MNKN 122MNKN	4	08-28-46		20	2		Н	SO Ec	23
	185 199 189 188 195	1.5		195 211 214 198 210	122MNKN 122MNKN 122MNKN 122MNKN 122MNKN	1.5 6 4.5 3	06-21-50 08-06-48 07-17-47 07-01-47 07-11-47	7.5 3 .5 1.5	22 25 20 20 20	2 2 2 1 2	2.9 8.3 40 13	н н н н	SO Ec SO Ec SO Ec SO Ec	24 25 26 27 28
0- 83.5-	180 342 75.	5	75.	205 362 5- 83.5	122MNKN 122CPNK 122MOCN	2 1.3 3	07-16-49 04-08-48 08-15-46	6 . 7	19 30 30	3 3 2	3.2 43	H H U	SO Ec SO Ec	29 30 31
	340 338	1.5		360 362	122CPNK 122CPNK	+ .5 +1.2	12-24-48 08-14-48	5.5 3.2	28 30	3	5.1 9.4	H H,S	SO Ec	32 33
	116 112 117	1.5		151 152 152 	122MOCN 122MOCN 122MOCN 122PCMK 112PCPC	3 3 1 	05-27-48 06-05-48 05-29-48 04-21-52	5 12 5	10 10 10	2 1 1 	2.0 .83 2.0	H H H,S	SO Ec SO Ec SO Ec SO Ec	34 35 36 37 38
0- 930-1 1,083-1 1,090-1 1,105-1	1,078 1,085 1,100	8 8 8 8	915 1,078 1,085 1,100	189 - 930 -1,083 -1,090 -1,105	112PCPC 112PCPC 122MNKN 125PLCN 217PTMC 217PTMC 217PTMC	2.6m 2.2m 3 51.7 1.9m	04-21-52 04-21-52 09-16-52 05-25-62 04-16-87	 1 98.3	30 415	2	30 4.2	H H H P	SO Ec SO Ec SO Ec SO Ec	39 40 41 42
	80	6		91	122MOCN	8	06-28-56	38	50	10	1.3	N	SO Ec	43
	121	2		133	122MOCN	5	01-16-63		14	2		Н	SO Ec	46
	1,121	8	1,121	-1,156	217PTMC 217PTMC	20	09-29-72	71	500	9	7.0	U P	SO Ec	47
1,156-1 916- 952- 976-1 1,096-1 1,130-1 1,269-1 1,307-1	916 922 970 1,090 1,112 1,249	8 8 8 8 8	922 970 1,090 1,112 1,249 1,295 1,312	- 952 - 976 -1,096 -1,130 -1,269 -1,307 -1,320	125PLCN 125PLCN 217PTMC 217PTMC 217PTMC 217PTMC 217PTMC	10	07-21-84	235	800	24	3.4	P	SO Ec	49
	170	2		200		1.0	00-20-81	_	10			С	SO Ec	50
	180 180 180 134	2 2 2 2		200 200 200 200 144	122MNKN 122MNKN 122MNKN 122MNKN 122MOCN	10 8 3 3	09-22-81 09-16-86 03-20-86 08-10-87 11-15-45	2 1 2 1	10 50 40 60 20	1 1 1 1 3	5.0 50 20 60	C H H H	SO Ec SO Ec SO Ec SO Ec	51 53 54 55

TABLE 14
RECORDS OF WELLS IN SOMERSET COUNTY—Continued

[ft = foot; in. = inch; gal/min = gsllons per minute; (gal/min)/ft = gsllona per minute per foot; -- = no data; m = maaaured; + = above land surfsce; SWD = Somerset Well Drilling]

Local well no.		Permit number	Owner	Contractor	Date well con- structed	Alti- tude of lsnd sur- face (ft)	Depth drilled (ft)	Diam- eter of caaing (in.)
SO Ed SO Ed SO Ed SO Ed SO Ed	2 3 4 5 6	SO-00-0800 SO-00-7920 SO-00-7644 SO-00-7503 SO-00-0817	COX, WILLIAM S CHAS D BRIDDELL INC MCCREADY, DORA STERLING, CLEMENT R STERLING, L T	CUSICK LAYNE-ATLNTC CUSICK CUSICK CUSICK	09-12-46 05-28-51 04-21-51 04-06-51 09-18-46	5 5 5 5	210 70 398 202 64	1.5 6 1.5 1.5
SO Ed SO Ed SO Ed SO Ed SO Ed	7 8 9 10 11	SO-00-8216 SO-00-7556 SO-00-6380 SO-00-5628 SO-00-3478	SOMERS, MELISSA SALTZ, S M BRADSHAW, RICHARD STEPHERS, JAMES F MRS SALTZ, SAM	CUSICK CUSICK CUSICK CUSICK CUSICK	07-14-51 04-11-51 08-08-50 04-19-50 12-06-48	5 5 5 5	210 198 211 212 215	1.5 1.5 1.5 1.5
SO Ed SO Ed SO Ed SO Ed	12 13 14 15	SO-00-4843 SO-00-2065 SO-00-2005 SO-00-0366	THORTON, JAMES DIZE, SHERMAN CULLEN, GEORGE T, JR NELSON, ALONZO	CUSICK CUSICK CUSICK CUSICK	10-08-49 11-15-47 10-29-47 05-25-46	5 5 5 5	198 200 195 189	1.5 1.5 1.5 3.5 2.5
SO Ed	16	SO-00-7555	STERLING, SILAS MRS	CUSICK	04-16-51	5	193	1.5
SO Ed	17	SO-00-0482	JUSTICE, MAUDE	CUSICK	06-02-46	5	188	3.5
SO Ed	18	SO-00-0481	STERLING, JACKSON	CUSICK	05-30-46	5	190	1.5 3.5 2.5
SO Ed SO Ed	19 20 21	SO-00-5685 SO-00-6231 SO-00-0742	TODD, WILLIS TYLER, FRED MADDOX, N	CUSICK CUSICK CUSICK	05-04-50 08-02-50 09-24-46	5 5 5	196 191 194	1.5 1.5 1.5 1.5
SO Ed SO Ed SO Ed SO Ed SO Ed	22 23 24 25 26	SO-00-3595 SO-00-3134 SO-00-1581 SO-00-0743	BLADES, ALVIN MADDRIX, PAUL E NELSON, ALONZO K MASSEY, CARLTON HINMAN, HOWARD	CUSICK CUSICK CUSICK CUSICK CUSICK	03-01-49 11-20-48 11-00-51 06-28-47 08-21-46	5 5 5 5 5	204 188 197 198 187	1.5 1.5 1.5 1.5
50 Ed 50 Ed 50 Ed 50 Ed 50 Ed	27 28 29 30 31	SO-00-2003 SO-00-4110 SO-00-3791 SO-00-4842 SO-00-4876	STERLING, BURNS ENNIS, ARZAH R CRISFIELD COUNTRY CLUB TAWES, WELLINGTON STERLING, ELIJAH	CUSICK CUSICK CUSICK CUSICK CUSICK	10-25-47 06-10-49 04-14-49 10-12-49 10-15-49	5 5 5 5	195 208 200 200 210	1.5 1.5 1.5 1.5
SO Ed SO Ed SO Ed SO Ed SO Ed	32 33 34 35 36	SO-00-2499 	PRICE, HOWARD DIZE, V MARINERS CHURCH HICKMAN, ERNEST BEDSWORTH, HENRY	WHITE	07-20-48 00-00-17 1877	5 5 5 5 5	54 6.8 6.85 86.1	2 1.5 28 2.3
SO Ed SO Ed SO Ed SO Ed SO Ed	37 38 39 40 41	SO-01-0637	PRUITT, WILLIAM WARD, IONA MORGAN, HETTIE STERLING, KENNETH WARD, WELLINGTON	CUSICK CUSICK	1892 00-00-22 03-00-51 08-18-52	5 5 5 5	10.6 7.7 9.4 203 201	1.8 2 2.5 1.25 1.5
SO Ed SO Ed SO Ed SO Ed SO Ed	42 43 44 45 46	SO-73-0286 SO-73-0483 SO-81-0090 SO-81-1020	CHRISTIAN HOLINESS CH LAWSON, ALFRED SOMERSET CO. RECR/PARKS ROACH, MELVIN U.S. GEOLOGICAL SURVEY	FORD FORD SWD SWD	07-26-74 06-20-75 06-28-82 01-05-87	5 5 10 5 5	200 75 210 8	2 2 2 2
SO Ed	47	SO-81-0917	TYLER, WILLIAM	SWD	07-15-86	5	210	4
SO Ed SO Ea SO Ea	48 49 1 2	SO-81-0765 SO-81-0941 SO-00-0854	MILLINER, TIMOTHY RIGGIN, EARL HANDY, JOHN T HALL, S	SWD SWD CUSICK	12-04-85 07-29-86 10-16-46	5 5 5 5	205 205 92 8.7	2 2 1.5 1.5
SO Ef SO Ef SO Ef SO Ef SO Ef	1 2 3 4 5	SO-00-6830 SO-67-0079 SO-73-0764	HOWARD, H M MILBOURNE, J E MORRELL, G O GRAY, S MADDOX, FRED YOUNG, GEORGE E, JR	WHITE GRAY FORD KAUFFMAN	10-00-50 00-00-32 00-00-51 11-24-66 08-04-76	5 5 5 10 5	73 12.5 95 15 104 120	2 24 1.25 1.25 2 4

Bott of cssi or csse inte vsl (ft)	ng d	Diam- eter of screen (in,)	scr in v	of reen or eene ter-	ı ed	Aquifer code	Water level (ft)	Date wster level mess- ured		Dis- charge (gs1/ min)	Pumping period (hours)	min)/	Uae of wster	Locs well	
	200 60 378 182 56	6 1.5			210 70 398 202 64	122MNKN 122MNKN 122CPNK 122MNKN 122PCMK	3.7 0 2 3	09-12-46 05-28-51 04-21-51 04-06-51 09-18-46	8.3 3 12 .5	25 3 100 30 10 15	2 48 3 3 2	50 12 10 .83	H,S N H H	SO Ed SO Ed SO Ed SO Ed SO Ed	2 3 4 5 6
	190 178 182 192 189				210 198 211 212 214	122MNKN 122MNKN 122MNKN 122MNKN 122MNKN	1.5 6 2 2	07-14-51 04-11-51 08-08-50 04-19-50 12-06-48	5 7.5 6 6 3	16 15 8 25 15	3 2 2 2 2 3	3.2 2.0 1.3 4.2 5.0	H H H H	SO Ed SO Ed SO Ed SO Ed SO Ed	7 8 9 10
0-	188 185 185 50 90				198 200 195	122MNKN 122MNKN 122MNKN	3 .5 1.5 1.8	10-08-49 11-15-47 10-29-47 05-25-46	13 2.: 3.:	5 25	3 2 2 1	2.7 8.0 7.1 63	H H H	SO Ed SO Ed SO Ed SO Ed	12 13 14 15
90-	179 183	1.5	179	-	189 193	122MNKN 122MNKN	2	04-16-51	5	30	2	6.0	Н	SO Ed	16
50-	50 90						1.8	06-02-46		4 25	1	63	Н	SO Ed	17
90- 0- 50-	178 50 90 180 186 181	1.5	178	-	188	122MNKN	. 2	05-30-46	2	25	1	12.5	Н	SO Ed	18
90-		1.5	180	-	190 196 191 194	122MNKN 122MNKN 122MNKN 122MNKN	1.5 2 2.5	05-04-50 08-02-50 09-24-46	1.:	20	2 3 1	19 2.9 40	H H H	SO Ed SO Ed SO Ed	19 20 21
	194 156 187 188 175	1.5			204 197 198 187	122MNKN 122MNKN 122MNKN 122MNKN 122MNKN	1.5 .5 1 2.5	03-01-49 11-20-48 06-28-47 08-21-46	1.:	5 22	2 5 2	7 8.8 15	H C H H	SO Ed SO Ed SO Ed SO Ed SO Ed	22 23 24 25 26
	185 188 188 185 190	1.5			195 208 200 200 210	122MNKN 122MNKN 122MNKN 122MNKN 122MNKN	1.7 3 0 4	10-25-47 06-10-49 04-14-49 10-12-49 10-15-49	1.: 4 4 21 28	3 25 22 25 22 20	1 2 2 2 2	19 5.5 6.2 1.1	C H,S H S	SO Ed SO Ed SO Ed SO Ed SO Ed	27 28 29 30 31
	40.5				52.5	122PCMK 112PCPC 112PCPC 112PCPC 122PCMK	5 .4m 2.0m 2.6m	07-20-48 03-25-52 04-21-52 04-21-52		10	1 		U H,S H H,S	SO Ed SO Ed SO Ed SO Ed SO Ed	32 33 34 35 36
	 191 186				 203 201	112PCPC 112PCPC 112PCPC 122MNKN 122MNKN	3.3m 2.4m 2.8m	04-21-52 04-21-52 04-21-52 08-18-52	 2	 25	 2	13	H H H H	SO Ed SO Ed SO Ed SO Ed SO Ed	37 38 39 40 41
	185 55 190 3	2 2 2 2			200 75 210 8	122MNKN 122MNKN 122PCMK 122MNKN 112PCPC	6 8 10 1.5m	07-26-74 06-20-75 06-28-82 04-15-87	2 1 2	10 12 10	3 3 15	5.0 12 5.0	N C,N H	SO Ed SO Ed SO Ed SO Ed SO Ed	42 43 44 45
+1- 100-	100 185	2	185	_	210	122MNKN	13	07-15-86	1	60	1	60	Н	SO Ed	47
	17.5 18.5 82 8.7	2 2	200		205 205 92	122MNKN 112PCPC 122PCMK 112PCPC	7 8 3 .6m	12-04-85 07-29-86 10-16-46 03-26-52	1 1 1	40 65 18	1 1 1.5	40 65 18	H C S H	SO Ed SO Ed SO Ee SO Ee	48 49 1 2
+2- 62-	65 88 62 91	2 2 2	91		73 104 101	122PCMK 112PCPC 122PCMK 112PCPC 122PCMK	3 1.5m 4	10-00-50 03-26-52 11-24-66 08-04-76	15	 10 40	 2 2	 2.7	H,S H,S H,S C	SO Ef SO Ef SO Ef SO Ef SO Ef	1 2 3 4 5



MAPS SHOWING WELL LOCATIONS

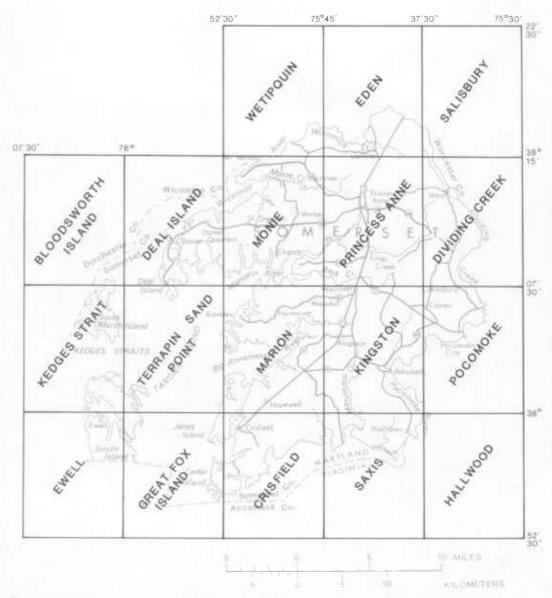


Figure 41.— Index map of 7 1/2-minute quadrangles.

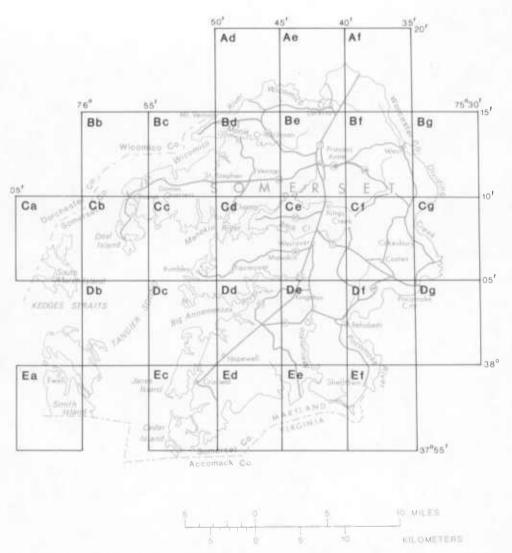
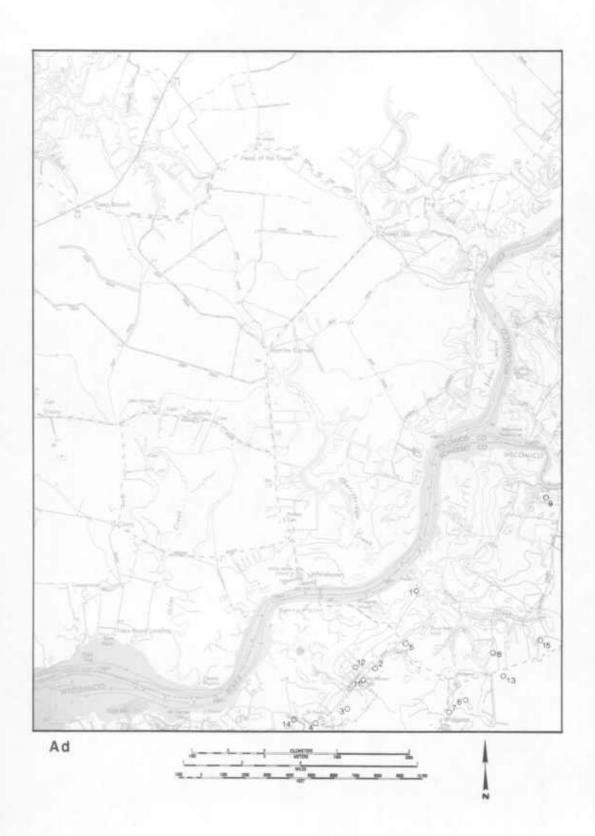
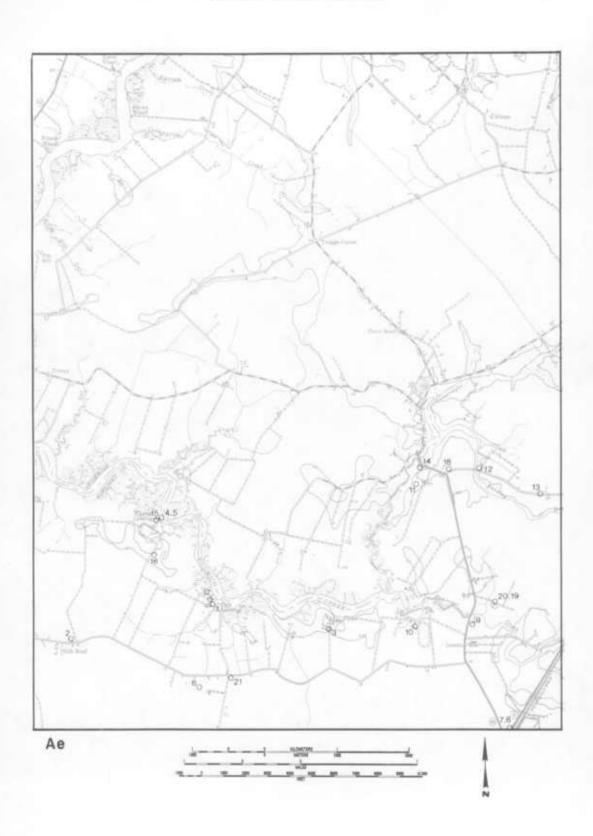
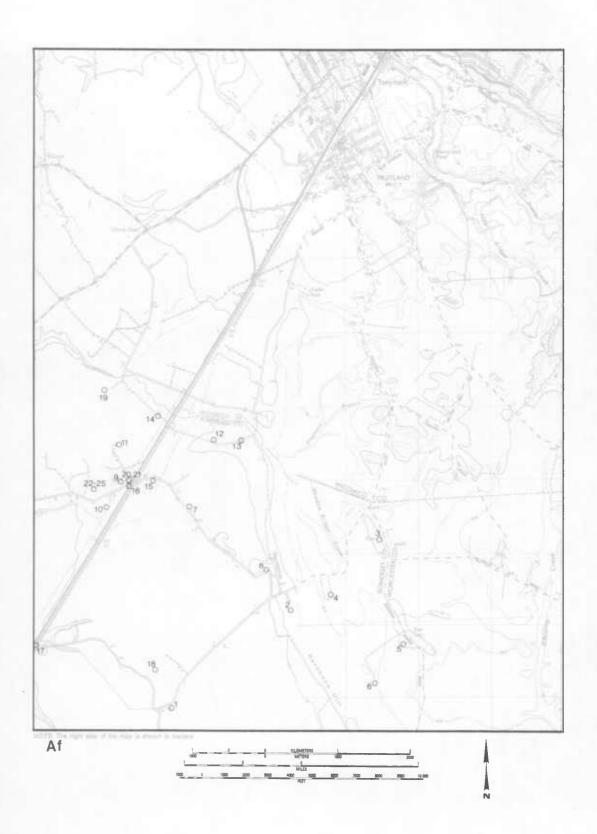
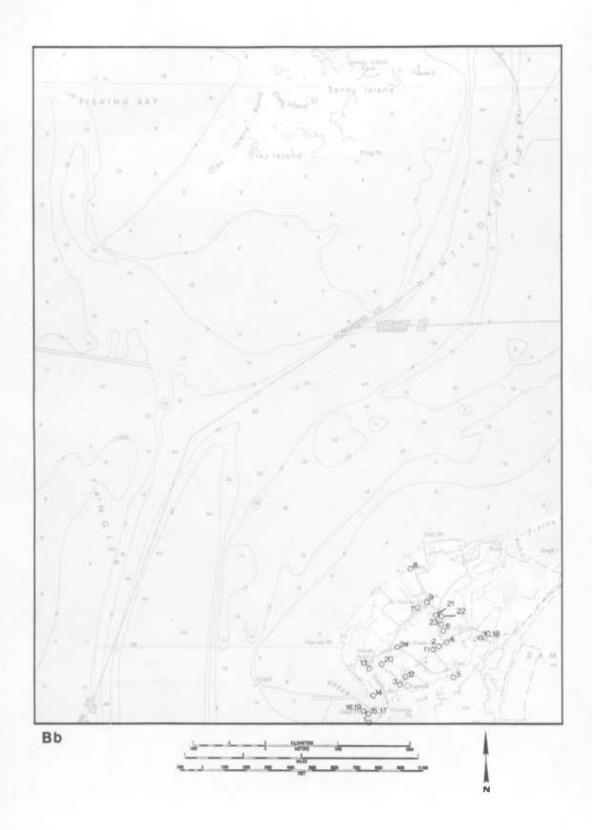


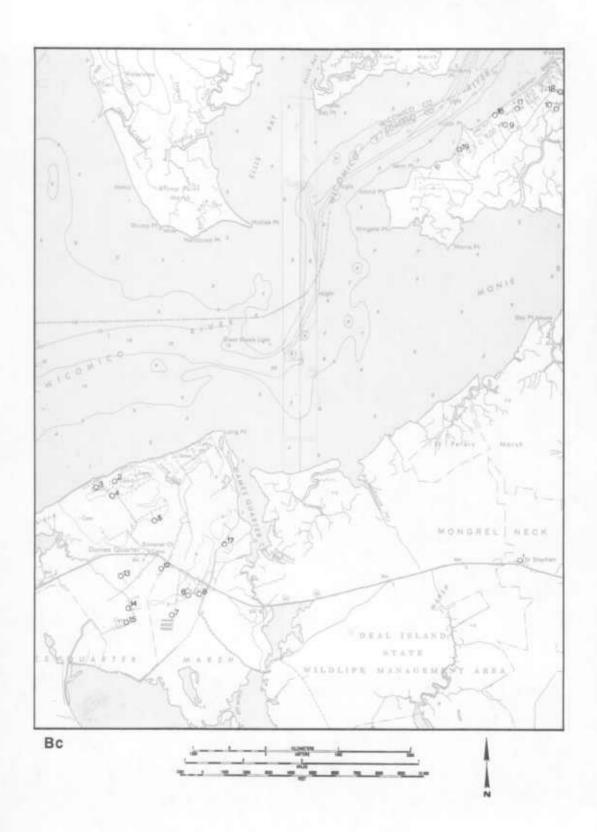
Figure 42.—Index map of 5-minute quadrangles.

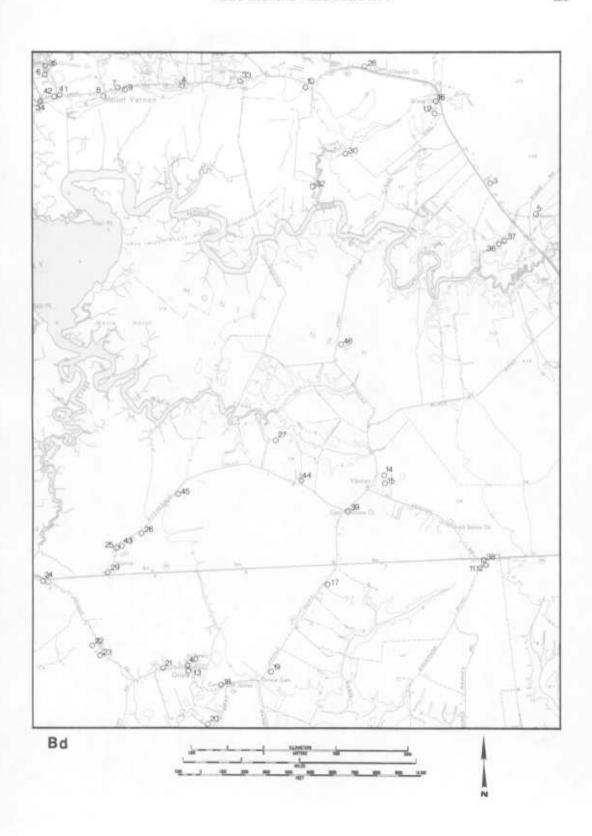


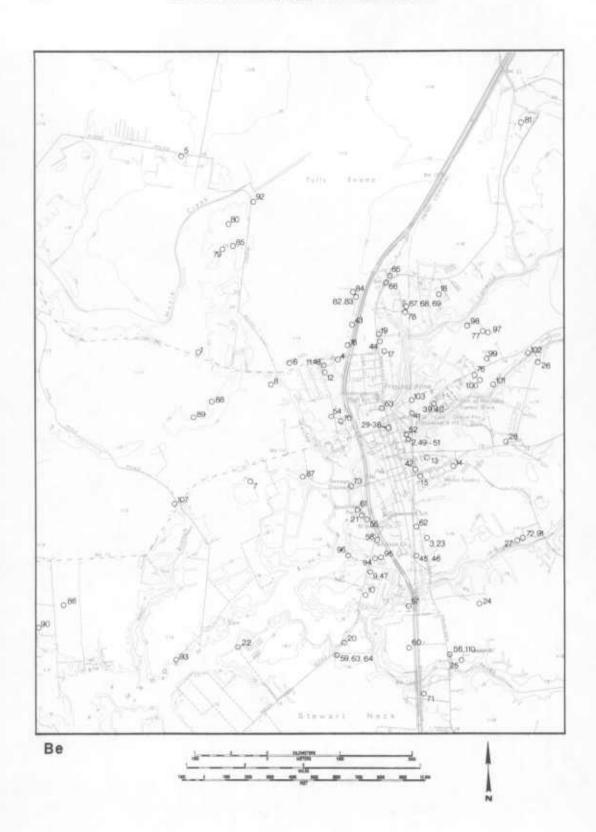




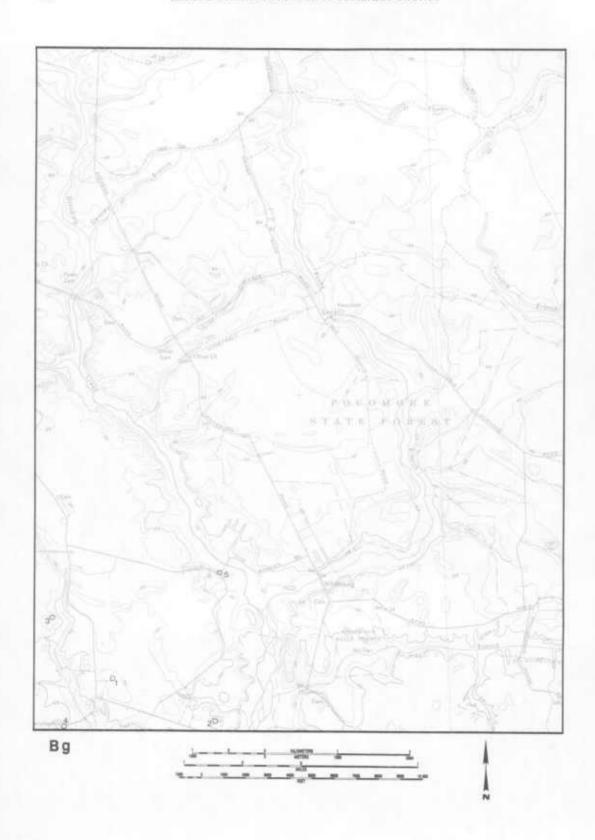


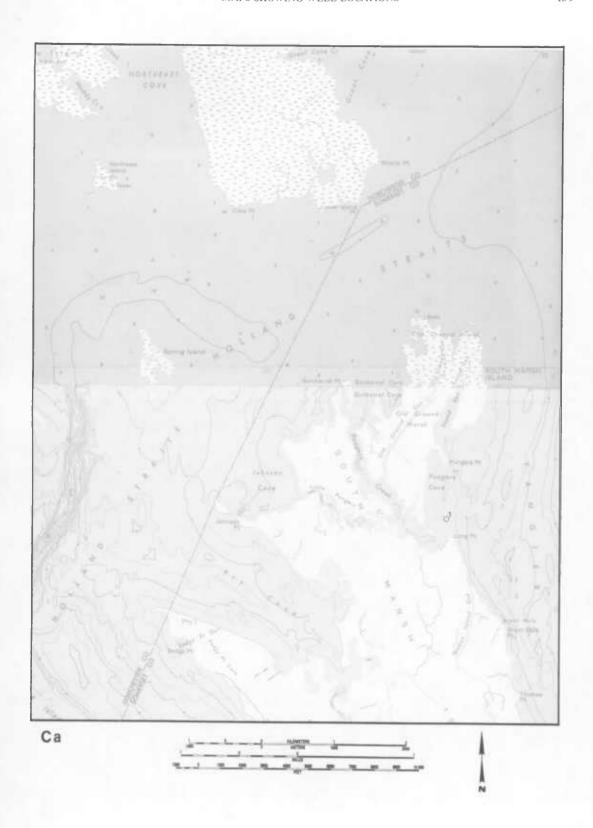


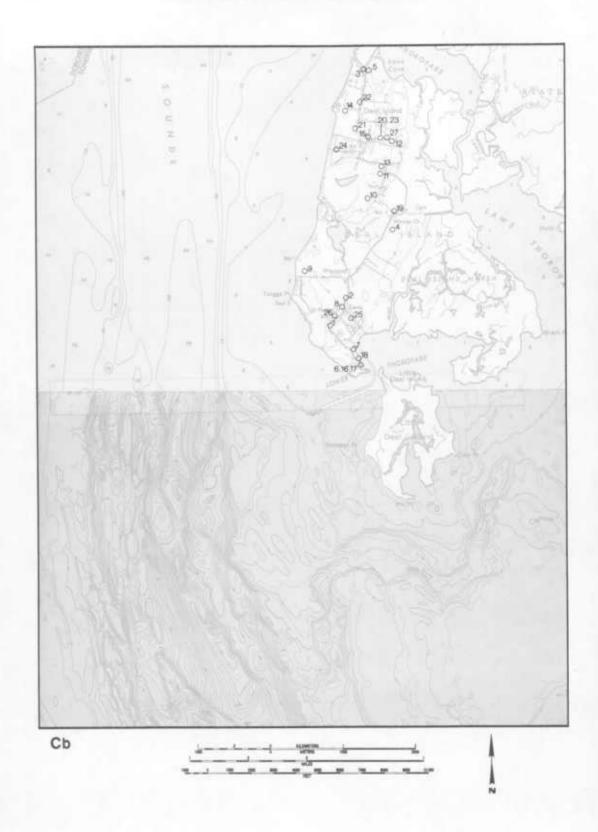






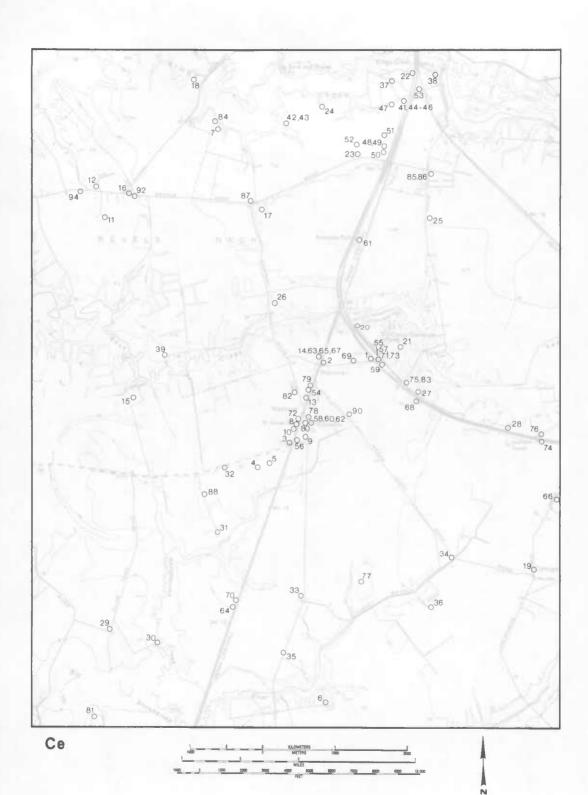




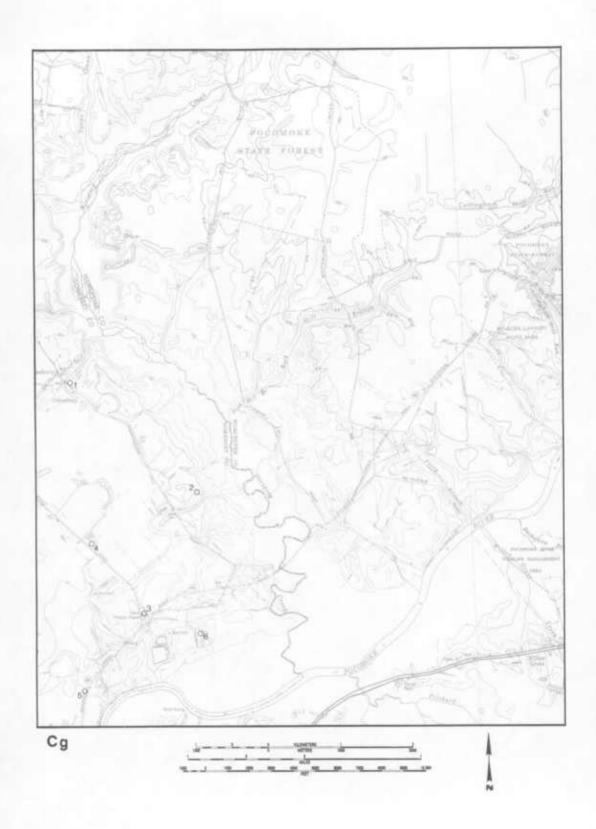












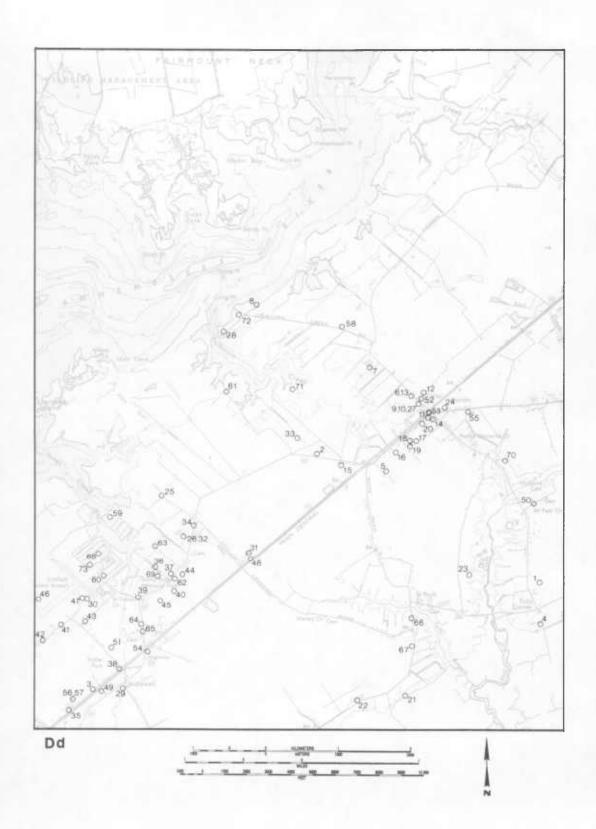


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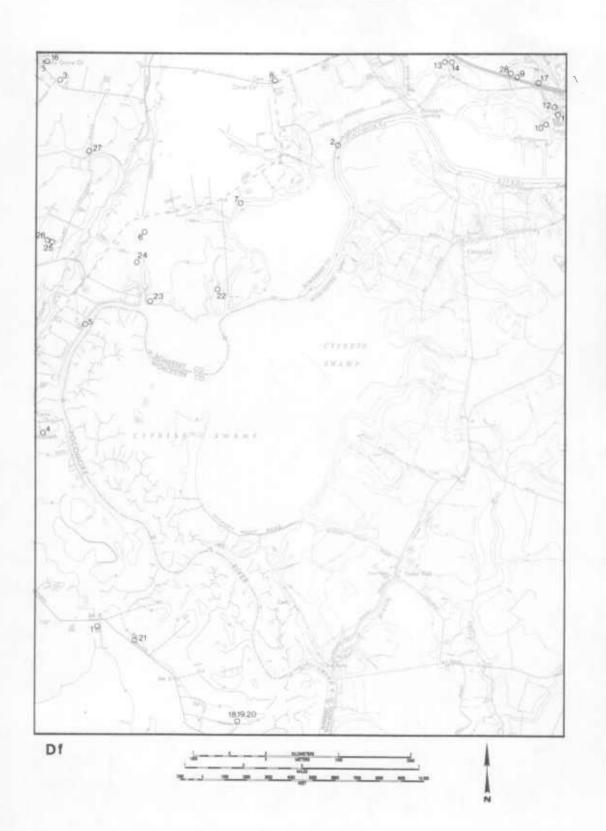






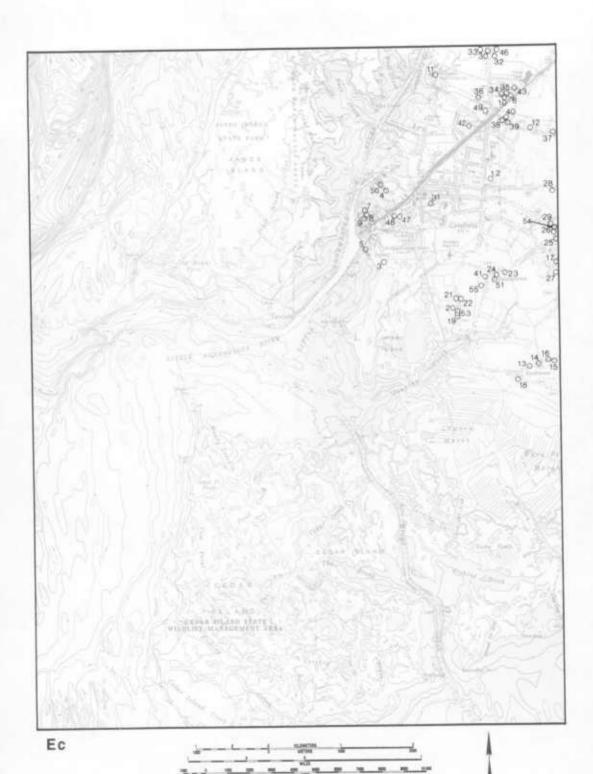




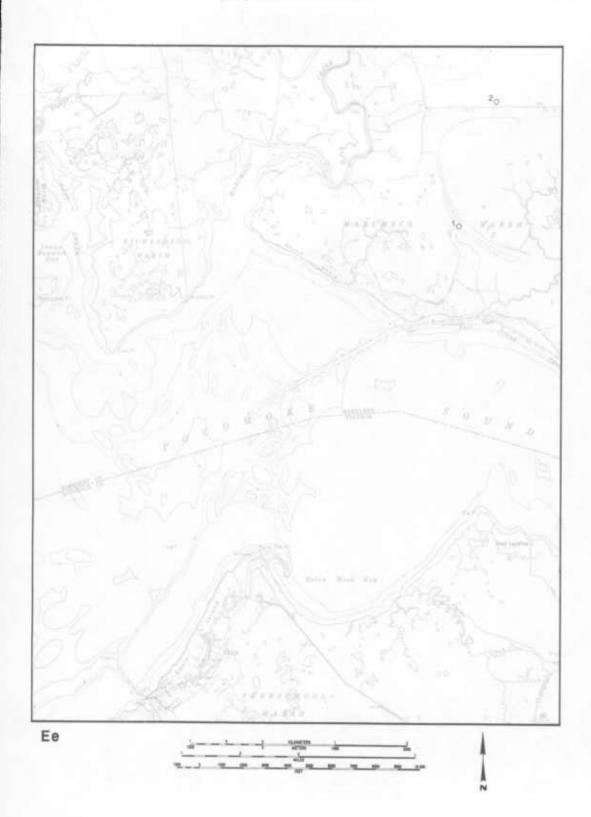


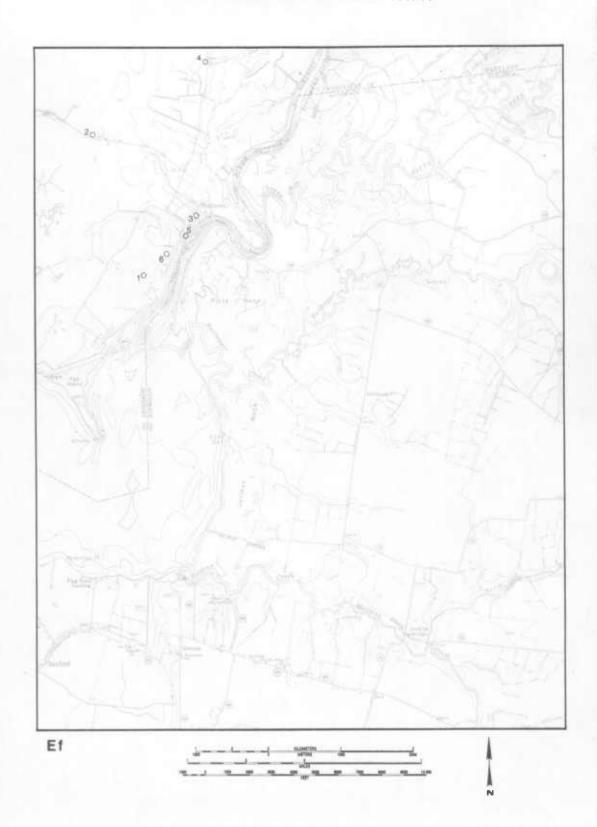


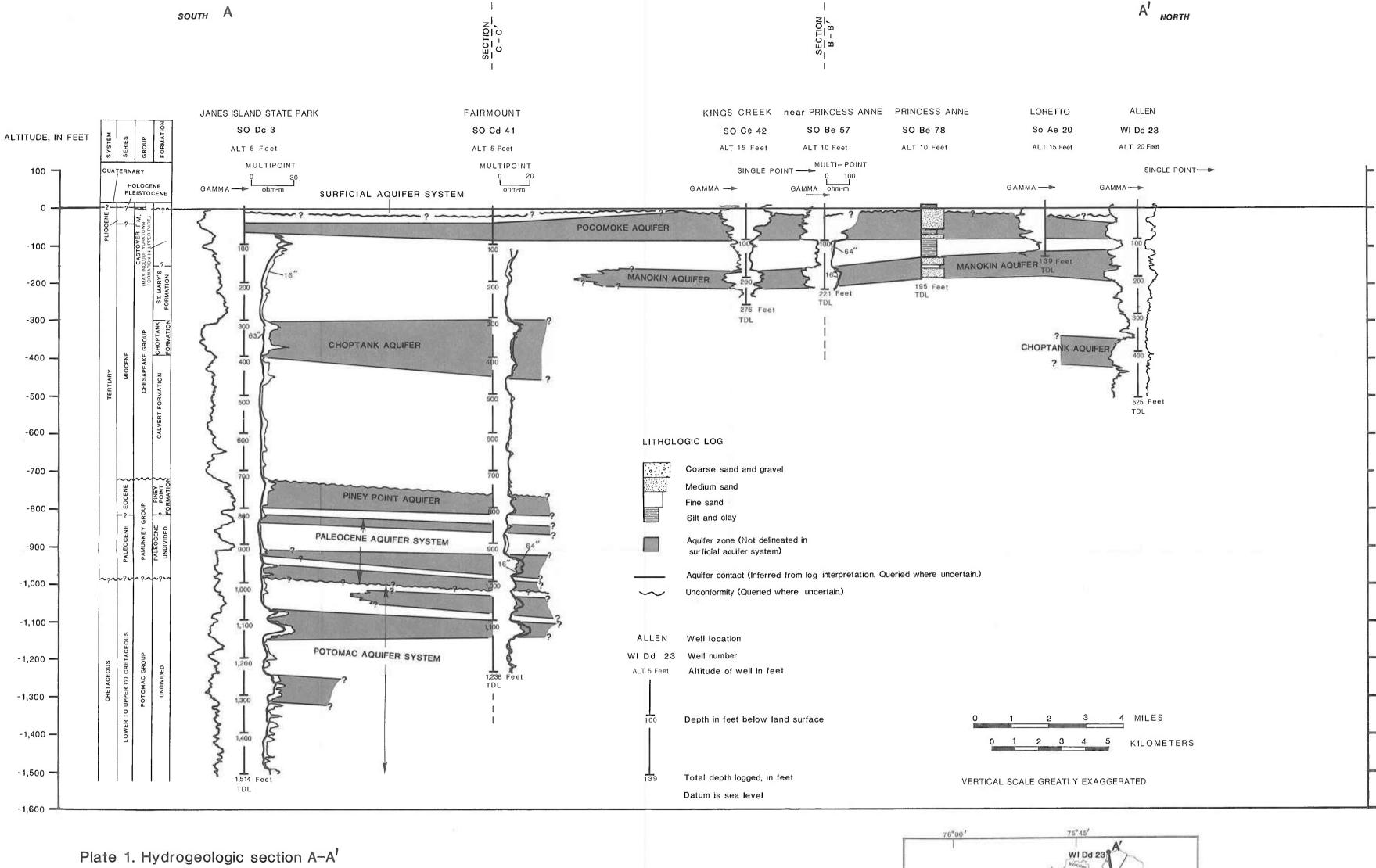


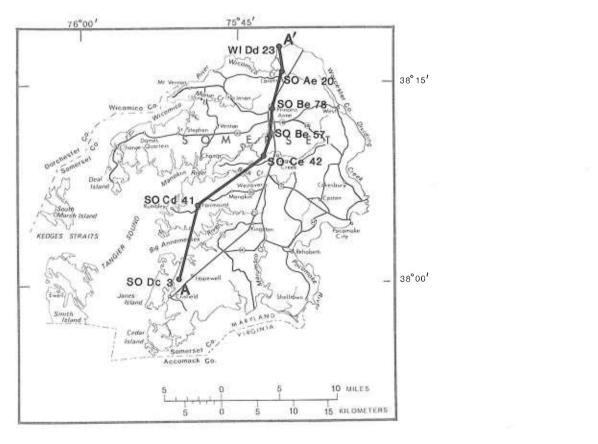


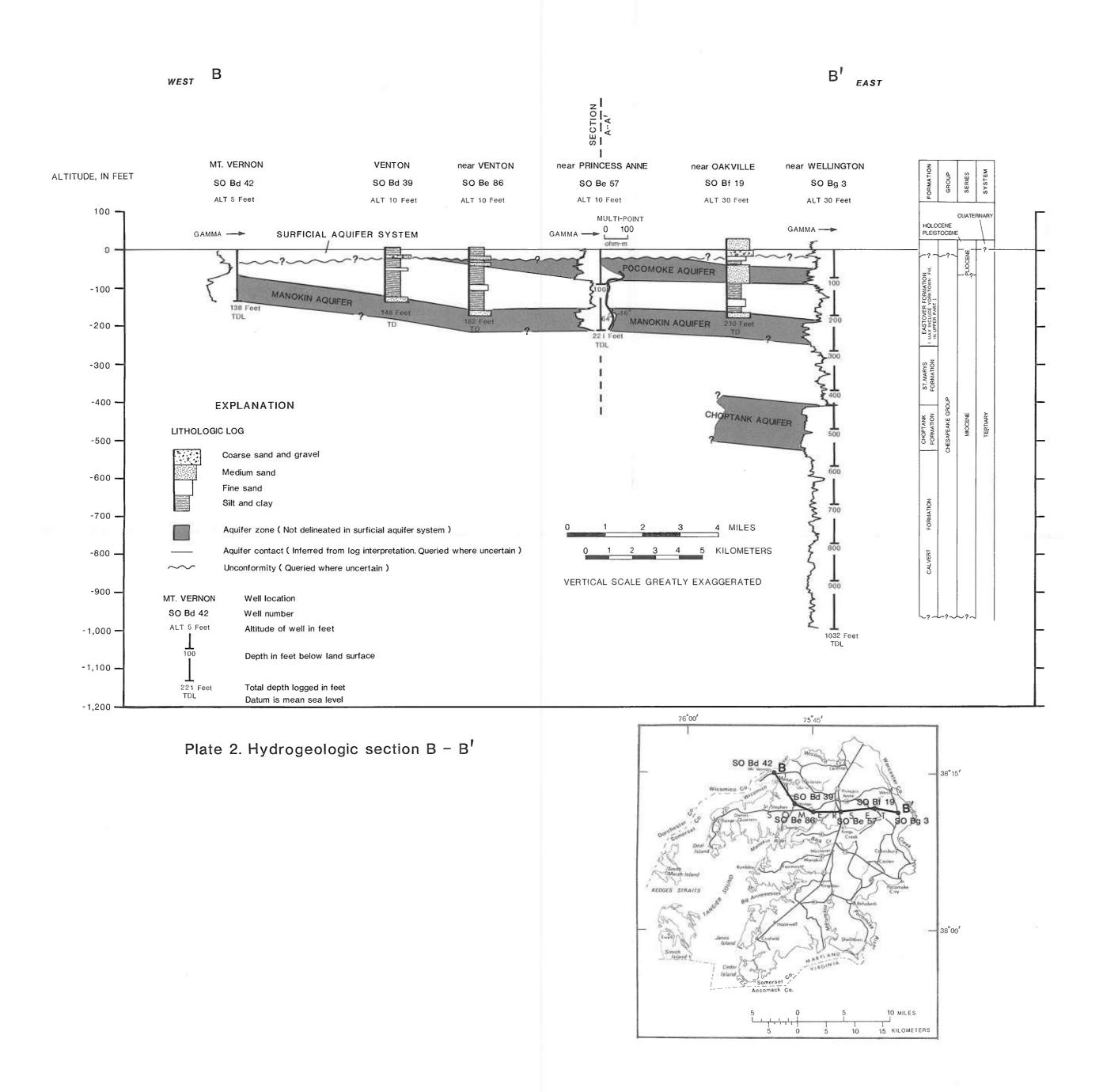


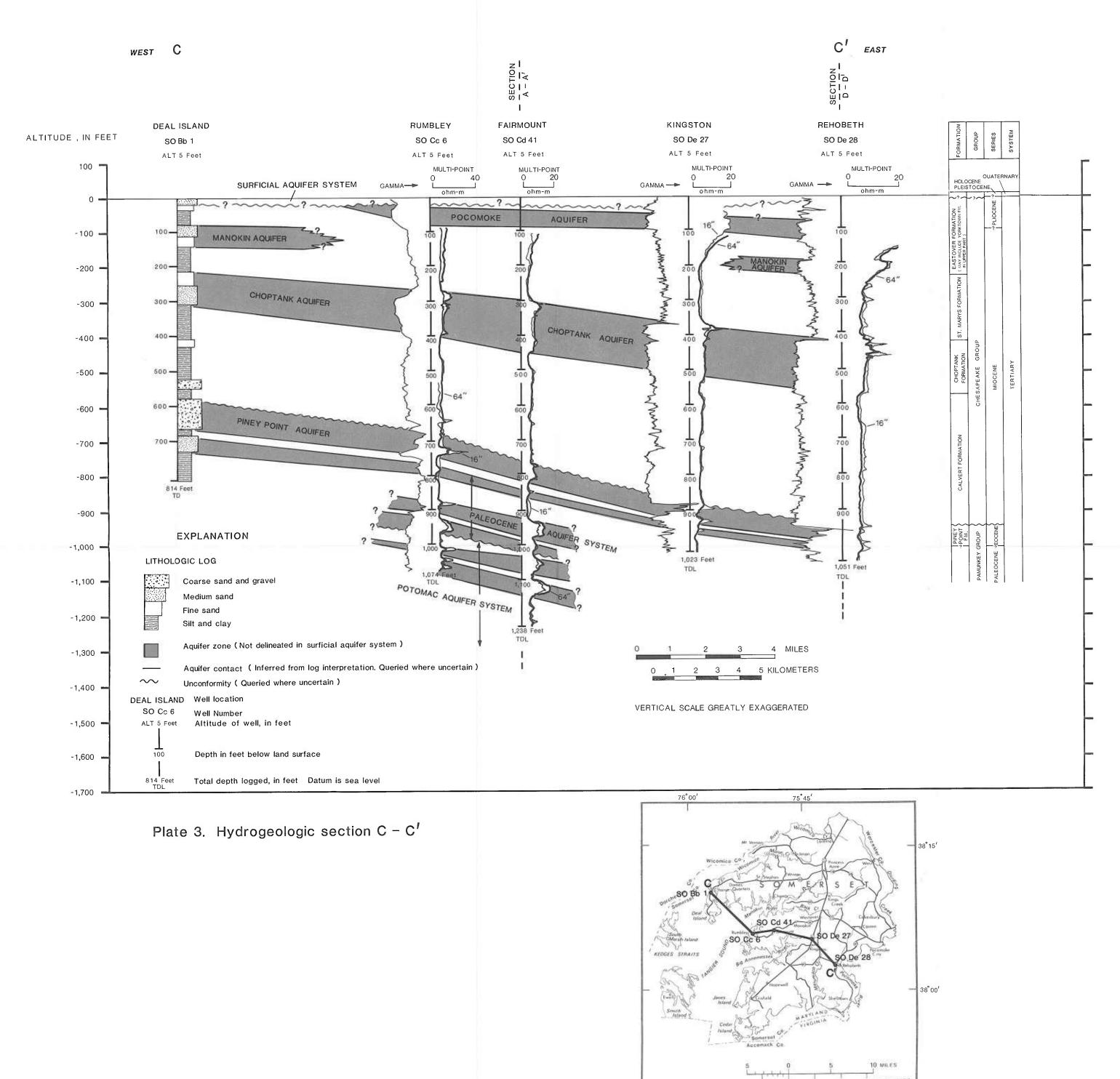


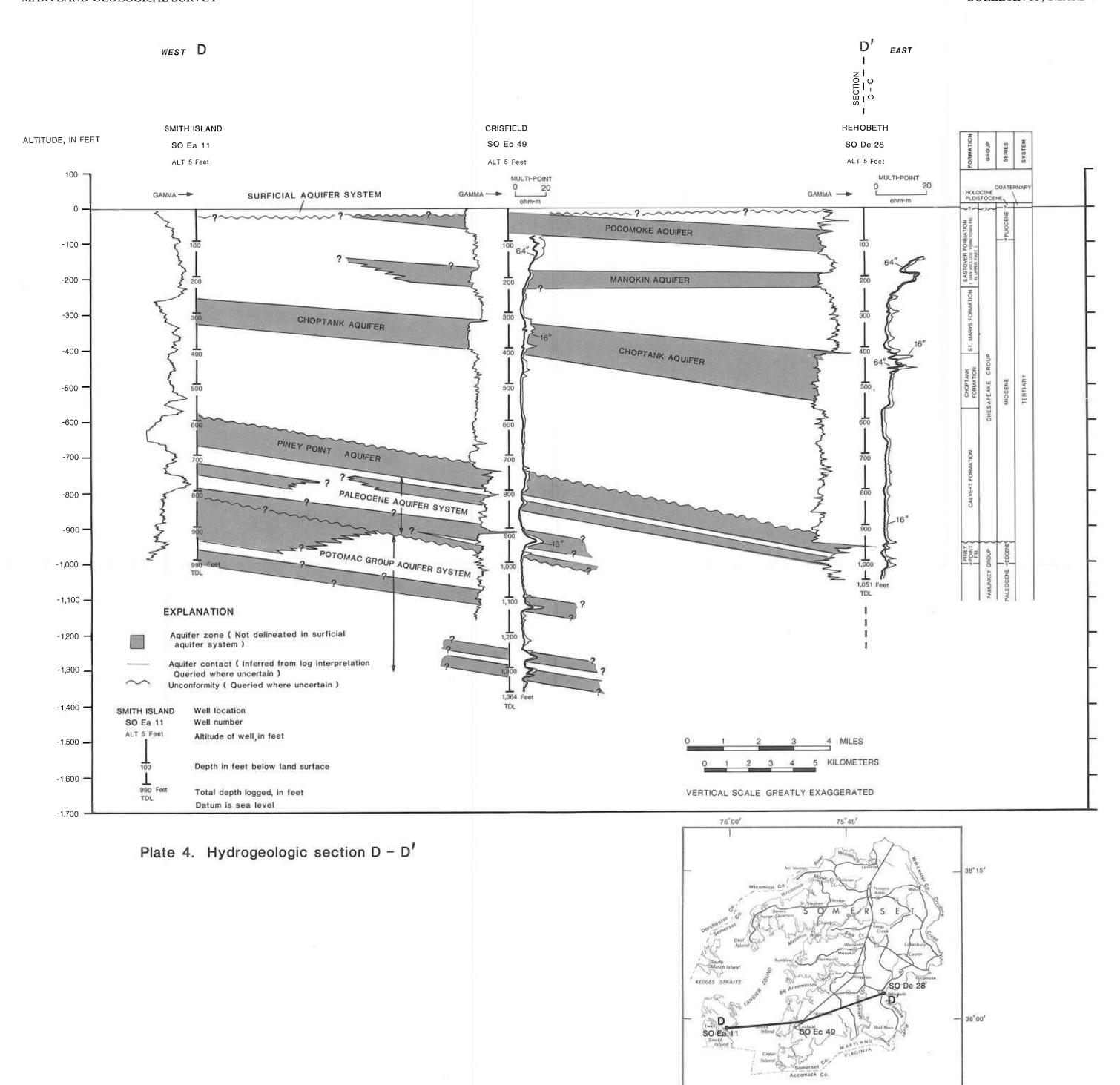












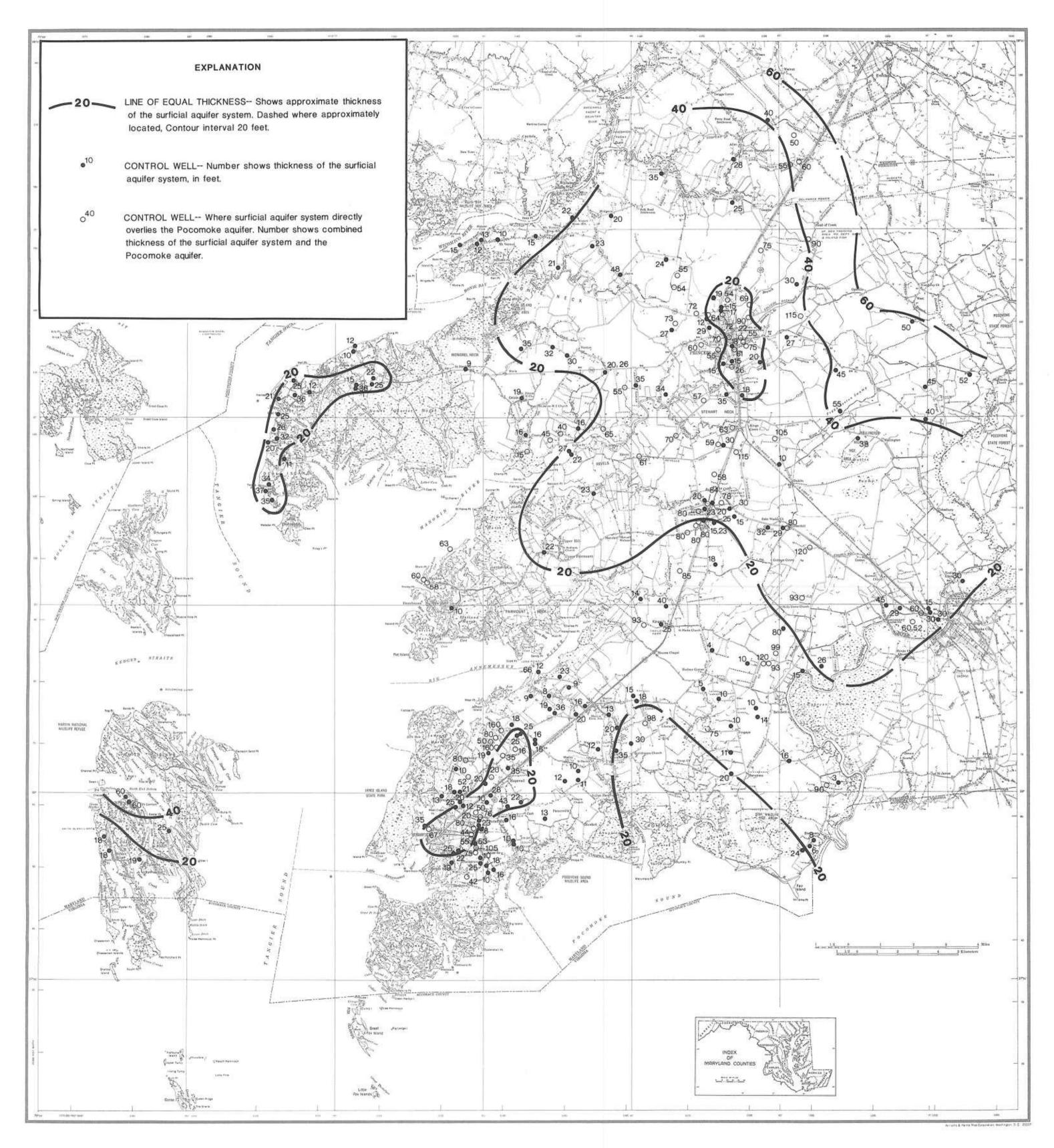


Plate 5.— Map showing thickness of the surficial aquifer system in Somerset County, Maryland.

MARYLAND GEOLOGICAL SURVEY

BULLETIN 35, PLATE 6

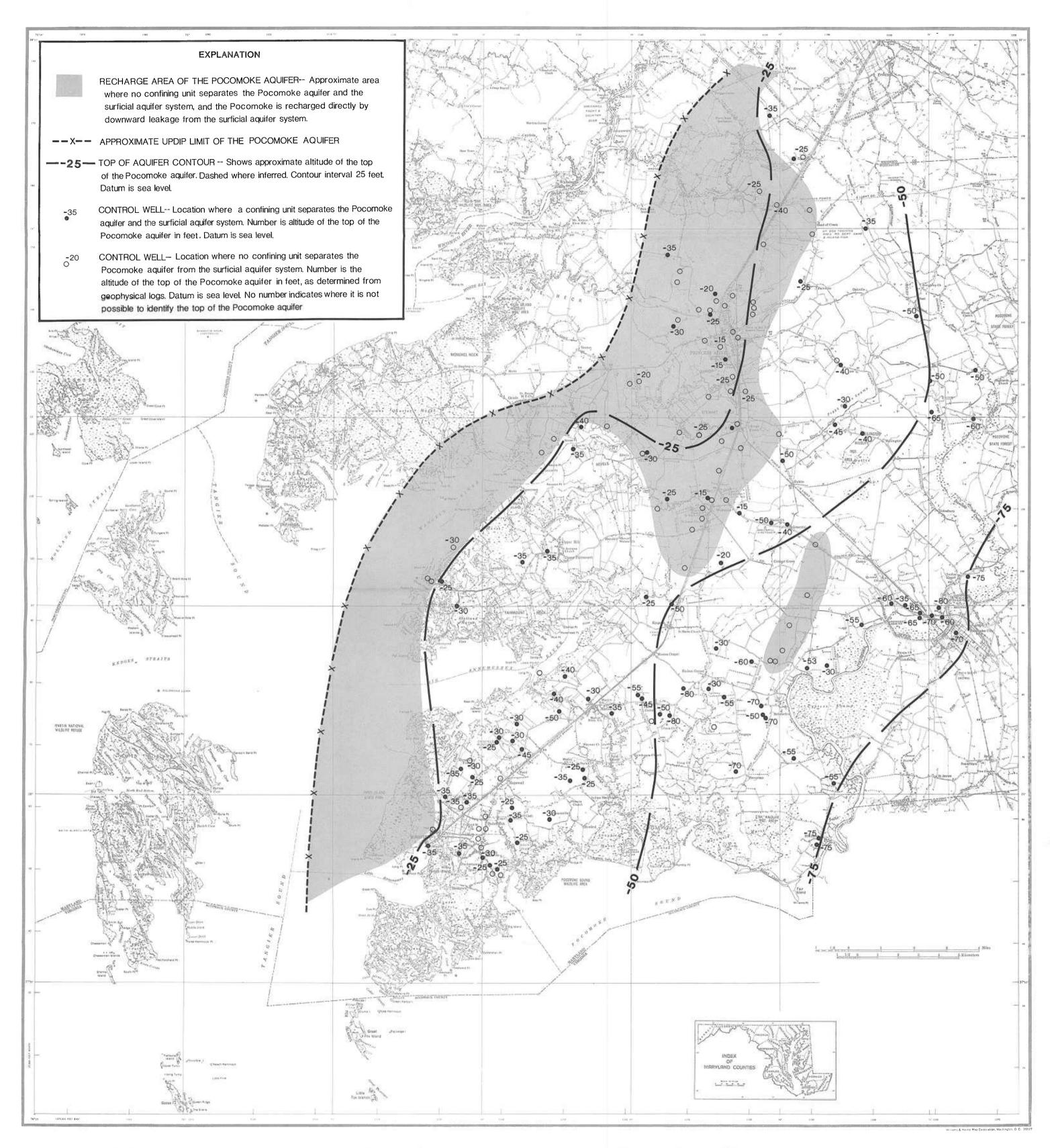


Plate 6.-- Map showing recharge areas and altitude of the top of the Pocomoke aquifer in Somerset County, Maryland.

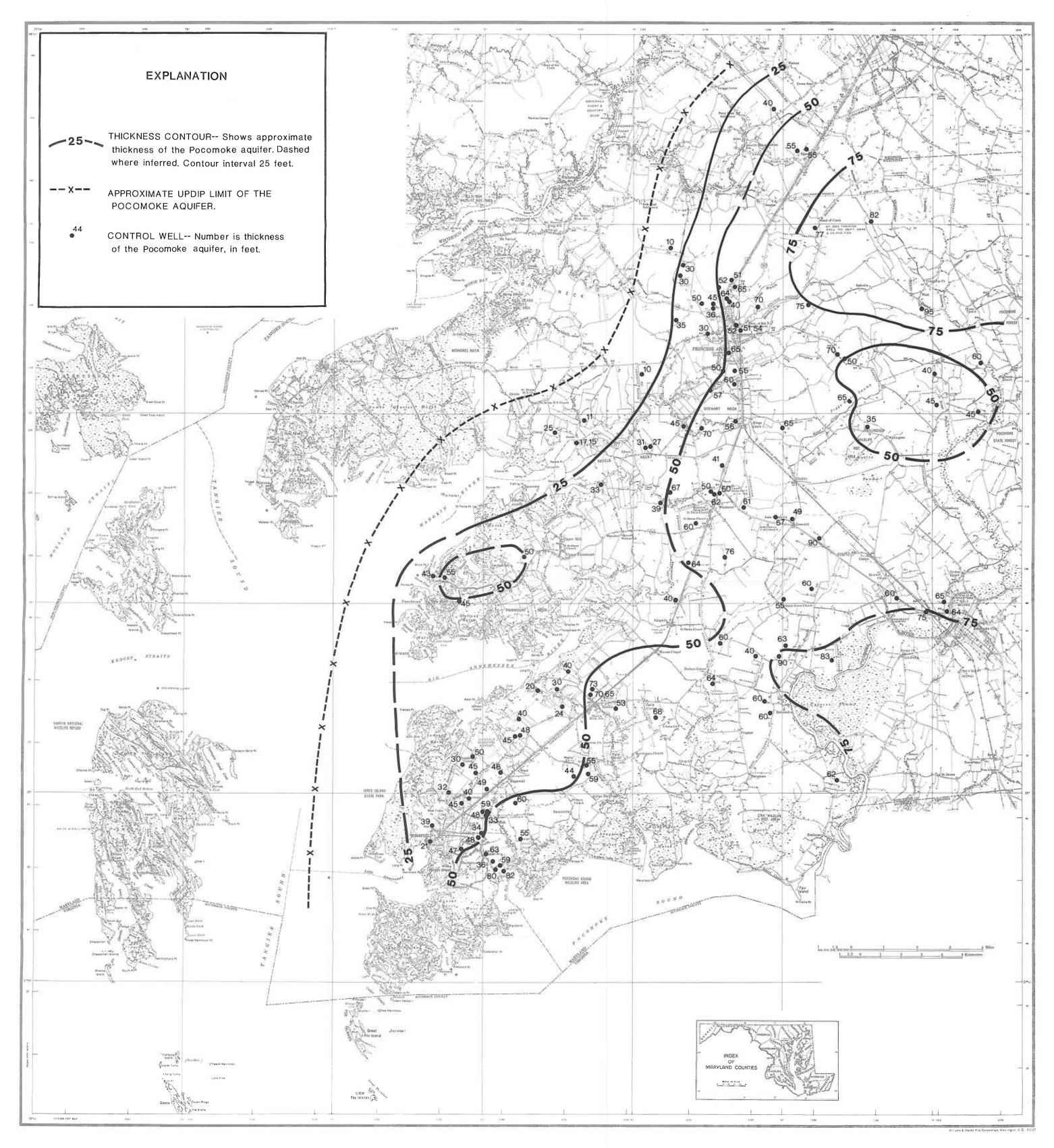


Plate 7. Map showing the thickness of the Pocomoke aquifer in Somerset County, Maryland.

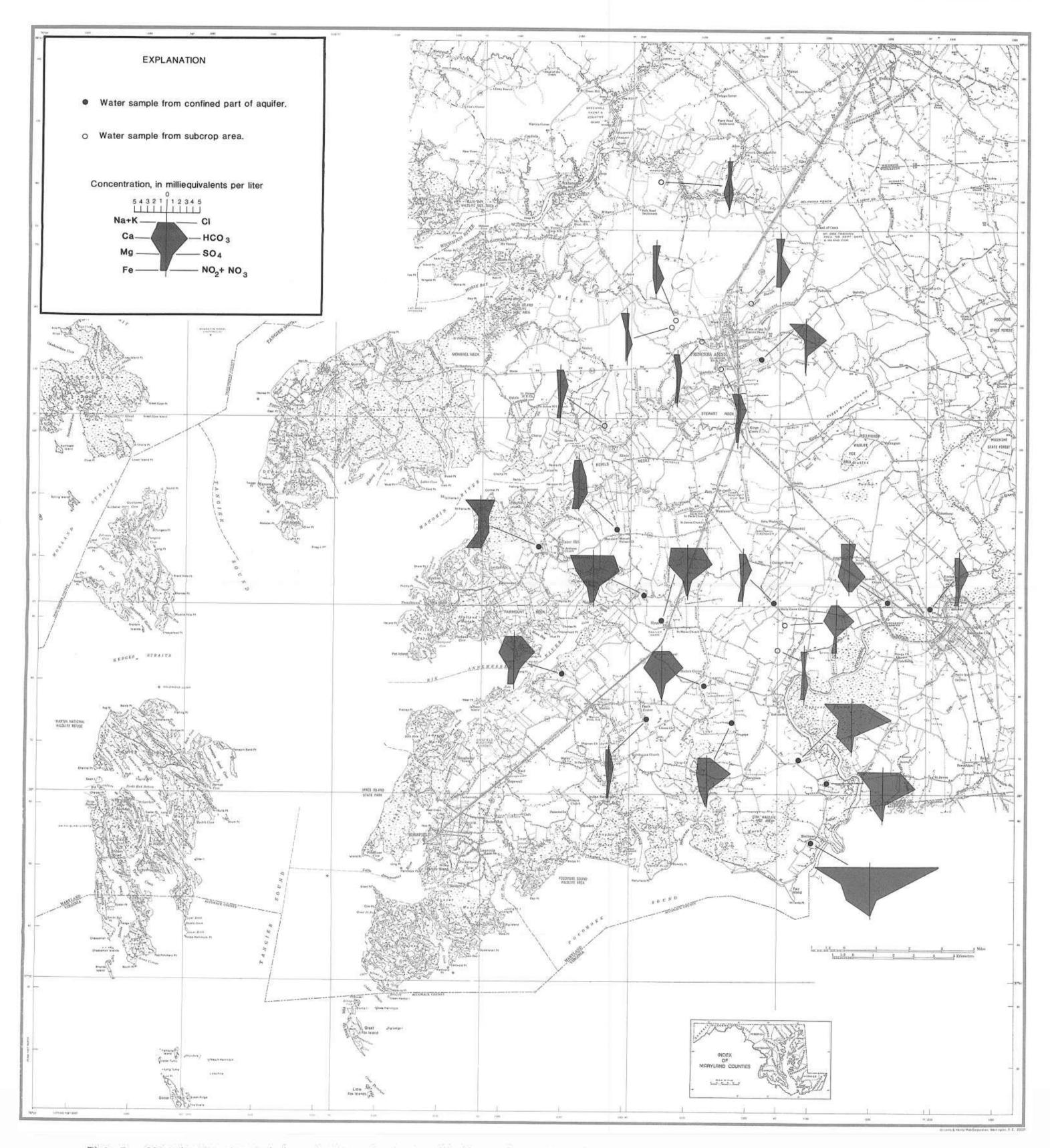


Plate 8.-- Map showing chemical characteristics of water from the Pocomoke aquifer in Somerset County, Maryland.

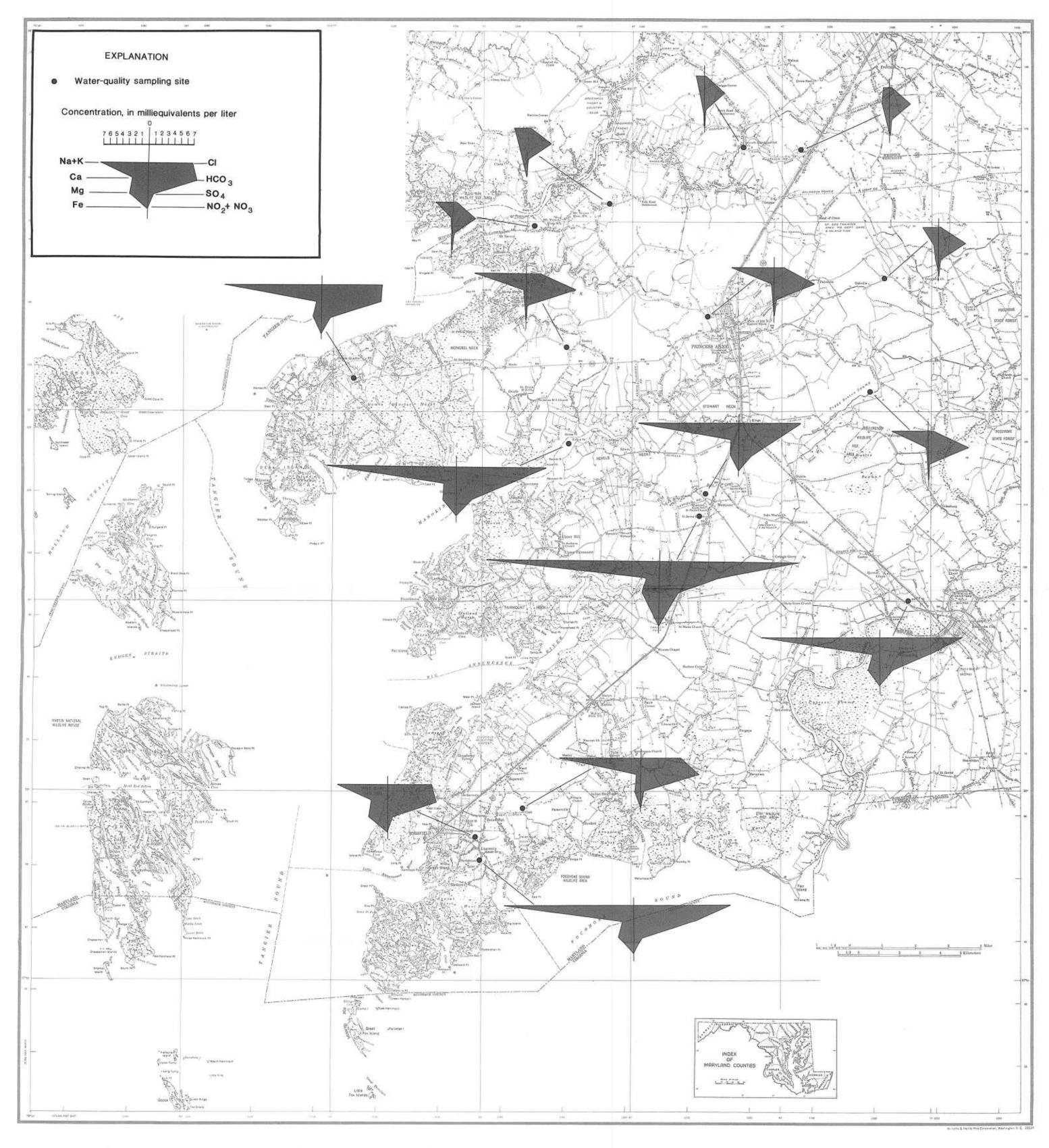


Plate 9.-- Map showing chemical characteristics of water from the Manokin aquifer in Somerset County, Maryland.